

A SIMULATION MODEL TO ASSESS THE HYDROLOGIC
PERFORMANCE OF THE TINAU WATERSHED, NEPAL

by
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DEDICATION

TO MY DEAR MOTHER, SHOBHA
WHO HAS NEVER STOPPED INSPIRING ME THROUGHOUT MY LIFE.

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ABSTRACT

A physical event-based computer methodology is presented for evaluating the hydrologic performance of the Tinau watershed in Nepal.

A stochastic model of the seasonal precipitation was developed and used to generate long-term daily synthetic rainfall data. Storm systems in the monsoon months of JUNE through SEPTEMBER are treated as frontal systems, and the storms of the remaining months of the year as independent thunderstorm events. A fitted-parameter conceptual model, called a generalized streamflow simulation system (GSSS), is used to simulate long baseflow recessions. The GSSS model transforms the synthetic rainfall data into daily streamflow using a soil moisture accounting process in a physically consistent manner.

The computer output from the model consists of statistics and cumulative density functions for monthly, seasonal and annual discharges, maximum daily high and low flows, and consecutive days of low baseflow. The model output can be used by decision makers to meet the objectives of proper land use and development.

CHAPTER ONE

INTRODUCTION

The Himalayan mountain range is considered to be largest, highest and youngest in the world. The annual south eastern monsoon produces heavy rainfall on the southern aspect of the range, which results in the washing off of productive fertile soil from the hill slopes. Nepal lies on this ecological range. Over 90 percent of the population of Nepal are engaged in agriculture, mainly subsistence agriculture. Over 350,000 people are added to the population of farmers each year. Unfortunately agricultural production has not been able to keep pace with population growth. Cultivation and animal husbandry have been expanded to unsuitable terrain and steeper slopes to meet the growing needs. The human activities place extreme pressure on forests and hill slopes already in delicate balance with the natural environment. In addition to accelerated soil loss, landslides commonly occur every year which destroy human lives, kill livestock and destroy crops. The depletion of soil from the hill lands produces sediment choked streams and river in the valleys and flood plains, causes floods during the rainy season, and droughts during the dry season.

It is estimated that on the average over 240 million cubic meters of soil is lost every year from Nepalese hills. The principal areas of this erosion are on public lands; e.g. forest, shrub and brush, and pasture lands. Land management is an urgent need in those areas, but decision makers and planners have difficulties due to the lack of accurate and up to date assessment of the water resources in that region.

Land use management and environmental protection are key issues of development strategies for the Tinau watershed. An accurate and comprehensive assessment of the present natural resource base, and its potential for development would help to optimize the benefits from agricultural land-use as well as reduce the damage to the natural environment. Because an adequate data base for making decisions in water resources development is not available, a land use management plan and policy has not yet been developed for the Tinau watershed.

Fleming (1978) found that the Tinau watershed was in the poorest hydrologic condition of any watershed in western Nepal. Winter cropping is not practised, and the valley floors remain dry and unirrigated during winter and spring. The small town, Butwal, in the plain is threated by frequent floods. Rieger et at. (1976) proposed a land management strategy for the development of water resources to increase

the productivity of the valley and for flood control downstream of the watershed.

In planning for the water resource, there is a need for a methodology that takes account the variability of rainfall, runoff and erosion from the watershed with time.

Objective

The primary objective of this study was to develop a method for evaluating the hydrologic performance of the Tinau watershed that can be used as a hydrologic tool for land management and policy making decisions.

The special objectives of the study were to develop a computer model, TINAU, that:

- 1) is simple to use and calibrate, and requires only a small amount of readily available data;
- 2) can be used as a physical event-based sediment yield model to assess the soil erosion losses from the watershed, and sedimentation on the downstream;
- 3) can be used as a hydrologic tool to estimate the probabilities of high flows and low flows in the river;
- 4) can be used as a water yield model to plan and implement a land use policy; and
- 5) is able to reflect the natural variation in the climate.

Approach

The approach of the TINAU model was to: (1) simulate a time series of daily rainfall using stochastic methods, (2) route the rainfall into a series of cascading soil moisture storage compartments, (3) drain the soil water to generate interflow and baseflow, and (4) evaluate the statistics of hydrologic performance in terms of monthly, seasonal, and annual streamflows, maximum and minimum daily discharges, and consecutive days of low flows. A simulation technique was used because of the short historical flows record available for the Tinau river, and because the most severe drought period and the greatest flow of the historical records are unlikely to be the maximum possible.

The following steps were made in determining the TINAU watershed model:

(1) Develop a physical event-based stochastic precipitation model to simulate a long-term series of daily rainfall data;

(2) Develop a fitted parameter, soil moisture accounting hydrologic model to transform the synthetic rainfall data into sequence of daily streamflow events;

(3) Produce statistical summaries and plot cumulative frequencies on monthly, seasonal and annual flows, daily maximum discharge and daily minimum discharge, and consecutive days of low baseflow recessions; and

4) Evaluate the results.

A separate submodel was developed and operated for each of steps 1 to 2 described above and are linked into the overall TINAU model. The development, calibration, and testing of each submodel is discussed in chapters three through five.

CHAPTER TWO

THE STUDY AREA

The Tinau Watershed is located in western Nepal. This study was made for the catchment of the Tinau River that lies in Palpa district. The catchment covers an area of 550 square kilometers. The area is roughly oval shaped, elongated east-west and lies between $27^{\circ} 42' N$ to $27^{\circ} 52' N$ latitude and $83^{\circ} 18' E$ - $83^{\circ} 42' E$ longitude (Fig. 1).

The Tinau River originates in the mid-hill Mahabharat range, drains the Madi valley and flows through the narrow gorge in the lower foot-hill Churia range. The main tributaries are the Jhumsa and Dovan Rivers. The Tinau River in the plain at Butwal creates flood hazards and deposits sediment on the fertile flood plains.

Topography

The topography of the watershed is characterized by its situation in the southern range of the Himalayas. The mountain range which runs east to west restricts the river course and largely determines the land-use patterns in the watershed. The watershed is in the foot-hills; the Churia hills are to the south and the middle Mahabharat range is in the north. The foot-hills rise gently from the plain (180m) to about 600m and converge abruptly into steep, almost

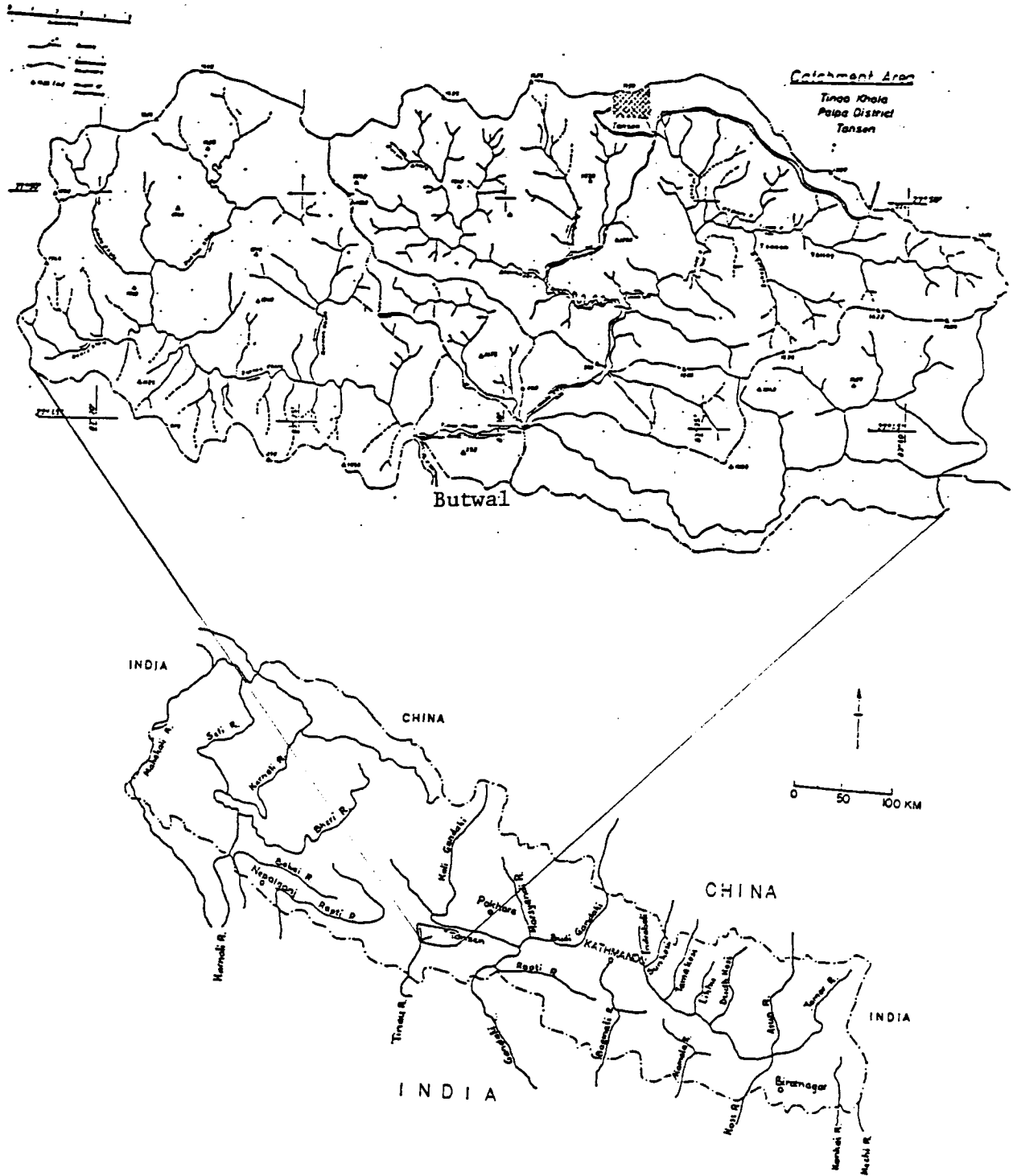


Figure 1. Map of NEPAL Showing Location of the Tinou Watershed, Palpa.

perpendicular escarpments (Churia hills). The mid-hill Mahabharat and inner Himalayan ranges comprise a rough terrain with steep slopes at altitudes of about 1800m.

Geology and Soils

The foot-hills consist of a thick detrital rocks, such as coarsely bedded sandstones, clays, conglomerates and sandy limestones. In the deep canyon, just north of Butwal, where the Tinau River cuts through the foot-hills, thick layers of sandstones vary with green-greyish to reddish clays, giving way in the north to breccious and conglomeratic layers and finally to dolomite or sandy limestones on the north boundry of the foot-hills. The mid-hills range consists mainly of quartzites, phyllites and sandstones with intermediate layers of schists. On the southern border of the Madi valley basin, a small strip of soft, pinkish schists extends from east to west.

However, the soils of the study area, developed from different parent materials, have some common characteristics. The soils vary from alluvial in the plains, foot-hills and mountain valleys; to laterite and lateritic soils in the hills, mountain terraces, hilltops and forested hills; and podzols under pine forest vegetation specially in the north-western sector. Hill soils are poorly developed (Pradhan et al. 1967), do not hold water well, and most of top soil has been washed downhill, especially on southern slopes.

Dominant soil texture is moderately to well drained sandy loam, underlaid by sandy loam, sandy clay loam or clay loam. Soil reaction is slightly alkaline to strongly acidic, and with low to high organic matter content.

The soil types of the north-eastern part, specially south facing slopes, the Madi plain and the upper Tinau valley are sandy clay loam to clay loam, slightly to strongly acidic in reaction. Those in the bottom lands are deep, stonefree loams. In the Dovan valley, lower Tinau valley sandy loam dominates. These soils are strongly acidic to moderately alkaline, medium to high in organic matter content, moderately well to well drained sandy loam.

Climate and Vegetation

Two major climatic regions are recognized in the study area: the hot and humid area of the plains and foot-hills; and the cooler or microthermal region of the Mahabharat range. The plains and foot-hills area below 900 m is a land of humid sub-tropical monsoon climate. The intricate seasonal monsoon rainfall associated with southwest winds of the summer monsoon and practically rainless winters extend into the mountain valleys. In the northern region, the climate changes markedly with elevation. The Mahabharat range and the inner Hamalayan chains have micro-thermal climate. Here, the winters range from moderately cool to severe, with warm and wet summers.

The rainfall of the area varies from 1400 mm to 2000 mm per year. Most of the rainfall is associated with the south west monsoon which breaks about mid June and spreads over the entire watershed within a few days. The south west monsoon usually lasts from June to September, and within this period the watershed area receives approximately 80 percent of the rainfall. Long-term annual rainfall distribution shows the following pattern: 6 months (November to April) are either very dry or have a rainfall below 50mm per month; 2 months (May and October) normally recorded below 75mm of rain; and 4 months (June to September) receive 200mm or more per month.

The physiographic range of the study area, which varies from 250 m to nearly 2000 m above sea level, creates a variation in temperature. The summers are hot with temperatures exceeding 38°C, and winters are cool and pleasant with a mean temperature of 12°C. The subtropical type climate prevails above 900 m where summers are cool and pleasant with a maximum temperature of about 32°C and winters are cold with a mean temperature at 10°C with occasional frost. Fairly dense moist deciduous forest occurs in the southern part of the foot-hills and Churia hills which lies in the tropical climatic region.

The main forest associated with this climatic region is the Shorea/Terminalia type. The tree species in this

forest are Shorea robusta, Terminalia spp., Acacia catechu, Salmalia malabarica, Dalbergia sisoo, Bauhinia vahlii, Mallotus philippinensis. Chirpine grows in the driest part of the Churia hills. Grass cover associated with this forest type is called Saccharum/Phragmites/Imperata (Dutt,1979). The main grass species are Saccharum spontaneum, Phragmatis karka, Imperata cylindrica, Cynadon dactylon, Heteropogon contortus, Eulaliopsri binata.

In the subtropical Mahabharat range two associations, the Shorea/Schima type, and the remnant of Schima/Castanopsis type are dominant. Dutt(1979) identified these two types in the Tansen area. The first type concentrates in an elevation range of 600 - 1200 m where Shorea robusta is the dominant species with Schima wallichii a companion species. This dominance of Shorea is related to edaphic factors, as the soil in this area is shallow and the floor composed of mainly rocky outcrops.

The Schima/Castanopsis type occurs at elevations of 1100 - 1500 m. Usually northern and north-eastern slopes are clad in dense forest and the southern slopes are bare, being too steep to maintain a soil cap for the growth of the forest or to allow the winter snow to accumulate. Main tree species in this forest type are Schima wallichii, Castanopsis indica, Fraxinus floribunda, Myrica nagii, Rhododendron arboreum, Englehardtia spicata, Lyonia ovalifolia, Prunus cerasoides, Morus spp., Brassicaopsis hainl, Ficus

religiosa, F. bengalensis, F. lacor, F. semicordata, F. glabberimma, F. clavata, Bauhinia variegata, Pyrus paschia. Pine (Pinus exelesa, P. longifolia) predominates in the north-western area. Shrubs that are dominant in this area are Maesia shista, Berberis aristada, Rubus ellipticus, Ziziphus spp. This area is rich in herbs and spices. The main spices and herbs are Osbeckia stellata, Artemisia vulgaris, Urtica dioica. Grammineae play an important role in soil conservation in this forest type. Two main species found are Dendrocalamus hamiltonii, Arundinaria spp. The grass cover associated with this forest type is called Themeda/Arundinella cover. The grass species in this association are Themeda anthera, Arundinella bengalensis, A. nepalensis, Chrysopogon fulvus, C. gryllus, Bothriochloa intermeia, B. pertusa, Heteropogon contortus, Dimeria fuscescens.

Agriculture and Land-use

The southern part of the watershed Churia hills is forested hilly land, where a little agriculture is practiced. The rest of the area is in the central hill farming system. The Tinau Watershed Project (TWP, 1980) has prepared a land-use map of the study area. The map indicates two predominant land-use types: south and central portion: forest-covered and erosion scared, no intensive agriculture; and northern section: more intensively cultivated, forest

cover reduced, shrubs and bushes dominate. In 1980 TWP had estimated the watershed area at 522.4 sq. km, which includes 197.2 sq. km of forest land, 127.4 sq. km of grass land, 168.17 sq. km of agricultural land with mainly rain-fed cultivation, and 12.78 sq. km of erosion, landslides, rivers and ponds.

Agriculture is the basis of life in the study area. Most of the population is engaged in agriculture, so almost every household has a piece of land. The main crops are rice, wheat, pulses, etc, in the lower valley bottom, and corn, millet, barley, sorghum, dry land rice, black gram, soyabean, mustard, pea, sunhemp, potato, ginger etc on the hill slopes and terraces. Most of the cultivation is rain-fed, except for a few irrigated areas in the Madi plain, Jhumsa, and Dovan valleys.

Livestock-keeping is an important farming system next to crop production. Cattle are raised mainly for ploughing and farm manures. Water buffaloes and bullocks are mostly stall-fed. During the wet season, animals are fed mainly fodder from trees, and grasses from private lands. In the dry season animals have to live on straw from different crops, and from grazing on pastures. TWP(1980) had estimated 6000 head of cattle and 2500 goats in the watershed, Grazing is uncontrolled and year round. The resulting overgrazing produces severe erosion during the monsoon season.

The human population of the watershed was 71,000 in 1980, and the rate of population growth was 2 percent. It was estimated that there were 11,141 households in 1978 (TWP, 1980), and 6.1 persons per family. In 1971, about 6500 people lived in and around one small hill town, Tansen, alone. Seasonal migration of the people prevails due to ecological and economic conditions. Seasonal movements take place twice a year, from July to September in the monsoon, and from December to February in the winter.

Data and information for this study area were collected from different agencies in Nepal. Daily rainfall, maximum-minimum temperatures for Tansen and Butwal were obtained from the Nepal Meteorological Service, and stream flow records of the Tinau river at Jhumsa and Butwal from the Hydrology Section, Ministry of Water Resources. Soil Sciences and Agricultural Chemistry, Department of Agriculture provided the soil survey report of Palpa district.

CHAPTER THREE

THE PRECIPITATION MODEL

Precipitation is a key process in most hydrologic models. Because of its variable nature in both time and space, it is a complex process. The stochastic model used in this study is an event-based daily precipitation simulation model. The model takes account of two seasonal precipitation patterns that occur in much of Nepal.

Duckstein, Fogel, and Kisiel (1972) and Duckstein, Fogel, and Davis (1975) developed stochastic models for rainfall patterns prevalent in the south western United States: convective type summer thunderstorms, and frontal type winter storms. These models were adapted to the two major rainfall patterns in south western Nepal: Dry Weather Precipitation and Monsoon Precipitation. Two assumptions were made: hot summer and cold winter thunderstorms occur as independent events; and wet warm storms occur in groups or sequences.

Initial Data Analysis

Daily precipitation records were available for the periods 1978 through 1985 in the Tansen area and 1972 through 1982 in the Butwal area. Similar daily precipitation

records were available from two main valleys, the Jhumsa and the Madi plains for 1984 and 1985. A comparison of precipitation amounts at Tansen was made with the other three areas, and it was observed that a correction was needed for spatial distribution of rain storms. The annual rainfall measured at Butwal was greater than that measured at Tansen by a factor of 1.504. It was calculated that the gauge catch at Butwal represented about 15 percent of the study area. Therefore, the daily rainfall measured at Tansen (T) was corrected by

$$\text{Rainfall} = 0.85 * T + (1.504) * (0.15) * T \quad (1)$$

For the calibration of the model parameters, daily precipitation data were analyzed by using computer programs &TN1 and &TN2. These two programs, originally written by Jones (1981), were modified for this study. Computer program &TN1 was written to determine the actual precipitation event interarrival day probability density function (PDF) and cumulative density function (CDF) of precipitation events on a monthly basis. Computer program &TN2 was written to determine the actual precipitation amount per event PDF and CDF. These programs also list a sequence of statistical parameters and 't' values, and a corresponding degree of freedom for all possible pairing of months. A student's t-test was conducted at the 0.05 level of significance for

grouping of statistically like months to separate seasonal precipitation types, hot and cold dry and wet monsoon.

After completing the t-tests and grouping months, JAN-FEB-MAR, APR-MAY, OCT, and NOV-DEC groups were found to be characterized by the hot and cold dry type seasonal precipitation. Computer programs &TN3 and &TN4 were developed to analyze these groups for estimating the parameters needed to fit the data to theoretical distributions. The months of JUNE through SEPTEMBER were treated separately as a wet monsoon precipitation type. A computer program &MONSOON was used to analyze and describe the parameters used in the wet monsoon seasonal model.

Dry Weather Model

An event-based stochastic model of convective storms (Fogel et al. 1971, and Duckstein et al. 1972) was used as the basis for simulating the hot and cold dry weather precipitation storms model. Two basic assumptions are made in this model:

- (1) storms occur independently of each other, and
- (2) only one precipitation event occurs in a day.

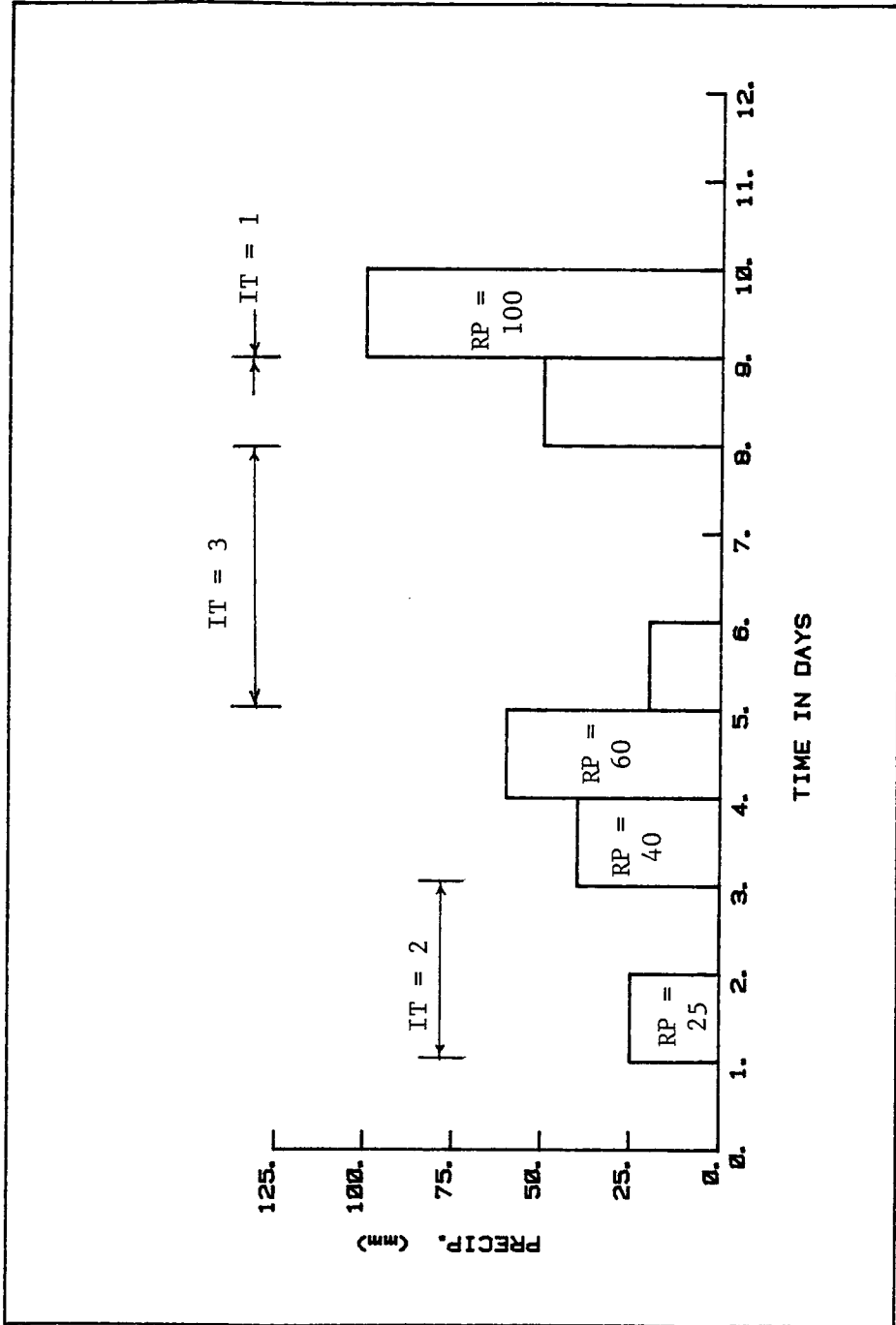
The two random variables used in simulating a sequence of dry weather precipitation events were the amount of precipitation per event (RP) and the interarrival time between rain events (IT). An event is defined to be a rainy day having a total precipitation of at least 0.1 mm. Two

consecutive precipitation events were defined to have one day interarrival time, and two precipitation events separated by two dry days to have a three day interarrival time.

A graphical representation of these variables is shown in Figure 2. Empirical frequency distributions were developed for these two variables, and each of the variables was fitted to a theoretical probability distribution. The Kolmogorov-Smirnov (K-S) test was used to indicate the best-fit between the frequency distribution of actual data and theoretical distributions (Benjamin and Cornell, 1970) at the 0.05 significance level.

Interarrival distribution groups 1 and 2 for the months of JAN-FEB-MAR and APR-MAY indicated a significant difference at the 0.05 level in the K-S test from the general theoretical distributions, and thus did not fit the theoretical distributions. In these situations, the actual distributions were shifted by assigning the first interarrival day PDF in the actual distribution to the first interarrival day (Jones, 1981). The shifted PDF, CDF were evaluated for all days beyond the first day. The new set of statistical parameters developed were used to generate new theoretical distributions which were tested against actual distribution using the K-S test as before.

The two parameter gamma distribution was found to fit actual distributions of precipitation amount and



IT = INTERARRIVAL TIME BETWEEN EVENTS (DAYS), RP = PRECIP. AMOUNT PER EVENT (MM). AN "EVENT" IS DEFINED TO BE A DAILY RAINFALL AMOUNT GREATER THAN 0.1 MM.

Figure 2. Hot and Cold Dry Weather Precipitation Model Components.

precipitation event interarrival time for three groups. The PDF for the gamma distribution is:

$$f_x(x) = \frac{\lambda^k x^{k-1} e^{-\lambda x}}{\Gamma(k)} \quad (2)$$

where, $\lambda = \bar{x}/s^2$ and $k = \bar{x}^2/s^2$.

In this distribution function, x is either precipitation amount per event or interarrival time, k is the shape parameter, λ is the scale parameter, $\Gamma(k)$ is the gamma function, \bar{x} is the actual distribution mean, and s^2 is the variance of the actual distribution.

For the monthly grouping, the best-fit theoretical distributions are listed in Table 1. The best-fit theoretical distribution of group precipitation and interarrival time, and their actual frequency distributions are shown in Figures 3 and 4 respectively.

Monsoon Model

The monsoon period of JUN-JUL-AUG-SEP was modeled differently from the rest of the year. This model was based on the assumption that rainfall events occur from the fronts or waves of large storm systems. Duckstein, Fogel, and Davis (1975) presented a mixed-event winter precipitation model that considered independently arriving storm sequences

Table 1. Probability Distribution Functions and Associated Parameter Values Used in Simulating Hot and Dry Seasonal Precipitation Model.

Group	Months	Distribution Type	Parameters Values
Variable: Interarrival Days			
1	1, 2, 3	Gamma	1-day fix: 0.4103 $\lambda = 0.1593, k = 1.849$
2	4, 5	Geometric	1-day fix: 0.4766 $x = 4.625$
3	10	Gamma	$\lambda = 0.0621, k = 0.2832$
4	11, 12	Gamma	$\lambda = 0.0283, k = 0.4841$
Variable: Precipitation Per Event			
1	1, 2, 3, 4, 5	Gamma	$\lambda = 0.1103, k = 0.8873$
2	10, 11	Gamma	$\lambda = 0.0490, k = 0.0523$
3	12	Gamma	$\lambda = 0.1504, k = 2.6304$

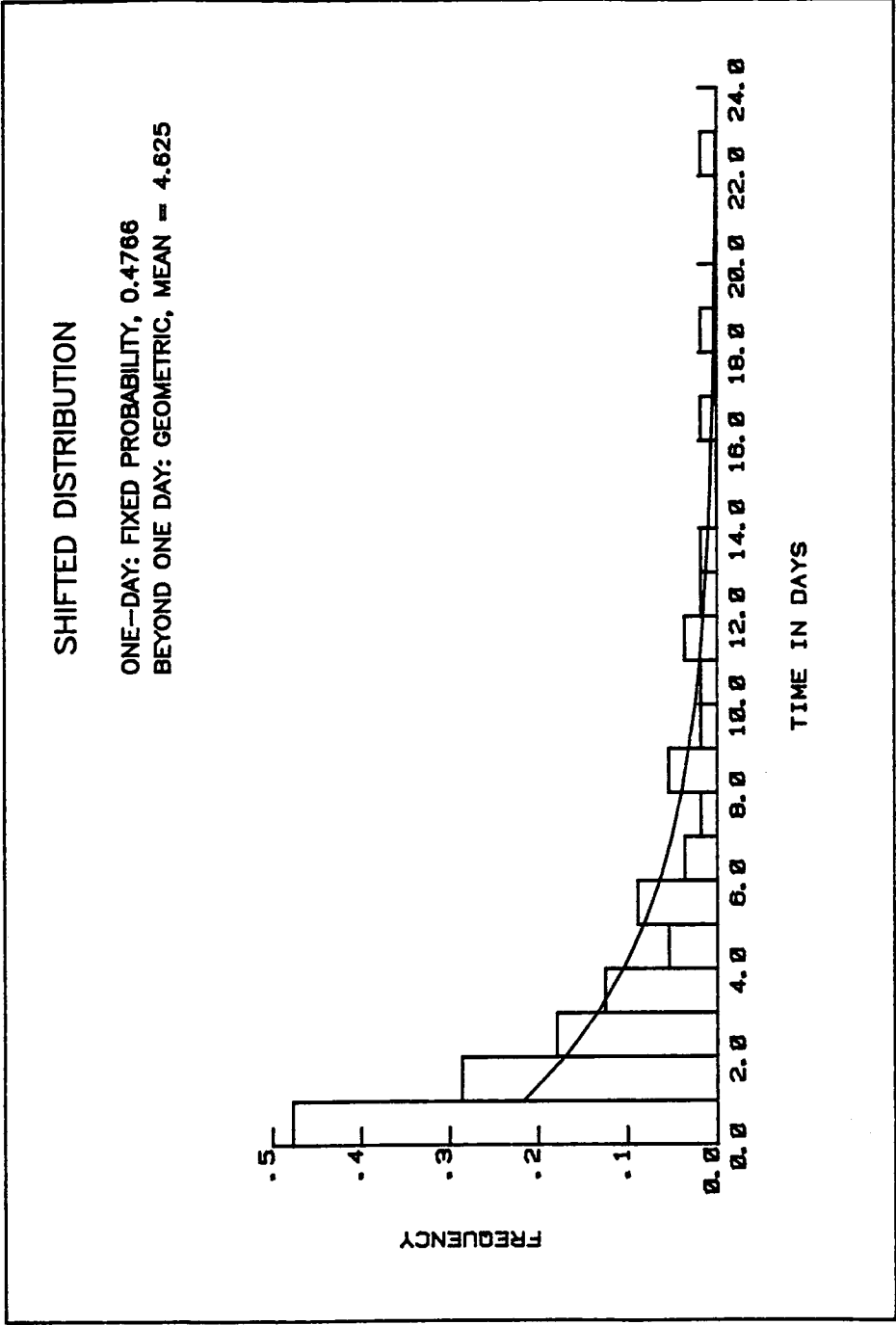


Figure 3. Comparison of Actual Data and Fitted Probability Distribution Function for Storm Interarrival Time in April and May.

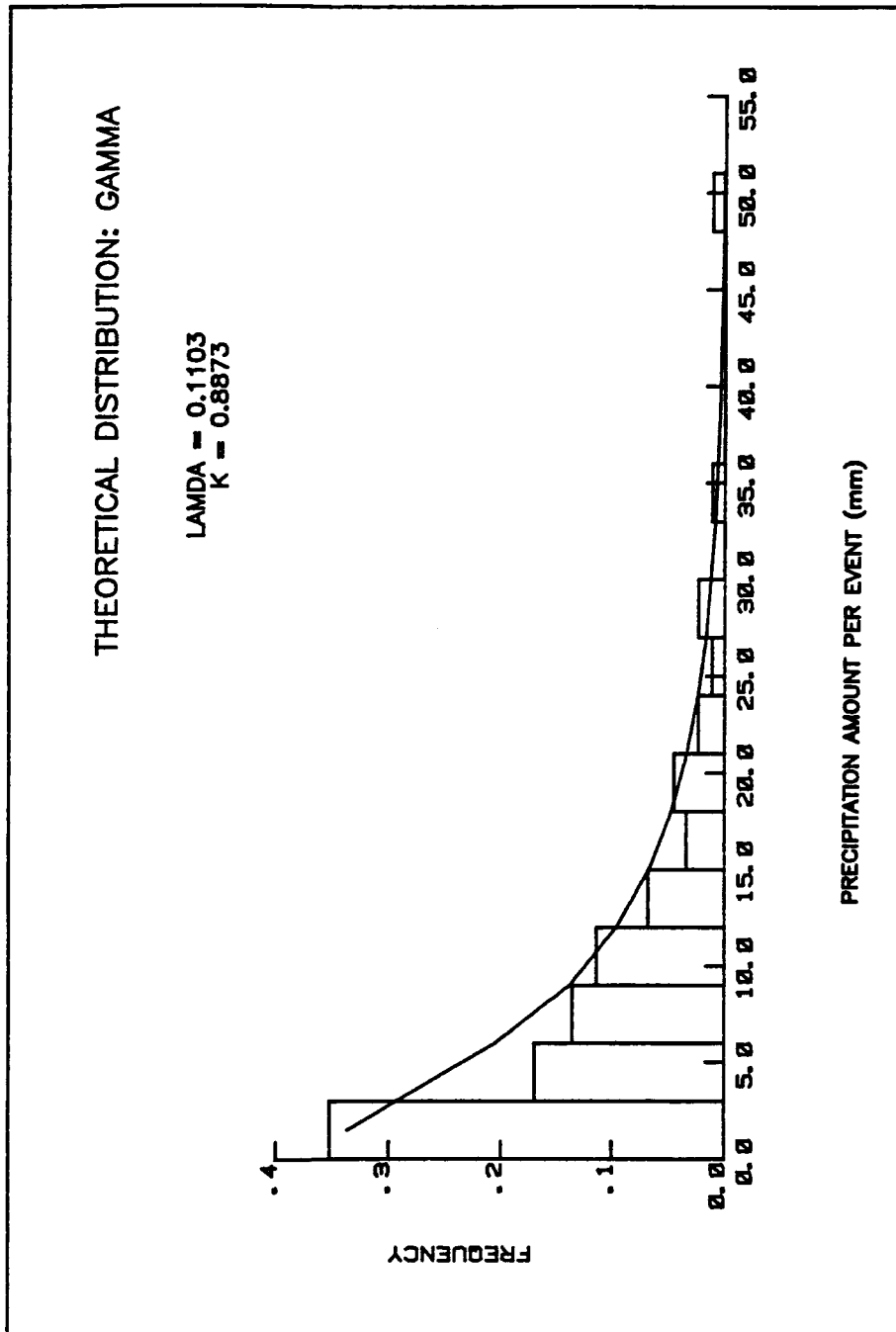


Figure 4. Comparison of Actual Data and Fitted Probability Distribution Function for Precipitation Amount per Event in January, February, March, April and November.

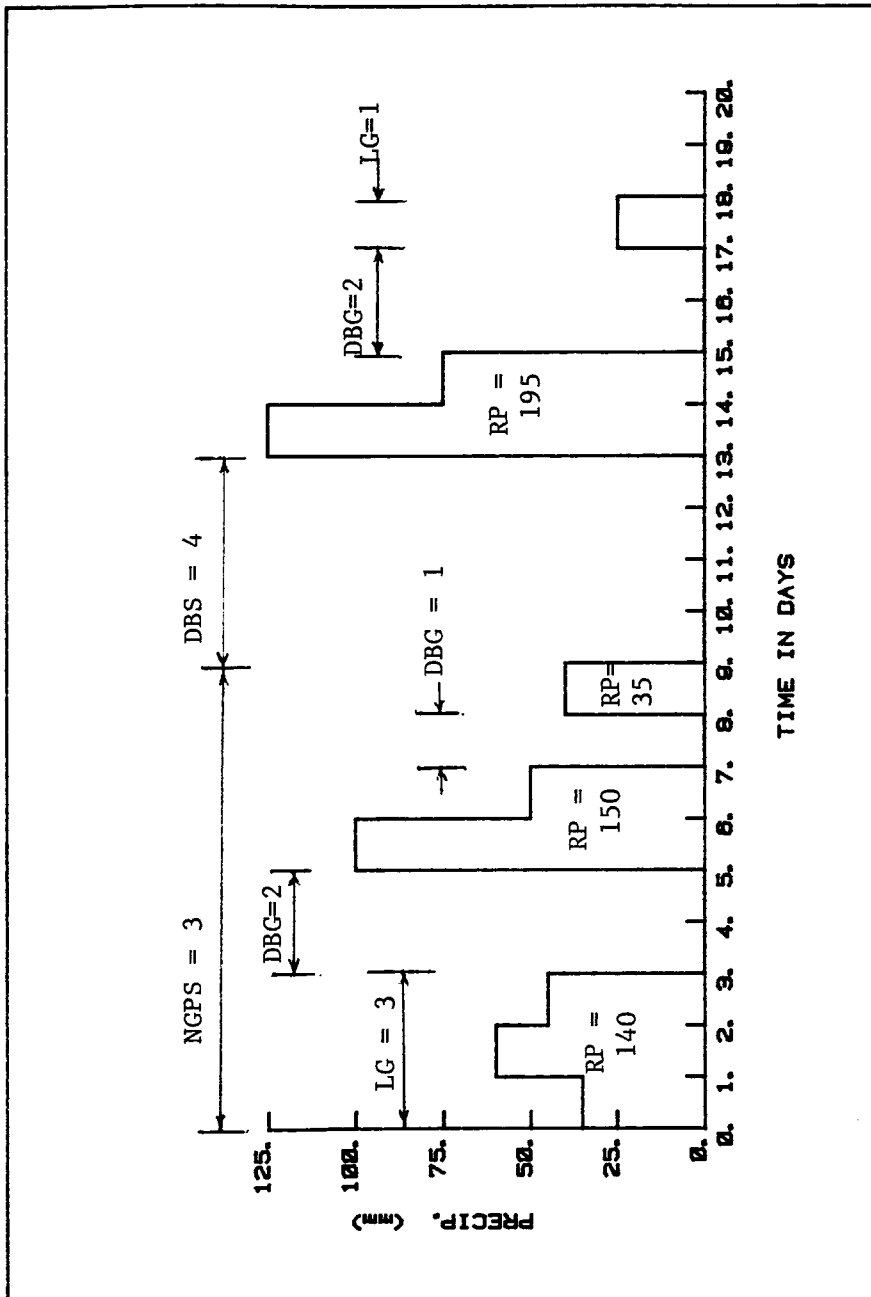
consisting of dependent groups. Monsoon storms often come in sequences in which one storm follows another within a short time interval, while at other times a storm appears to arrive in an independent manner. The element of precipitation unit in this model was precipitation group, defined as consecutive days of rainfall during which daily precipitation amounts exceeded a minimum of 0.1 mm. Consecutive groups separated by no more than a specified number of dry days constituted a sequence. Hekman (1977) estimated this maximum number to be three in mountainous areas of Arizona, U. S. In this study, rainfall events after five dry days were empirically considered to be in another sequence.

The random variables used in this model are schematically shown in Fig. 5, and defined as follows:

- (1) DBS, interarrival time between sequences (days)
- (2) DBG, interarrival time between groups (days)
- (3) LG, group duration (days)
- (4) NGPS, number of groups per sequence, and
- (5) RP, total precipitation amount per group (mm).

In this model, events correspond to groups of consecutive rainy days. Two assumptions made concerning the operation of the model are:

- (1) precipitation amount per group is uniformly distributed over group length, and



DBS = Dry Days Between Sequences DBG = Dry Days Between Groups
 NGPS = Number of Groups Per Sequence LG = Length of Group
 RP = Precipitation Amount Per Group (MM)

Figure 5. Wet Monsoon Type Precipitation Model Components.

(2) group precipitation is also independent of group duration.

Frequency distributions were developed for the five variables, and hypothetical or theoretical distributions were fitted to each of the actual distribution groups. Goodness of fit tests were conducted to compare actual and hypothetical distributions. The K-S test at a significance level of 0.05 was used to determine best-fit between the actual distribution and the theoretical.

Geometric distributions were used to describe the variables (DBG) number of dry days between groups and (LG) length of group. Dry days between sequences (DBS) were fit to the exponential distribution and number of groups per sequences (NGPS) were fit the gamma distribution. The cumulative density function (CDF) of geometric and exponential distributions are as follows:

Geometric distribution (CDF)

$$F_x(x) = 1 - (1-p)^x, \text{ where } p = 1/\text{Mean} \quad (3)$$

Exponential distribution (CDF)

$$F_x(x) = 1 - e^{-\lambda x}, \text{ where } \lambda = 1/\text{Mean} \quad (4)$$

In these distribution, x is one of the random variables DBG, DBS, NGPS, LG, and Mean is the actual distribution mean.

Precipitation amount per group did not fit any of the theoretical distributions, so this variable was grouped into different classes having intervals of 10 mm rainfall. These Class groups fit the gamma CDF's, so a mathematical relation was established to estimate precipitation amount in mm.

$$PPT = (10 x - 5) \quad (5)$$

where, PPT = precipitation amount in mm

x = grouped precipitation amount class, and

10 and 5 = the class interval and the interval mean,

respectively.

A summary of model parameters used in simulation is listed in Table 2, and best-fit theoretical distributions for all variables are schematically presented in Figure 6, 7, 8, 9 and 10.

Operation of the Model

The precipitation model subroutine PANI (See Appendix A) used in this study combined two submodels for dry and wet type rain storms, and operated as an event-based stochastic model of daily rainfall in the Tinau watershed. This model PANI generates a one year time series of the Julian calendar (Jan-Dec). The model operates by a series of Monte Carlo simulation (Benjamin and Cornell, 1970), using random numbers generated by the computer to sample the

Table 2. Probability Distribution Functions and Associated Parameter Values Used in Simulating Monsoon Precipitation Model.

Variable	Distribution Type	Parameter Values
1. Days Between Sequences (DBS)	Exponential	$x = 8.722$
2. Days Between Groups (DBG)	Geometric	$x = 1.813$
3. Days in a Group (LG)	Geometric	$x = 3.730$
4. Groups per Sequence (NGPS)	Gamma	$\lambda = 0.1388$ $k = 0.7609$
5. Precipitation Amount per Group (RP) (grouped class)	Gamma	$\lambda = 0.0672$ $k = 0.5648$

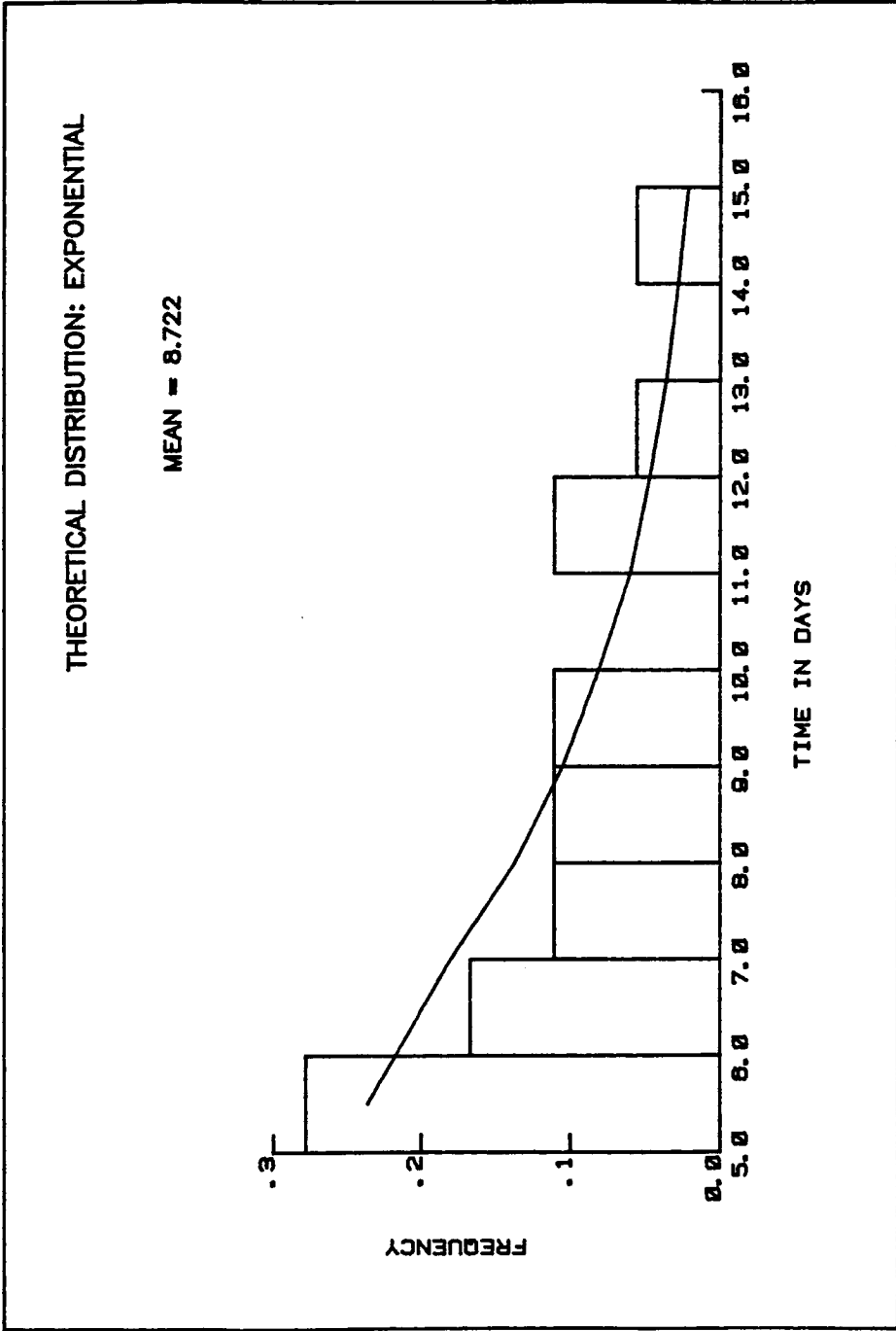


Figure 6. Comparison of Actual Data and Fitted Probability Distribution Function for Dry Days Between Sequences (DBS).

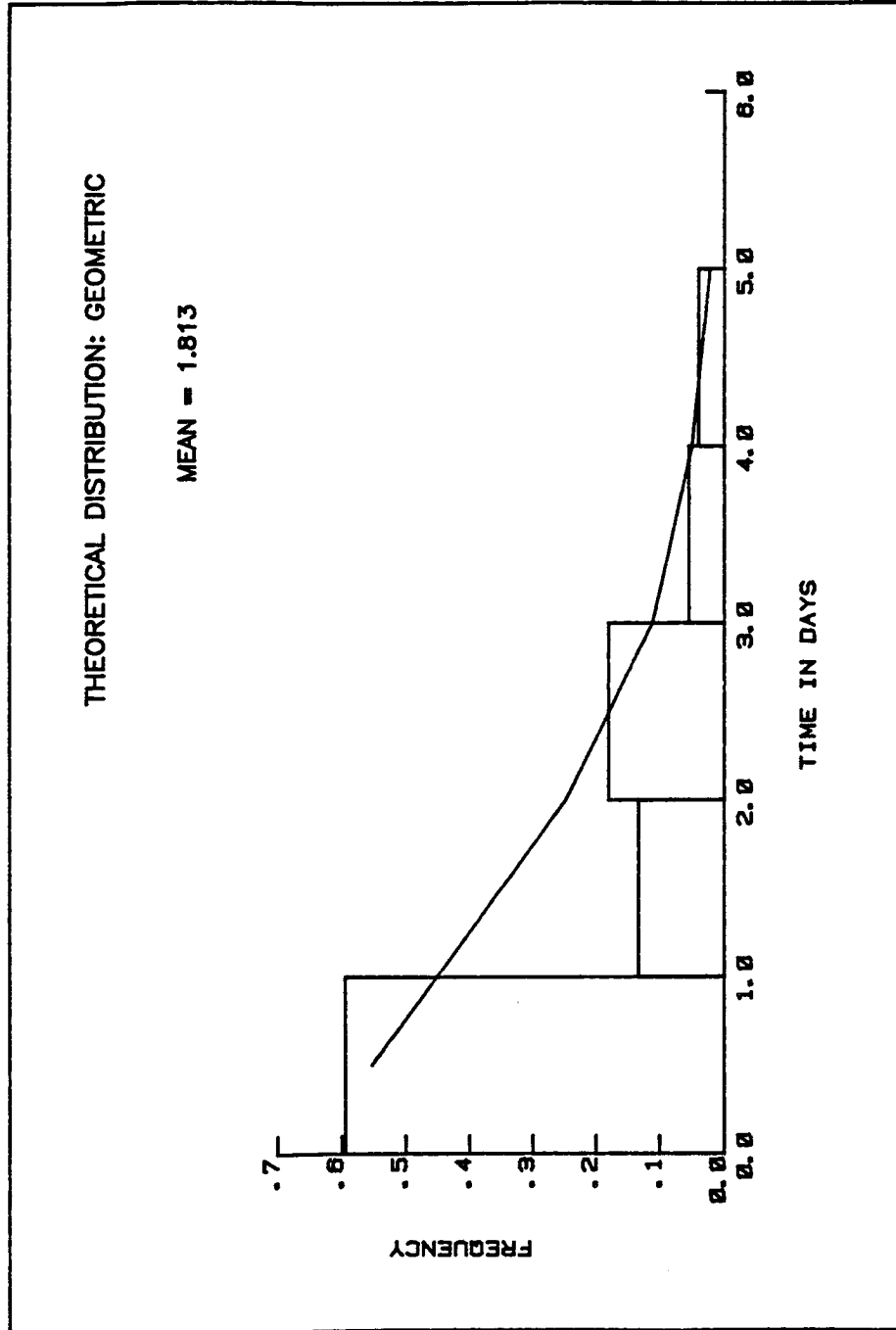


Figure 7. Comparison of Actual Data and Fitted Probability Distribution Function for Dry Days Between Groups (DBG).

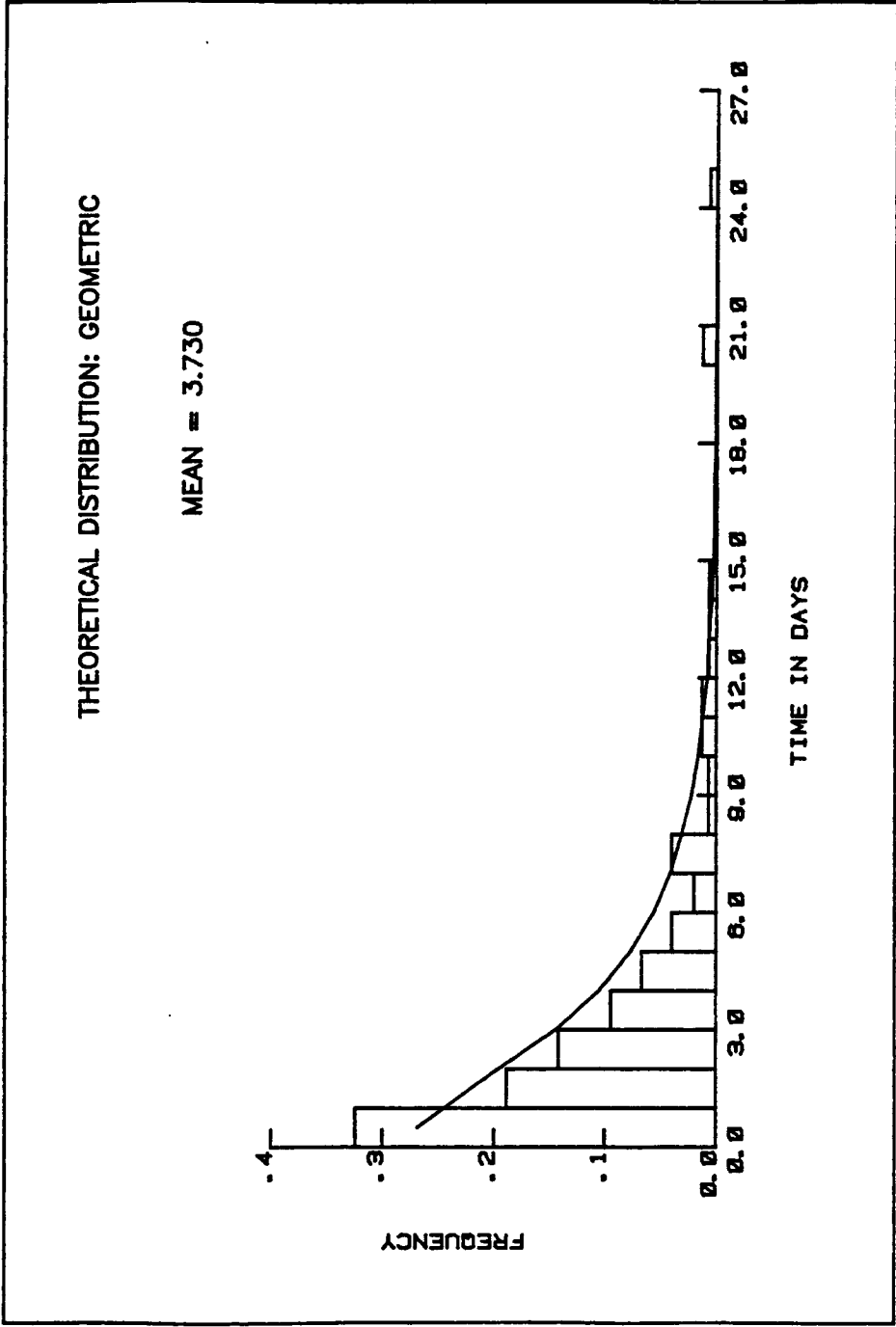


Figure 8. Comparison of Actual Data and Fitted Probability Distribution Function for Group Duration (LG).

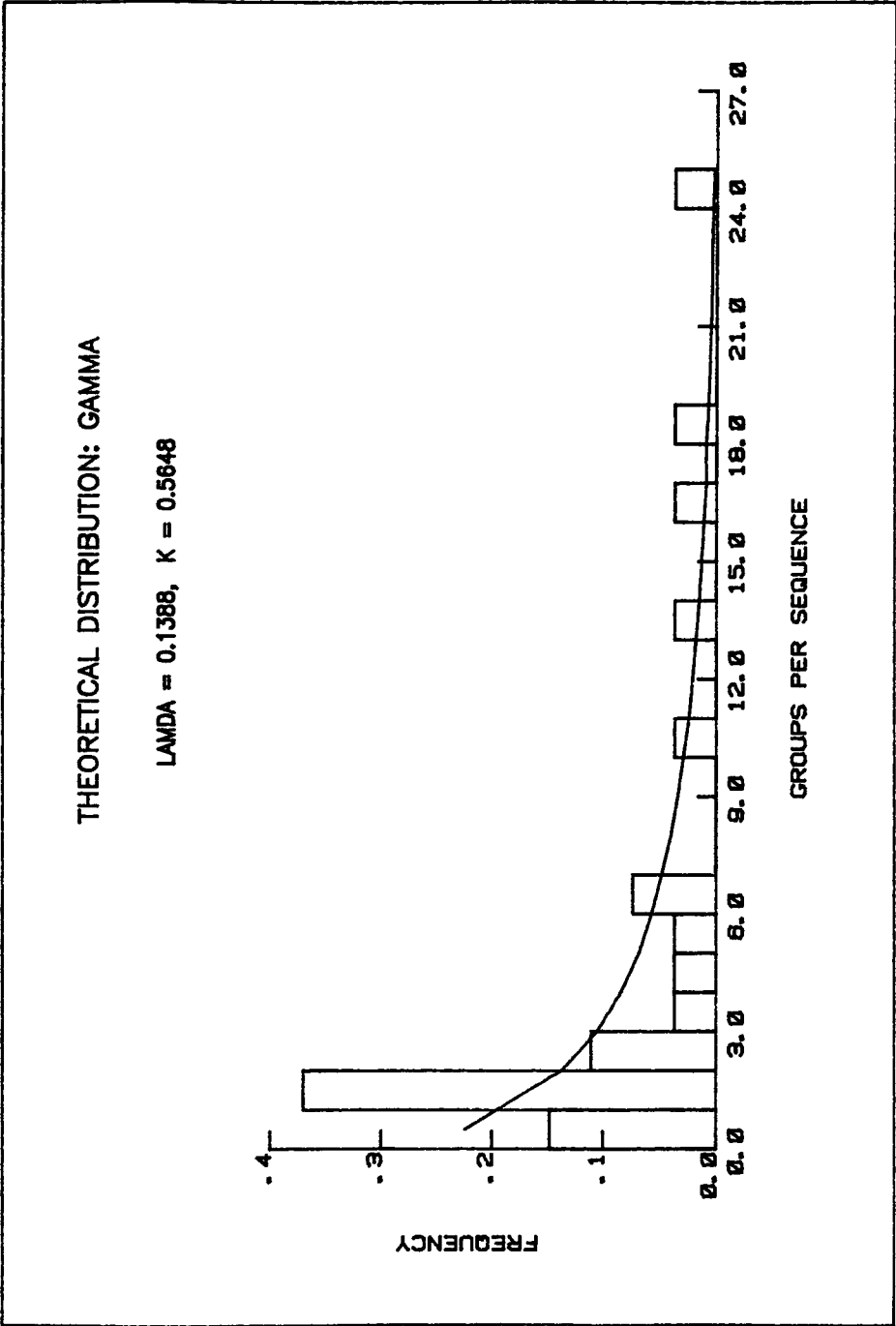


Figure 9. Comparison of Actual Data and Fitted Probability Distribution Function for Number of Groups per Sequence (NGPD).

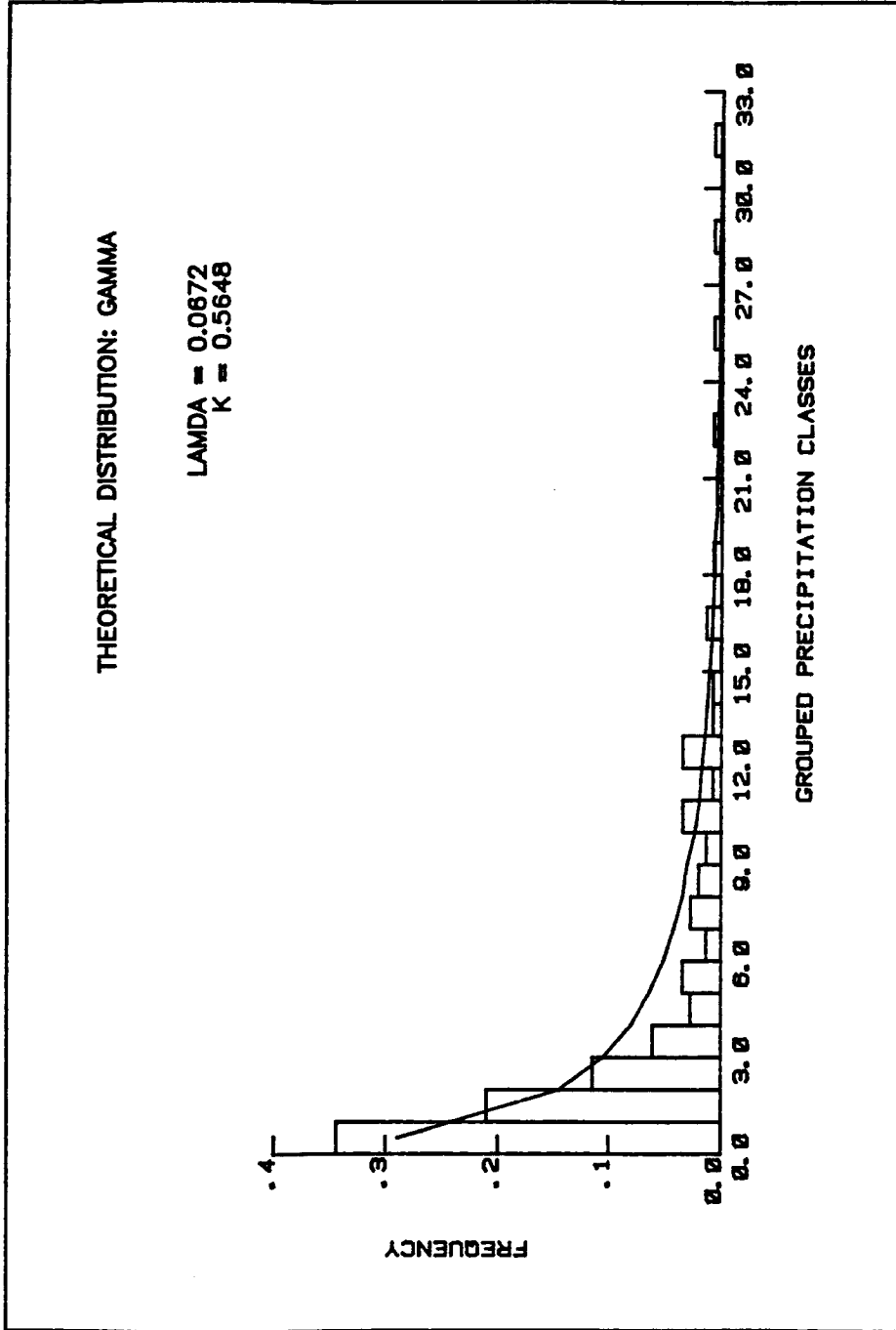


Figure 10. Comparison of Actual Data and Fitted Probability Distribution Function for Precipitation amount per Group (RP).

appropriate theoretical cumulative distribution function for each month and model parameters.

The cumulative probability function for any random variable is uniformly distributed over the interval 0 to 1, and random variables are generated from this probability distribution (Haan, 1982). These variables are generated by: generating a random number uniformly distributed between 0 and 1; setting that random number to CDF of the random variable distribution; and equating the distribution equation for the random variable occurrence. An example of the geometric distribution is as follows:

$$\text{CDF} = F_X(x) = 1 - (1-p)^x \quad (6)$$

where, $x = \text{Occurrence}$

$$p = 1/\text{Mean}$$

and $\text{CDF} = F_X(x) = \text{a random number.}$

In this equation, the occurrence x is determined by substituting values of CDF and p . Occasionally this inverse transformation could not be matched for a certain distribution as in the case of gamma. In these cases a gamma random variable generating function was used to describe interarrival time and precipitation amount per event.

The model was operated, starting from the first day of the Julian Calendar (Jan 1), in the following sequences.

(1) A random number is generated and that number is set to evaluate interarrival day CDF. The CDF equation is solved to determine interarrival day. The Julian calendar day is advanced keeping track of each seasonal group.

(2) On the day indicated for rainfall event, another random number is generated and CDF for precipitation per event is evaluated to give rainfall amount for that day.

As the calendar advanced, a check was made to determine if it entered the monsoon season. When the entrance was noticed, the following steps were operated.

(3) Generate a random number and enter it into the dry days between sequences distribution function. The calendar is advanced and counts the days of a year, accordingly.

(4) For the sequence indicated for rainfall groups, generate and enter another random number into the number of groups in a sequence function to evaluate number of groups.

(5) Generate and enter another random number into the group duration distribution function, and evaluate group length.

(6) Generate a random number and enter it into the distribution function for precipitation amount per group.

(7) Evaluate the precipitation amount on each day of the group by dividing total precipitation amount per group by number of days in a group.

(8) When more than one group was observed in a sequence then step 5 was repeated. And when there was only one group

in a sequence then another sequence was determined repeating step 3.

The calender is still in progress, so a check is made whether the date has the monsoon season or not. If it is the hot and cold dry season repeat steps 1 and 2 till the are repeated untill day of the calendar.

The calibrated model was verified to test its ability to estimate long-term seasonal precipitation patterns. The test was conducted by simulating 100 years of precipitation events. Mean number of precipitation events per seasonal group and mean precipitation amount per event or group were calculated and compared with actual data. A summary of the comparisons are presented in Table 3.

Table 3. Comparison of Precipitation Characteristics
Between the Historical 8 Year Record and the
Simulated 100 Year Record for the Tinau Watershed.

Month	Mean Monthly Precipitation		Mean Precip. per Event		Mean Number of Events per Month	
	Actual-Simulated (mm)	Actual-Simulated (mm)	Actual-Simulated (mm)	Actual-Simulated (mm)	Actual-Simulated (Number)	Actual-Simulated (Number)
JAN	21.2 15.6	9.9 7.5	2.1	2.1		
FEB	11.0 21.3	9.7 8.2	1.2	2.5		
MAR	18.1 26.6	6.9 7.6	2.7	3.4		
APR	33.4 46.5	8.0 7.9	4.3	5.9		
MAY	95.5 105.1	9.7 12.3	9.9	8.5		
OCT	48.8 58.8	13.4 11.4	2.7	5.1		
NOV	4.5 10.8	5.2 7.5	0.9	1.4		
DEC	21.9 33.6	17.5 18.1	1.3	1.8		
Sub-Total	254.4 317.3		25.1	30.7		
Monsoon seasonal precipitation and groups (JUN, JUL, AUG, SEP)						
	1482.8 1486.7	80.2 82.6	18.5	18.0		
Total						
Annual rainfall (mm); Actual = 1737.2, Simulated = 1805.0						
Annual rainevents (days); Actual = 94.1, Simulated = 97.8						

CHAPTER FOUR

THE STREAMFLOW GENERATION MODEL

The U.S.D.A., SCS runoff equation has been successfully used to transform basin precipitation into daily streamflow in a semi-arid areas. This formula evaluates direct runoff, but is unable to simulate baseflow. So in humid and sub-humid watersheds with perennial streams, a soil moisture accounting model is needed to simulate long recession flows in dry season after the wet monsoon. The Tinau watershed lies the in sub-humid zone with an average annual rainfall of 1700 mm. The monsoon and post-monsoon storms contribute perennial flows in the Tinau river. One of the objectives of this study was to develop a methodology which required only readily obtainable information, so it was necessary that the model be able to transform synthesized long-term daily climatic data into a corresponding record of daily streamflow volumes.

The transformation of rainfall to streamflow was accomplished by the use of a deterministic hydrologic model. The criteria used in selecting a hydrologic model for this study were: the input requirement for the calibration and operation of the model must be limited to daily precipitation, daily temperature, and daily runoff volumes;

calculated parameters must be minimized; the model must be capable of accurately modeling baseflow recessions; and if possible there should be a precedent for its use.

The Generalized Hydrologic Model (GHM) contained in the Generalized Streamflow Simulation System (GSSS) developed by the Joint Federal-State River Forecast Center in Sacramento, California (Burnash et al. 1973) met the above criteria, and was selected for this study. The GHM is a fitted parameters, conceptual model which uses a system of soil moisture storage and drainage components to simulate perennial streamflow. The conceptual logic of the GHM was designed to approximate the physical processes which occur on a watershed. Burnash et al. (1973, pp.11) describe this hydrologic model in the following manner:

It (model) is based on a system of percolation, soil-moisture storage, drainage, and evapotranspiration characteristics which are intended to represent the significant hydrologic processes in a rational manner. Each variable included in the model is intended to represent a discrete and understandable characteristic required for effective hydrologic analysis. The definition of model parameters is achieved by establishing a soil-moisture computation which allows the determination of basin streamflow from basin precipitation. Effective moisture storage capacities in the soil profile are estimated not by sampling of the soil profile, but by inference from the rainfall and discharge records.

The structure and basic components of the GSSS runoff model are presented in Figures 11 and 12. The model operates by routing basin precipitation into a series of cascading soil

A GENERALIZED
HYDROLOGIC MODEL

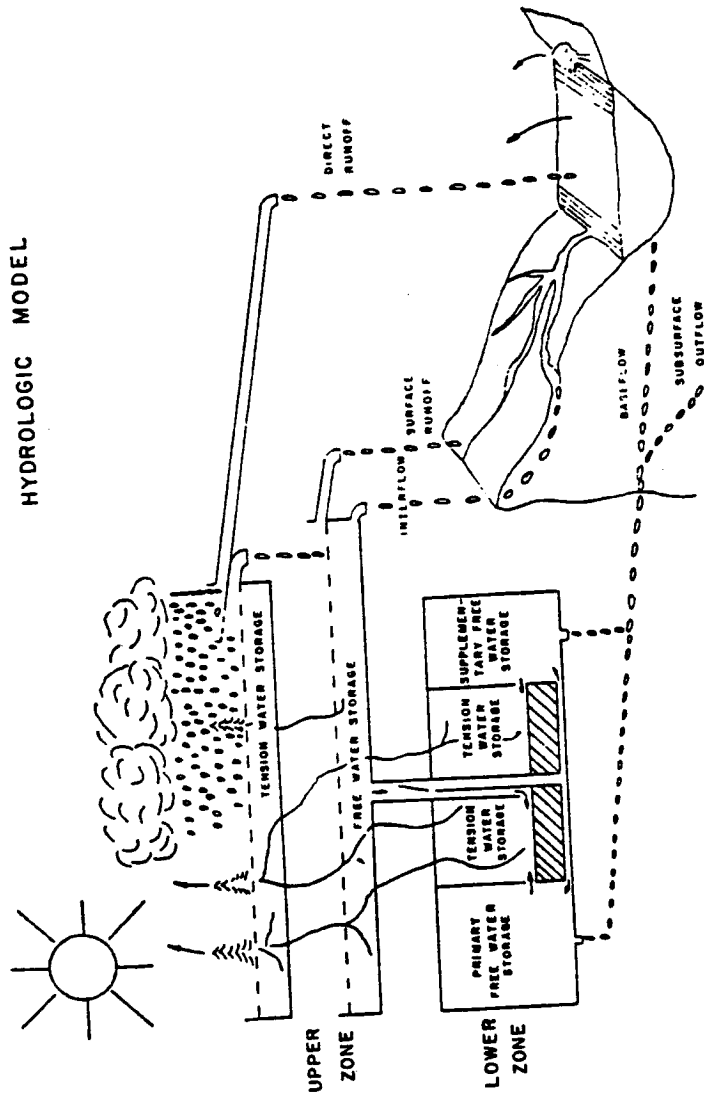


Figure 11. Conceptual Structure of the Hydrologic Cycle Used in the Generalized Streamflow Simulation System (From Burnash et al., 1973).

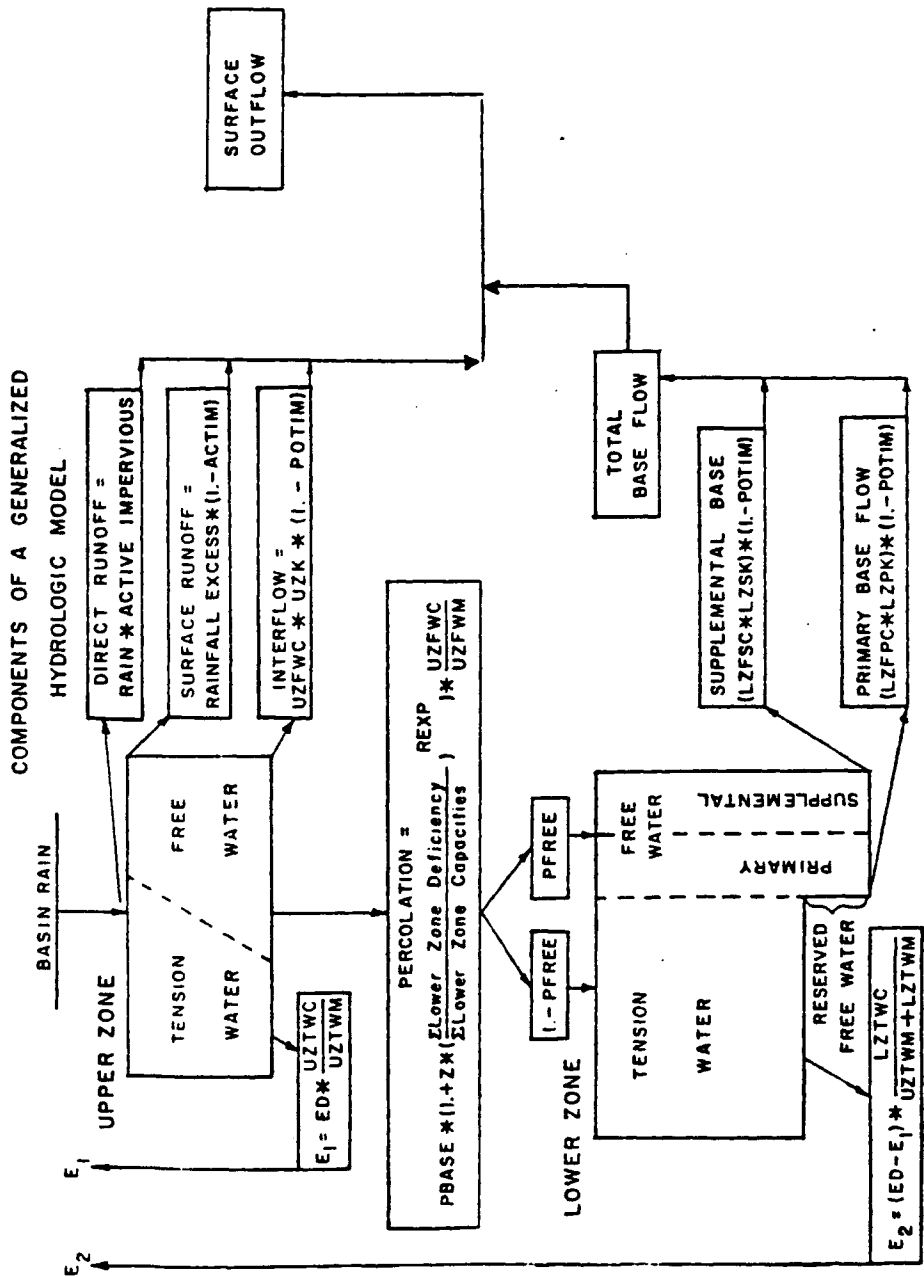


Figure 12. Principal Components of the Generalized Streamflow Simulation System (From Burnash et al., 1973).

moisture compartments, and slowly drains the soil water to form interflow and baseflow. The drainage depends upon the maximum soil water storage capacity of these compartments and the soil water content at present at them.

Rainfall encounters both permeable and impermeable surface areas. Impermeable areas, such as lakes and streams, transform all rainfall into runoff, while permeable areas transform direct runoff only when rainfall exceeds infiltration capacities. The GSSS allocates permeable areas into two layers upper and lower zone. Each zone is composed of tension water and free water storages. Upper zone tension represents the precipitation volume required under dry conditions to satisfy interception requirements and to provide sufficient moisture to the upper soil mantle for horizontal drainage and to begin deep percolation. When the upper zone tension volume is filled to its capacity, the excess precipitation is routed into upper zone free water storage. In this storage, horizontal interflow and vertical percolation to the lower zone originate. Percolation rate is controlled by the upper zone free water contents and capacities of lower zone moisture volumes. Interflow occurs only when the precipitation exceeds the maximum percolation rate. When the maximum rates of both interflow and percolation are exceeded and upper zone storage are full, surface runoff results.

Water percolated from upper zone reaches to the lower zone, where it is divided among three storages: tension, primary free, and supplementary free. Lower zone tension water capacities associated with suction forces absorb most of the percolating water. The percolated water that is not absorbed as lower zone tension water is distributed to primary and supplementary free water storages in response to their relative capacities. The horizontal output from the lower zone free water storages constitutes the baseflow. The differential release rate of the free water zones enable the modeling of a wide variety of streamflow recession curves.

The GSSS model expresses the watershed as a set of storages of determinable capacities, which hold water temporarily and which gradually recede as their contents are diminished by vertical percolation, evapotranspiration and/or lateral drainage. The important parameters in the model are evaluated by analyzing historical precipitation and simultaneous streamflow records. Burnash et al.(1973) utilized twenty four parameters in operating this model, and presented initial parameters estimation method. The critical parameters of the model are defined as follows:

Soil Moisture Storage Capacities

UZTWM, Upper Zone Tension Water Capacity

UZFWM, Upper Zone Free Water Capacity

LZTWM, Lower Zone Tension Water Capacity

LZFSM, Lower Zone Free Supplementary Water Capacity

LZFPM, Lower Zone Free Primary Water Capacity

Soil Moisture Drainage Rates

UZK, Lateral drainage rate from the upper zone
expressed as a fraction of the contents per day

LZSK, Drainage rate for lower zone supplementary water

LZPK, Drainage rate for lower zone primary water

Percolation Rate

ZPERC, Controls the percolation rate from upper to lower
zone water

REXP, Exponent which controls change in the percolation
rate

There are fourteen other parameters in the GSSS model; they are all less sensitive as compared to parameters mentioned above or self-adjusting. A list of all these parameters are presented on subroutine DOVAN (see Appendix A).

Initial Parameter Estimation

The critical parameters, such as drainage rates and water storage capacities, were estimated for the GSSS runoff model operation. Examples of the procedures of estimation of soil moisture drainage rates LZPK and LZSK, and upper zone tension water content UZTWC are explained below.

The lower limits of soil moisture drainage parameters LZSK and LZPK were estimated from historical

hydrographs analysis. These parameters were used to calibrate the simulated streamflow recession curves. The assumption made in calibration was that long-term recession curves can be modeled as the sum of drainage from two lower zone soil moisture, such that:

$$\text{Baseflow} = \text{LZPK} * \text{LZFPC} + \text{LZSK} * \text{LZFSC} \quad (1)$$

where LZFPC and LZFSC are the current soil moisture contents of the soil moisture storage compartments, and LZPK and LZSK are as defined above. The moisture contents of the soil moisture storage compartments are updated daily.

The primary baseflow recession characteristics were estimated by a simple exponential decay function:

$$QP_t = QP_0 * K^t \quad (2)$$

where, K is the recession coefficient of primary baseflow for time unit used, usually daily.

t is the number of time units in days.

QP_0 is a discharge when recession is occurring at primary rate.

QP_t is the discharge t time units later.

Solving for decay factor K,

$$K = (QP_t / QP_0)^{1/t} \quad (3)$$

The lower zone primary moisture drainage rate is simply estimated as follows:

$$LZPK=(1-K) \quad (4)$$

$$\text{and, } LZFPM = QP_{\max}/LZPK \quad (5)$$

where LZPK and LZFPM are lower zone primary soil moisture drainage rate and capacity, and QP_{\max} is the maximum primary baseflow that can be inferred from hydrograph analysis.

An initial estimate of these parameters was derived by substituting actual runoff data for QP_t and QP_o in the equation (3) and then using the equation (4) and (5). These drainage rate parameters represent maximum baseflow capacity for the watershed, thus, the historical recession curves chosen for this analysis are those from the largest monsoon flow events.

In estimating these parameters, five monsoon rain-storm flow recessions (1978-1982) were selected. Each was plotted on separate log-normal paper and fit with three straight lines representing early, middle, and late streamflow responses. These linear recessions were combined to form composite recession curves. The composite recessions are presented in Figure 13. The early recession represents streamflow from the upper zone soil moisture storage and that from lower zone compartments, the middle section represents streamflow from two lower zone compartments, and

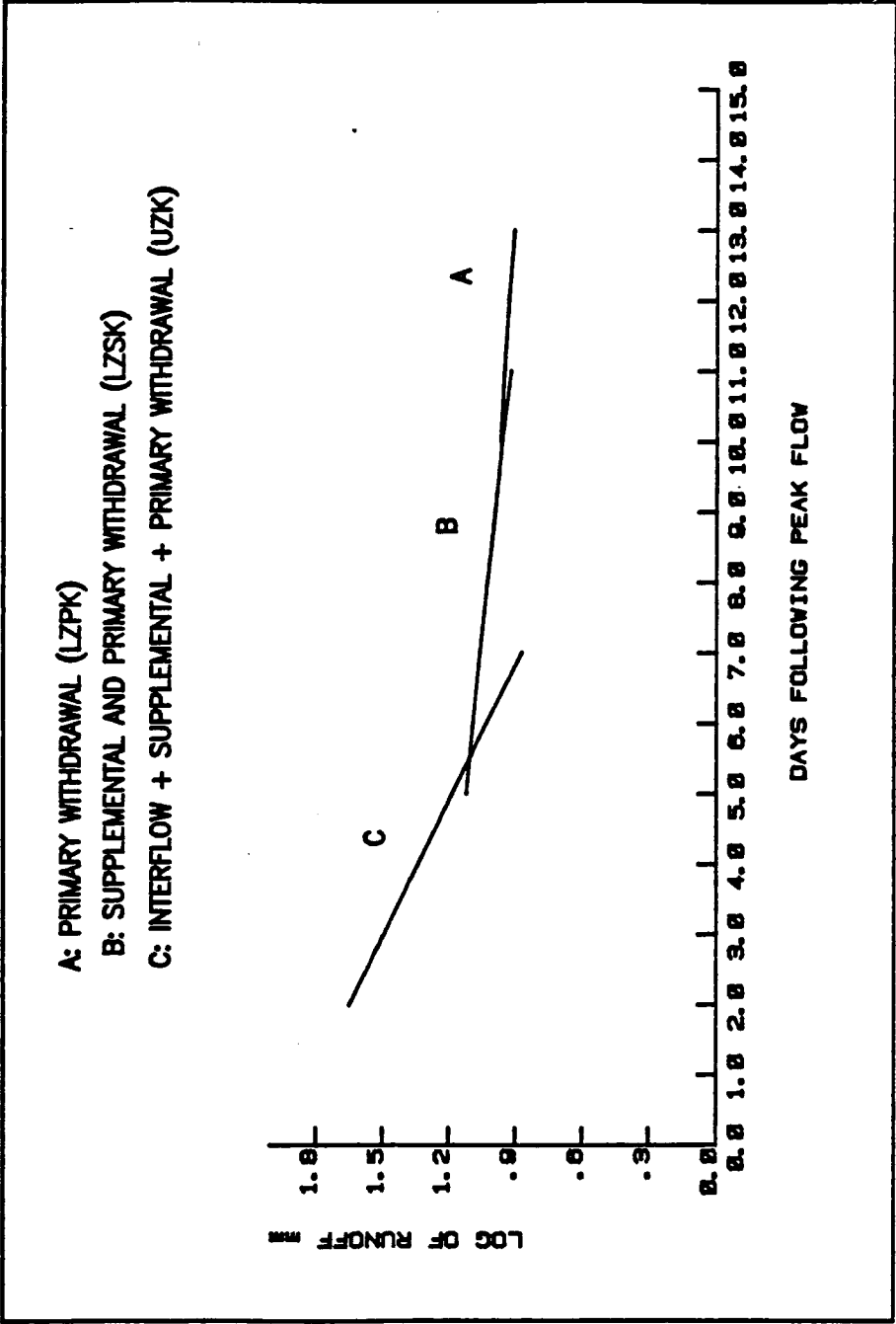


Figure 13. Composite Recession Curves Developed from Five Monsoon Rainstorm Recessions at the Tinau River, Butwal.

late recession curve represents streamflow from the lower zone primary storage.

An initial value of 0.03 per day was estimated for LZPK, and 363.33 mm for LZFPM. These values were later adjusted in the final model calibration.

The supplemental baseflow characteristics were evaluated as follows (Burnash et al.,1973):

$$K_s = (Q_{s1}/Q_{s_{max}})^{1/t} \quad (6)$$

where, K_s is rate of change of supplemental flow

Q_{s1} is lower zone supplemental discharge arbitrarily used at 2/3 time from peak to Q_{P_0} , and

$Q_{s_{max}}$ is lower zone supplemental discharge at time zero

These supplemental rates were estimated by extending the late recession curve backward underneath the middle curve Figure 14. Drainage rates were calculated as:

$$LZSK = (1-K_s) \quad (7)$$

$$\text{and, } LZFSM = Q_{s_{max}}/LZSK \quad (8)$$

where, LZSK is rate of change of supplemental storage, and

LZFSM is maximum apparent supplemental storage.

Values of Q_{s1} and $Q_{s_{max}}$ from the middle recession curve were substituted in the equations (6) and (7) to evaluate an initial estimate of LZSK = 0.158 and LZFSM = 49.37 mm, which were later adjusted in the final calibration.

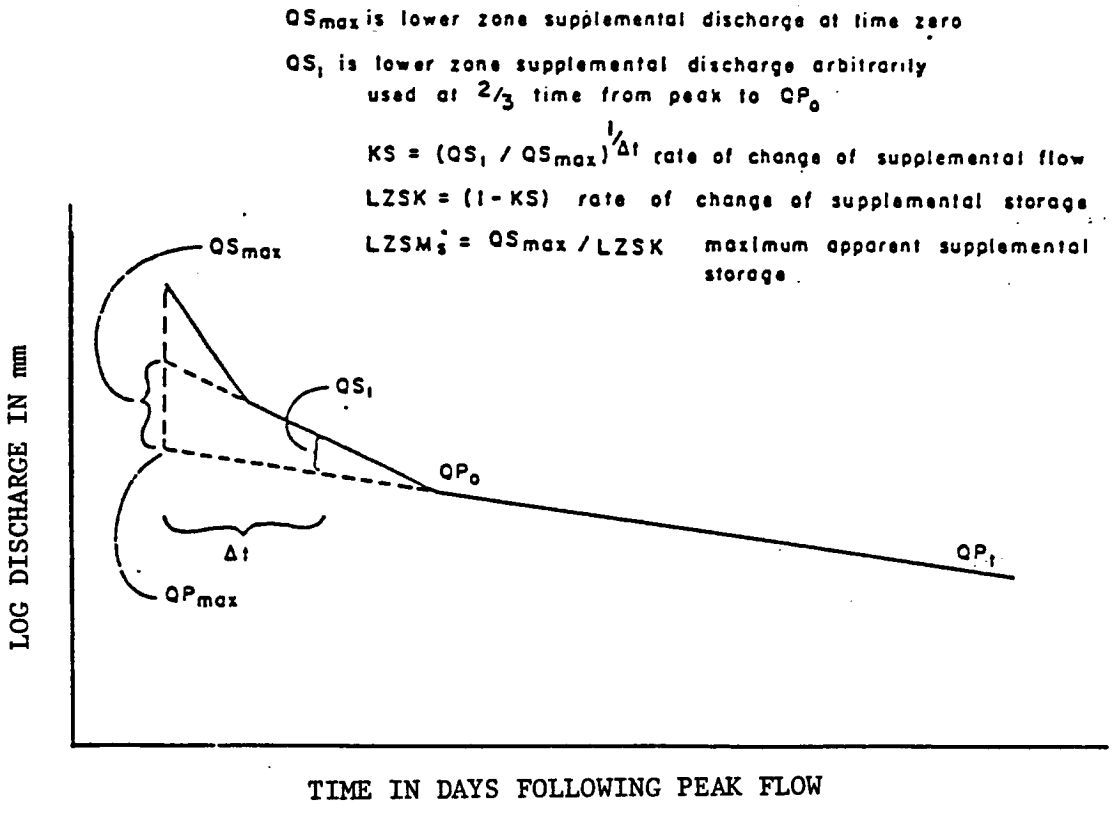


Figure 14. Initial Approximate of Lower Zone Supplemental Water Storage and Depletion Coefficient (From Burnash et al., 1973).

The upper zone soil moisture storage capacity parameter UZTWM is critical for determining both the timing of initial monsoon runoff and the volume of peak one-day flows. The UZTWM simulates upper soil layer which absorbs rainfall until its moisture storage capacity is filled, then spills the excess as surface runoff. Every day, a certain amount of this moisture storage volume is lost to evapotranspiration, percolation, and interflow.

Burnash et al. (1973) have described the initial estimation of UZTWM. For this study a portion of the historical record, following a dry period when evapotranspiration had depleted the upper soil moisture was used to estimate UZTWM. The UZTWM is that volume of rainfall which is retained before the runoff includes any increase beyond that which can be produced by rain on the impervious fraction. Three sequences of extended dry periods followed by several consecutive days of heavy rain were analyzed, one is as follows:

Day	Basin rain (mm)	Observed Discharge (mm)
Jun 4, 1982	0.0	0.99
Jun 5, 1982	34.9	1.04 *
Jun 6, 1982	24.9	1.04 *
Jun 7, 1982	7.0	1.04 *
Jun 8, 1982	0.0	3.87 **

* Direct runoff

** Surface runoff

The sudden flow on June 8, 1982; indicates that soil moisture storage was full and surface runoff exceeded the direct runoff. The UZTWM was estimated as the amount of rainfall before surface runoff begins.

$$\begin{aligned} & \text{Rain (day one) + Rain (day two) + Rain (day three)} \\ & + \text{Rain (day four) = UZTWM} \\ & 0.00 + 34.9 + 24.9 + 7.0 = 66.8 \text{ mm} \end{aligned}$$

This value of UZTWM was an initial approximation. In another sequence, it was 76.00 mm which was the minimum estimate, because it was not known what rainfall amount on the day after direct runoff should be added to the minimum estimate.

The remaining soil moisture parameters in the GSSS runoff model are estimated by trial and error optimization during successive calibration runs of the model. Burnas et al. (1973) gave starting values and ranges for trial and error parameters estimation. An optimization process was used for preliminary fitting to historical record in this study.

Evapotranspiration Estimates

The GSSS runoff model requires a daily estimate of potential evapotranspiration (PET) as an input. Evapotranspiration can be the dominant use of the moisture supplied by precipitation, and since it is one of the most difficult process to evaluate in hydrologic analysis, evapotranspira-

tion is frequently a principal source of error in streamflow simulation. Two year (1984-1985) of incomplete pan evaporation data were available. Because of incompleteness and the difficulty in transforming data from the pan site to the the study site the data could not be used. Therefore the Blaney-Criddle method was used to estimate monthly evapotranspiration demand.

Daily maximum and minimum temperature at Tansen were available from year 1980-1984, and monthly maximum and minimum temperature at Butwal from 1978-1984. A consumptive use estimation equation used in this study required these data. The Blaney-Criddle equation was developed to estimate consumptive use in South-Western United States. It estimates mean monthly PET from mean monthly temperatures, the latitude of the site, and a crop factor (Dunne and Leopold, 1978):

$$PET = (0.142 T_a + 1.095) * (T_a + 17.8) * K * D \quad (9)$$

where, T_a is mean air temperature for the month

PET is mean monthly potential evapotranspiration

D is monthly fraction of annual hours of daylight

K is a crop fractor chosen from a table of suggested values.

In equation (9), D was evaluated from a table for the latitude of Tansen, and a weighted K was evaluated for each month with respect to land-use and ground cover in the

watershed. Mean temperatures were calculated from the temperature data from Tansen and Butwal. Yearly pan evaporation recorded at Tansen is 775 mm, and that at Butwal is 1708 mm. These values were used to verify the Blaney-Criddle method. The results of the comparison are presented in Table 4.

The GSSS model suggests that a weighting factor for monthly PET rates be employed to account for seasonal plant growth characteristics. The monthly weighing factors chosen for this study were based on the crop rotation pattern and the vegetative cover in the watershed. The GSSS model uses the daily fraction of adjusted monthly PET, and daily loss calculated as the product of the adjusted PET values and a soil moisture volume to soil moisture capacity ratio.

Hydrologic Model Calibration and Testing

The GSSS model was calibrated by a series of trial simulation runs, in which the model parameters were adjusted until the simulated daily runoff sequence approached the actual daily runoff record (McDowell, 1985). The actual daily precipitation data were needed as input for each simulation. Simultaneous daily runoff data were used to validate the model calibration.

A year sequence of climatological and runoff data were used for model calibration. Six years (1977-1982) of daily streamflows from the Tinau river at Butwal were used for calibration. The actual and simulated streamflow records

Table 4. Summary of Potential Evapotranspiration Estimates for the Tinau Watershed, Palpa.

Month	Blaney-Criddle Estimate (mm)	Adjusted factor for GSSS	Adjusted monthly PET (mm)	Adjusted Daily PET fraction (mm)
January	55.8	0.60	33.5	1.1
February	70.0	0.60	42.0	1.5
March	125.0	0.70	87.5	2.8
April	135.7	0.70	95.0	3.2
May	139.6	0.80	111.7	3.6
June	96.0	0.85	81.6	2.7
July	93.0	0.85	79.1	2.6
August	83.7	0.85	71.1	2.3
September	75.0	0.75	56.3	1.9
October	71.3	0.70	49.9	1.6
November	60.0	0.70	42.0	1.4
December	49.6	0.60	29.8	1.0
Total	1054.7 mm		779.5 mm	

displays on a HP-1000 computer and HP-7221B graphic plotter respectively. Criteria for the comparison of daily streamflow records included volume of peak flow days, shape and duration of recession flows, and total annual flow volumes. All comparisons were evaluated subjectively.

The water years selected for the final model calibration were 1980-1981 and 1981-1982. Two criteria used in calibration of the model were: model results should match recorded data; and the estimates of parameter values should be consistent with watershed characteristics. For the calibration years, rainfall was found to be uniformly distributed over the watershed as indicated by a close correlation between total rainfall recorded in Butwal and in Tansen.

An interactive version of the GSSS runoff model was written and modified to facilitate rapid parameter adjustment. Critical parameters were adjusted by trial and error, using the initial parameter values described in the preceding sections as a first trial. Parameters were adjusted using the conceptual structure and function of the GSSS model. For example, if the simulated streamflow was high, the soil moisture storage parameters were increased, and if it was low the storage parameters were reduced. If a simulated streamflow sequence gave low peak flows and exaggerated base flows, percolation and/or lower zone soil moisture

storage the parameters were reduced. This adjustment increases the amount of water in the upper zone available for surface runoff and evapotranspiration, and reduces the water in the lower zone available for base flow. Burnash et al. (1973) described the parameter calibration and its effect on simulated streamflow. Definitions and the final estimated values for the major parameters of runoff model are presented in Appendix B.

The calibrated runoff simulation model was tested on a continuous and simultaneous six year sequence (1977-1982). It was considered important to test the model on a continuous multi-year data sequence to check for long-term trending errors. But, due to unavailability of streamflow and precipitation data, a long-term calibration could not be made. The performance of the calibrated runoff model is illustrated by the comparison made in Figure 15.

The calibrated streamflow simulation model was used in the main watershed model TINAU. The listing of the interactive version of the GSSS runoff model, used for calibration of the model with a one year data set (starting from June to May) is presented in Appendix A, as a subroutine DOVAN.

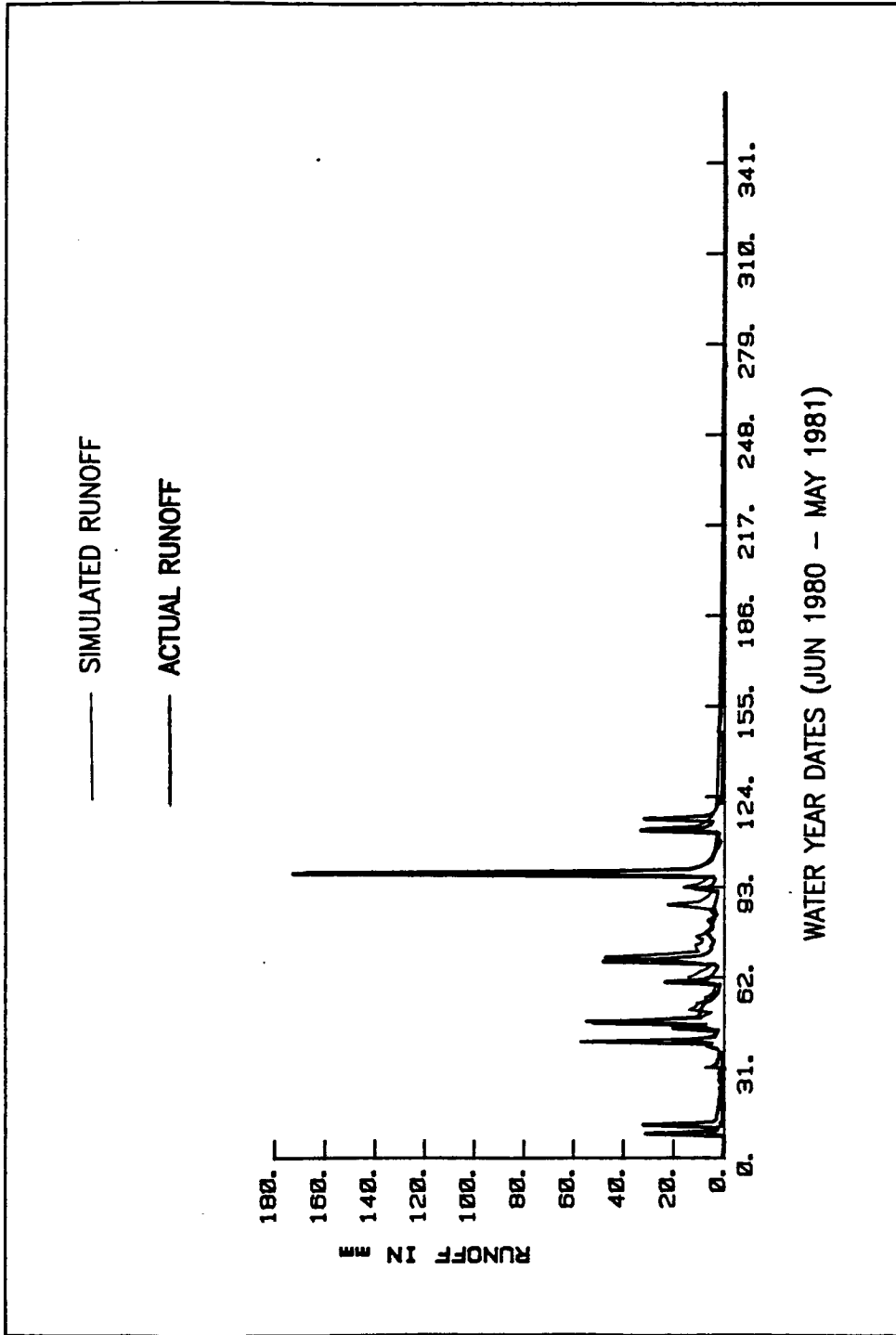


Figure 15. Comparison of Simulated and Actual Runoff at the Tinau River, Butwal for the Water Year 1980-1981.

CHAPTER FIVE

THE WATERSHED MODEL

All the component models of precipitation, streamflow generation described previously were linked to form a watershed model TINAU. TINAU is driven by the component model for stochastic precipitation simulation. The deterministic runoff model transformed this simulated rainfall data into daily streamflow. The TINAU performs a daily water budget and calculates the monthly and yearly runoff output statistics. It also computes consecutive days of low flow. A complete listing of the FORTRAN IV program for the TINAU is given in Appendix A. The TINAU was run for 100 years. Sample output of the monthly runoff frequencies, yearly and seasonal runoff frequencies, storm daily high and low discharge frequencies, and frequencies of consecutive days of low flow are given in Appendix C.

A flow chart of the TINAU program is illustrated in Figure 16. The model program is set up in discrete steps. The stochastic components of the precipitation model are independent of each other, but each step in the deterministic streamflow generation model is dependent upon the previous step. Each step in the deterministic model is driven by the stochastic model.

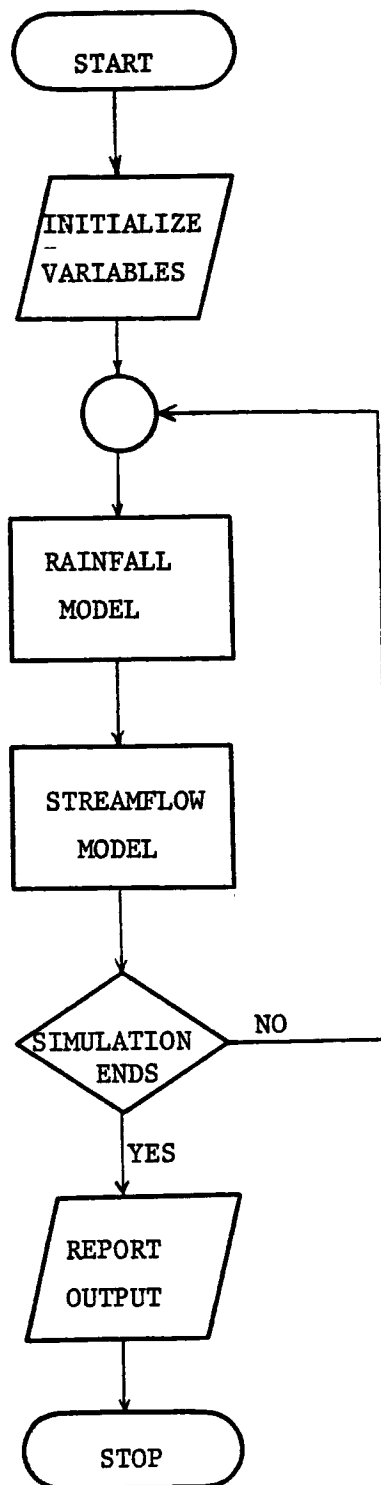


Figure 16. Flow Chart of the Watershed Model TINAU.

The TINAU program operates by simulating one year at a time, then storing each year's results for subsequent statistical analysis. The stochastic precipitation sub-model PANI is operated beginning with the first day of the Julian calendar (Jan 1). PANI was calibrated, run, and verified using Julian dates. PANI generates input variables, dry days between rainfall events and rainfall amount for the event according to a random process, which is guided by the statistics and the distributions derived from the actual data. PANI has capability to simulate a time series of rainfall data, and assign zero rain for each rainless dry day. Since, the deterministic streamflow generation model operates on the water year (June-May), the synthetic rainfall data was rearranged accordingly.

The deterministic streamflow generation model DOVAN uses mathematical equations to transform the daily synthetic rainfall data in runoff values. DOVAN accounts for soil moisture and evaporation, using daily evapotranspiration demand and soil water parameter values in a physically consistent processes. A statistical update is made for each day to determine the frequency of runoff amounts. The mean and standard deviation were calculated from the these frequencies on a monthly, seasonal and annual bases. The frequency of maximum and minimum flows were used to evaluate annual daily maximum and minimum events.

CHAPTER SIX

RESULT AND DISCUSSION

To evaluate the hydrologic performance of the Tinau river at Butwal, the watershed model TINAU was run for a 100 year simulation period. The results presented here are concerned with total annual, monthly and seasonal discharge, maximum daily high and low discharges, and consecutive days of low flow. These quantities were selected for evaluation because of their relevance to watershed management and water utilization in the valleys upstream, and to water yield and storm flows downstream.

Monthly Discharge

The monthly discharge of the Tinau river is influenced by the monsoon season. The probabilities of amount and timing of high and low discharge determine the reliability of water supplies during the non-monsoon season and the flood risks that may be involved during the monsoon months. A knowledge of the probabilities can help in the planning of watershed management activities, conserving water when it is plentiful and that release water when it is in short supply.

The simulated Monthly discharge of the Tinau river at Butwal for the non-monsoon season is presented in Table 5. The values in the table are the occurrence probability of monthly flow events equal to or less than the particular values. The distribution curves developed indicate that minimums flow of 0.15 mm or less could occur in January, April and May and that the probability of flows less than 1.5 mm is low in the post monsoon month of October. The simulation also indicated that 50 percent of the time the stream flow expected is less than 1.7 mm for all non-monsoon months, except October. A cumulative probability frequency plot for March, presented in Figure 17, and indicates that 90 percent of the time the stream flow in March is equal to or less than 0.71 mm. Results of simulated monthly discharge frequencies, means, and standard deviations are listed in Appendix C.

Annual and Seasonal Discharge

The simulated monsoon and annual discharges at Butwal varied over a wide range. The total monsoon discharge for the driest year of simulation was 120 mm, and for the wettest year it was 2,150 mm. The probability frequency analysis indicated that 90 percent of the time during the monsoon season the flow was equal to or less than 1,350 mm and 50 percent of the time equal to or less than 740 mm. The occurrence of flows equal to or less than 190 mm was only 5

Table 5. Monthly Discharge Probability of the Tinau River, Butwal

 Monthly Flow Frequency Probability
 - percent -

Month	10	20	30	40	50	60	70	80	90	99
JANUARY	0.52	0.63	0.70	0.78	0.85	0.92	0.99	1.08	1.20	1.65
FEBRUARY	0.42	0.46	0.51	0.56	0.61	0.68	0.75	0.80	0.97	1.60
MARCH	0.25	0.30	0.37	0.42	0.46	0.50	0.55	0.59	0.71	1.00
APRIL	0.21	0.25	0.27	0.30	0.33	0.36	0.39	0.44	0.49	0.70
MAY	0.15	0.17	0.21	0.23	0.25	0.27	0.33	0.43	0.53	2.50
OCTOBER	1.50	1.78	2.03	2.25	2.50	2.75	3.00	3.70	5.00	22.00
NOVEMBER	1.08	1.32	1.44	1.58	1.71	1.84	1.96	2.13	2.34	3.25
DECEMBER	0.77	0.90	1.01	1.11	1.20	1.30	1.39	1.51	1.67	3.00

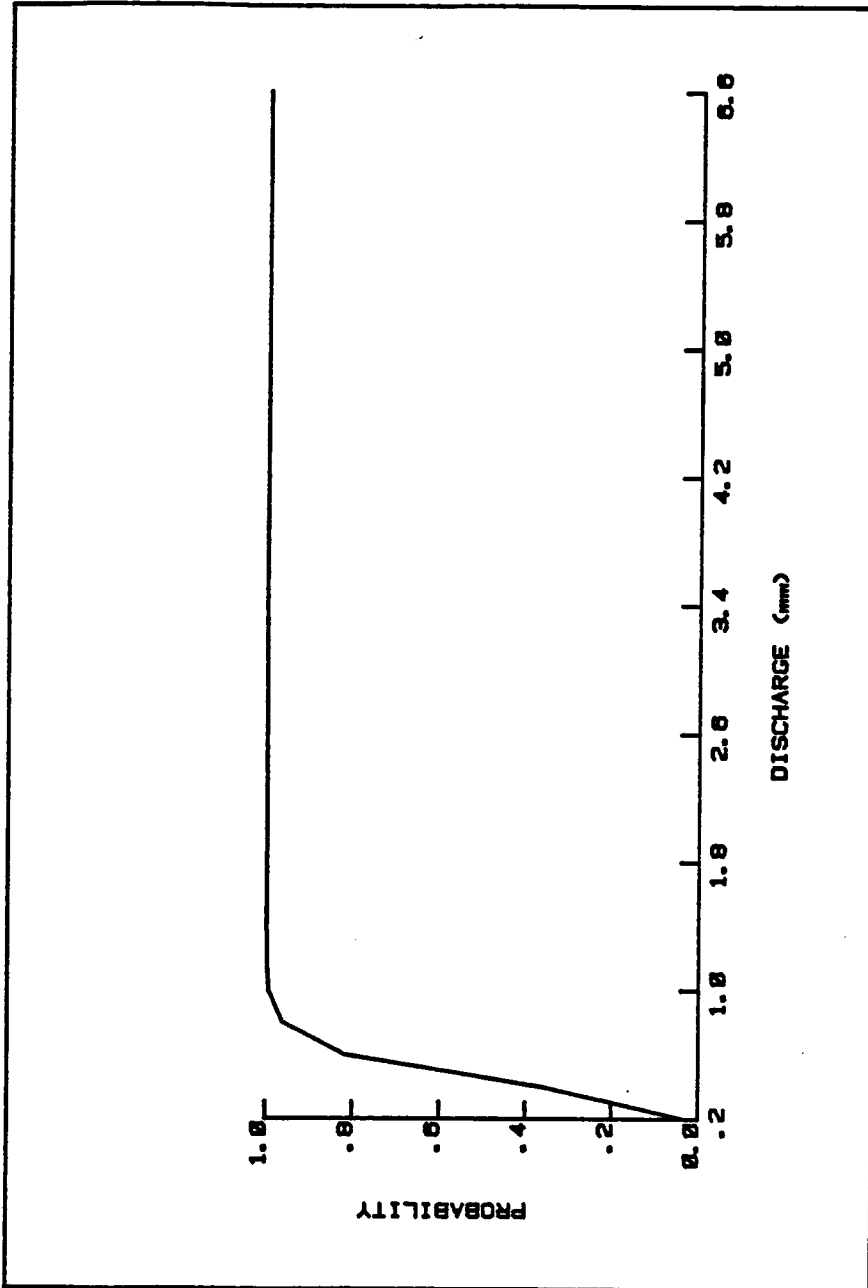


Figure 17. Probability of Total Discharge of the Tinau River for March.

percent. The simulated total annual discharge was 165 mm for the driest year and for the wettest year 2,375 mm. The annual flows were equal to or less than 1,660 mm 90 percent of the time and 50 percent of the time equal to or less than 1,040 mm. The probability of an annual discharge of 360 mm or less was only 5 percent. A summary of total seasonal and annual discharge occurrence probability is presented in Table 6, and the cumulative probability frequencies are illustrated in Figure 18 and 19.

Maximum and Minimum Discharge

The amount and timing of peak discharge are affected by the flow during the monsoon season. The probability of peak flow can be used to determine the design and stability of flood control structures in the plains, the safety of water supply for human and agriculture use, and in problem of power generation. A knowledge of low flow probabilities aids in optimizing the operation of power plant, water supply and irrigation.

Annual maximum discharge and minimum discharge were evaluated for the 100 simulation period. The maximum discharge noted was 4,440 m³/s. The actual highest flows recorded at Butwal were 3,300 m³/s in 1970, and 3,500 m³/s in 1981. Simulated results indicated that 50 percent of the time maximum flows should be equal to or less than 840 m³/s. The probability of exceeding 2040 m³/s was found to be only 10

Table 6. Discharge and Low Flow Days Probability of the Tinau River, Butwal.

		Probability of Occurrence									
		- percent -									
		10	20	30	40	50	60	70	80	90	99
Total Discharge Equal to or Less Than											
		- mm -									
Monsoon (JUN-SEP)		300	440	560	670	740	890	1020	1170	1350	1940
Total Annual		475	710	830	950	1040	1180	1305	1490	1660	2310
Daily Storm Flow Equal to or Less Than											
		- cubic meter per sec -									
Annual Maximum Flow		145	360	510	645	840	1080	1270	1560	2040	4200
Annual Minimum Flow		0.91	1.07	1.18	1.30	1.38	1.50	1.64	1.78	1.88	2.31
Consecutive Days of Flows Less Than 1.5 cu. m. per sec.											
		2.6	4.0	8.3	11.5	16.5	22.0	25.7	33.7	51.5	120

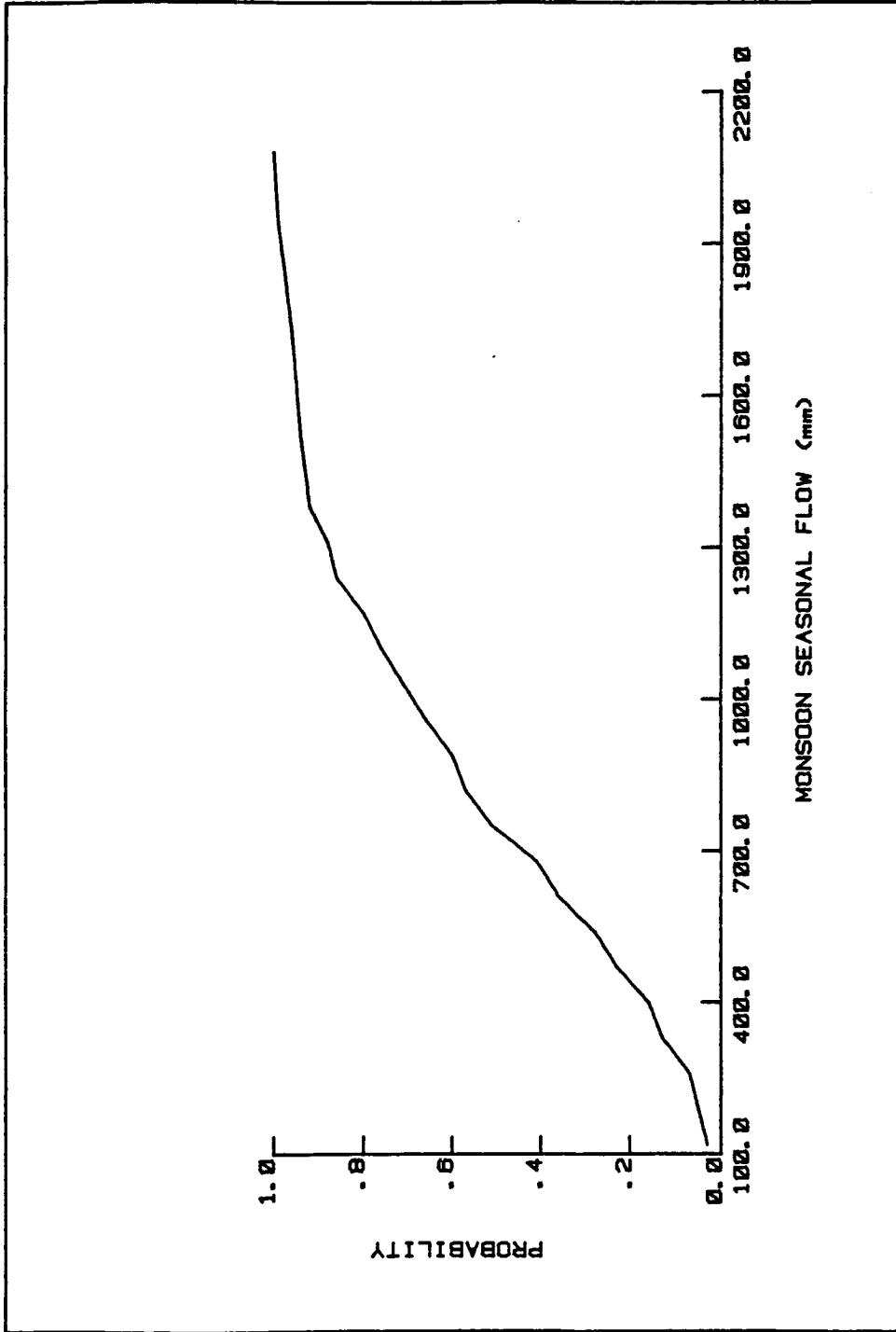


Figure 18. Probability of Monsoon Seasonal Discharge from the Tinau Watershed, Palpa.

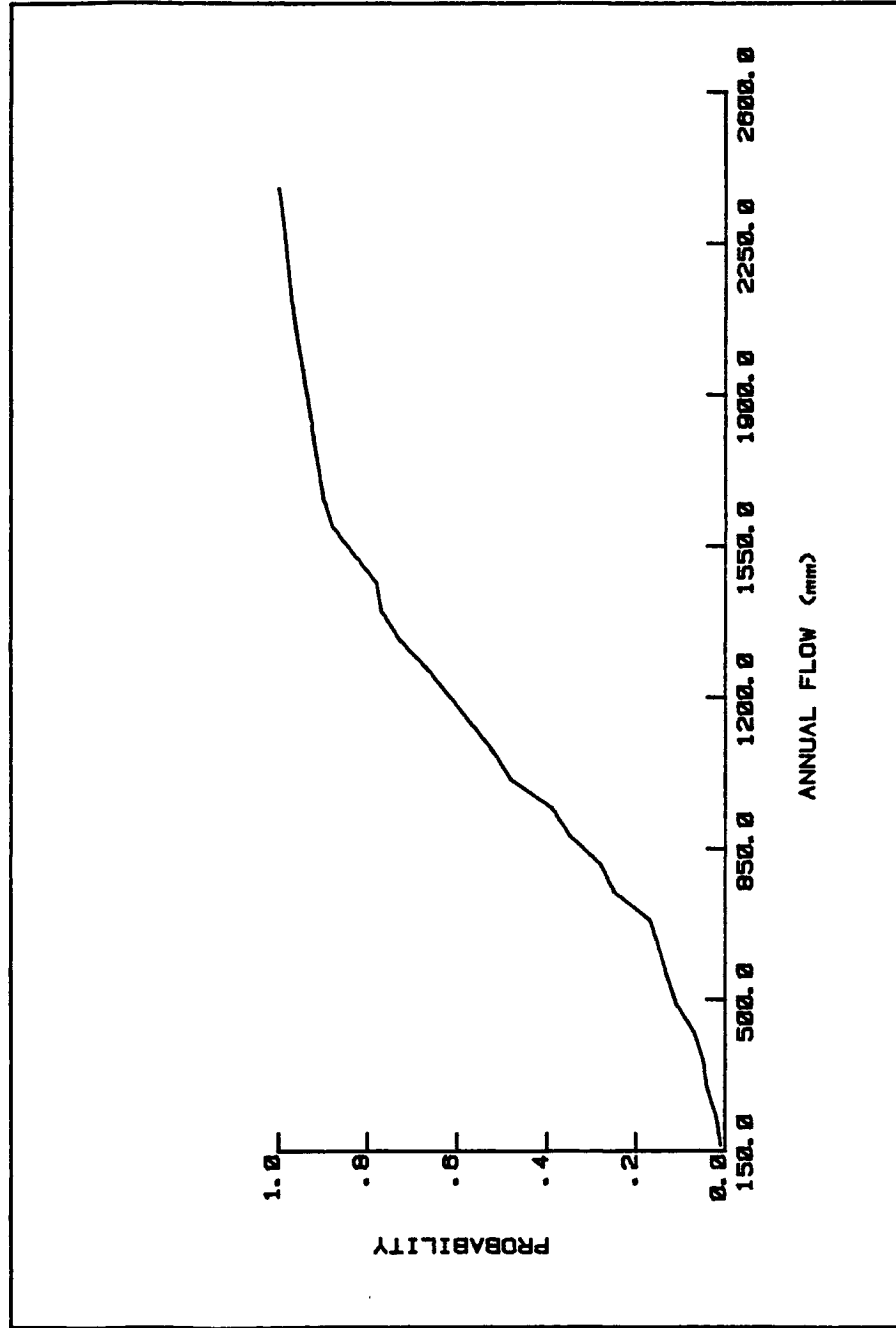


Figure 19. Probability of Annual Discharge from the Tinau Watershed, Palpa.

percent. A probable maximum discharge of only 120 m³/s occurred in the driest year of the simulation. Cumulative frequency curves of annual maximum discharge are given in Figure 20. Discharges equal to or less than 2.5 m³/s were considered to be low flows. Low flow occurred mainly in late March, April, May, and early June. The probable minimum flow noted for the simulation period was 0.253 m³/sec, the actual low flow of the historical record was 0.94 m³/s in 1981.

The simulation indicated that 90 percent of the time the annual minimum flow is less than 1.88 m³/s. A cumulative probability frequency of the annual low flows is illustrated in Figure 21. The probabilities of occurrence of annual maximum and minimum flows equal to or less than a given amount are presented in Table 6. A summary of simulation results of high and low discharges are listed in Appendix C.

The peak of a storm flow could be much higher than the maximum daily flow as analyzed here. Thus, flood control devices and other water utilization structures should be designed and constructed accordingly. An alternative safe water supply scheme may be designed to share low flow with power plant and irrigation in the dry period.

Consecutive Days of Low Flow

Consecutive days of low flow are assumed to be those periods when the daily stream discharge is less than 1.5 m³/s. Although a constant flow of 1.5 m³/s is satisfactory

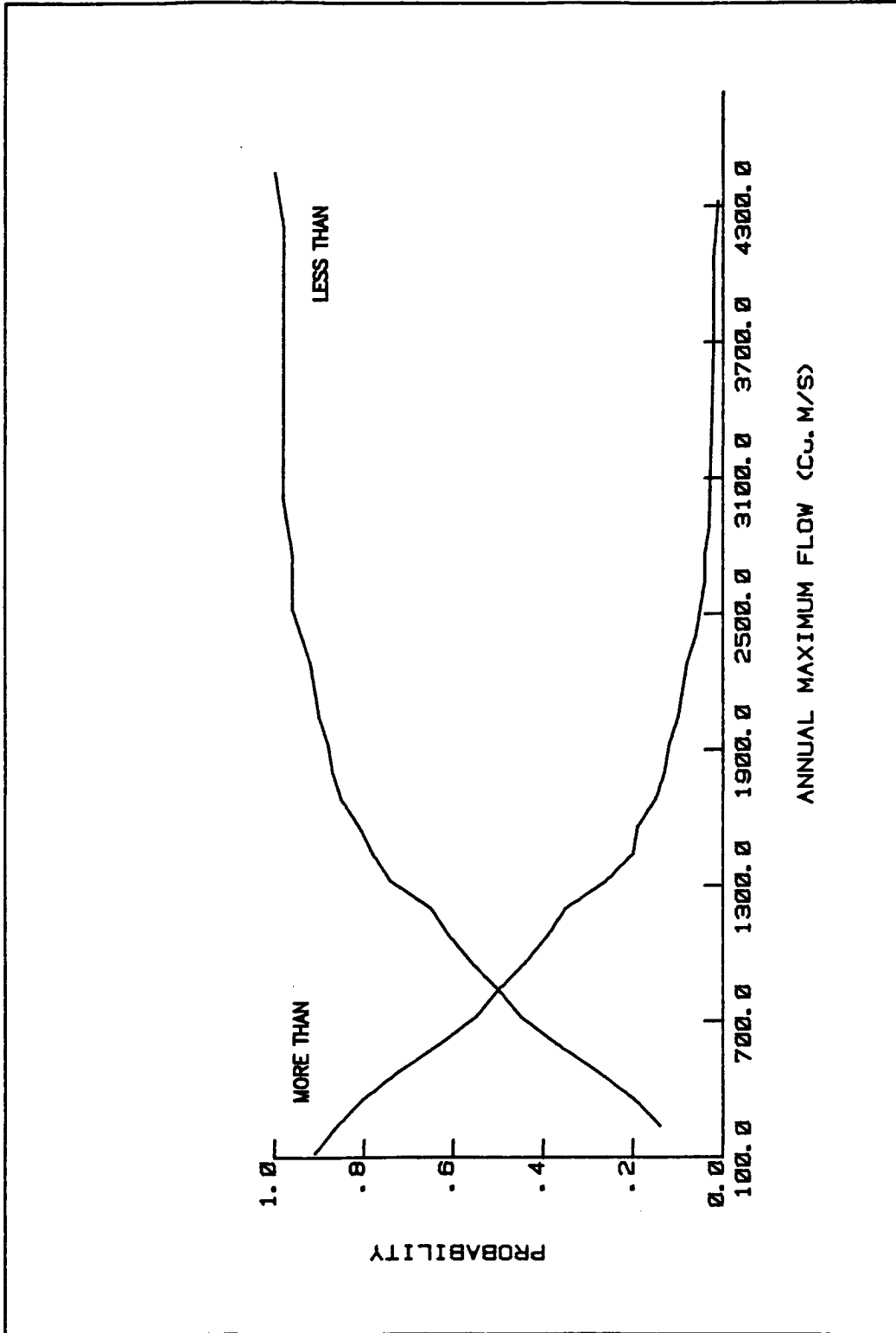


Figure 20. Probability of Annual Maximum Flows of the Tinnau River, Butwal

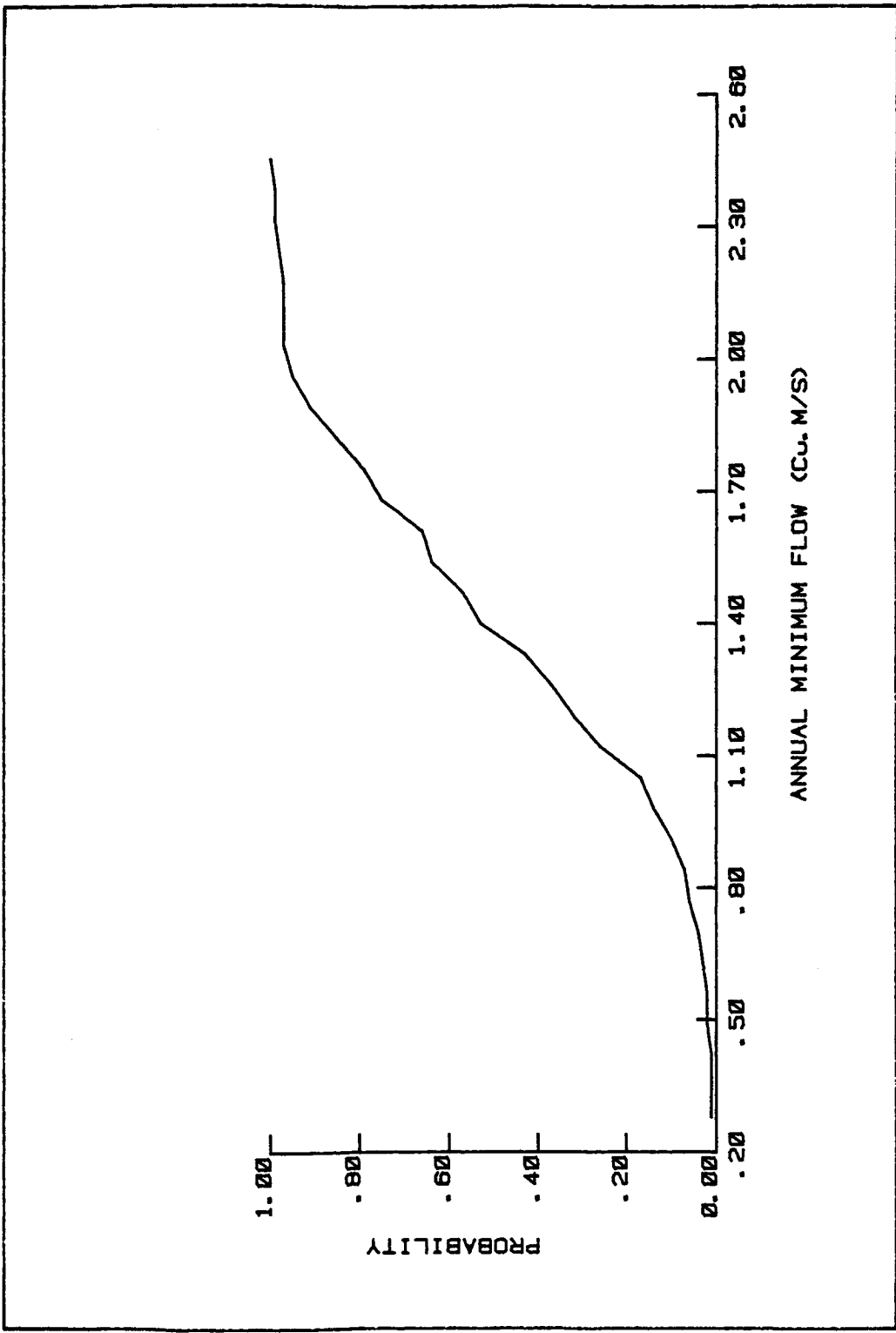


Figure 21. Probability of Annual Minimum Flows of the Tinau River, Butwal.

to maintain power production downstream at Butwal, flows less than this could affect the operation of the hydropower station, irrigation schedules and water supplies in the plains. A summary of probable consecutive days of low flow is given in Table 6 and the cumulative probability frequency is illustrated in Figure 22.

Simulation results showed that in 100 years, 1930 days have streamflow less than $1.5 \text{ m}^3/\text{s}$, 5.3 percent of the time. The mean consecutive days of low flow was 24.42 with a maximum of 157 days for the 100 year simulation. It was noted that 90 percent of the time the duration of low flows was 51 days or less. Most of the low flow occurred in late March, April, May and in early June. However, in relatively dry simulation year low flow began in mid January and February, and in years when monsoon started late, low flows began in May and extended until July. The actual flow data showed that if rainfall were high in the post-monsoon months flows below $2.5 \text{ m}^3/\text{s}$ were produced only from late May until late summer when the monsoon came early and low flow began in early December.

Application of the Model

The watershed model TINAU can be applied to a variety of problems in managing the Tinau watershed, Palpa. The event-based approach presented in the model is a hydrologic tool to assess storm discharge and sediment yield from the

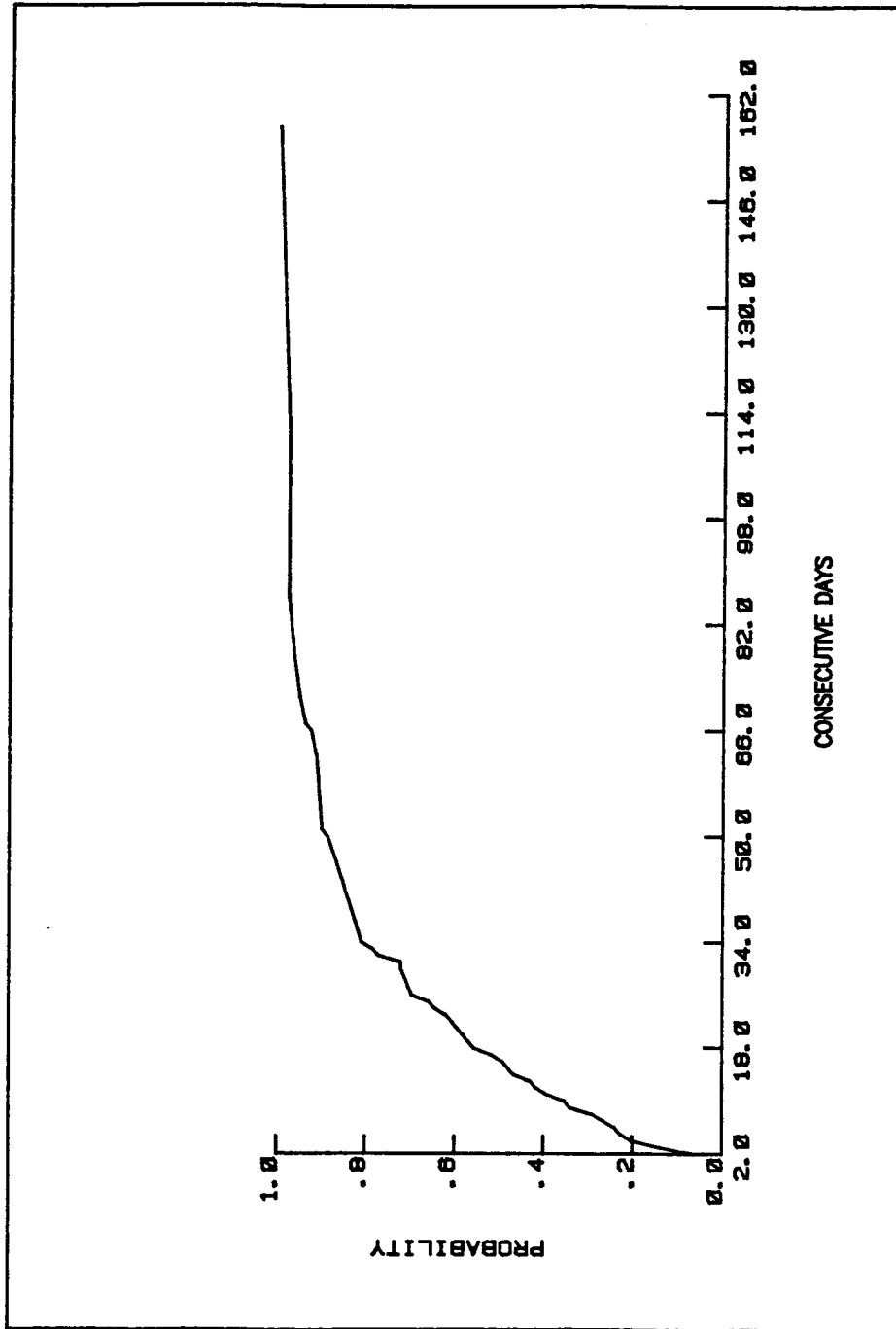


Figure 22. Probability of Consecutive Days of Low Flows of the Tinnau River, Butwal.

watershed. Watershed management and water utilization activities in the three main sub-catchments Jhumsa, Dovan and Madi are critical questions. The TINAU model can be operated in these three subcatchments separately to evaluate water and sediment yield, extreme events, and consecutive days of low flow. The model was designed to operate in snow free area and may be used in the various activities listed below:

a) Flood control activities: The Tinau river creates flood damages in the Butwal area. The maximum daily discharge from the model output can be used to design flood control structures. The model can also be used as a tool to determine risks in flood forecasting and damage assessment process.

b) Sediment yield: The model generates daily runoff from the watershed. Due to unavailability of rainfall duration information, the model was not designed to evaluate instantaneous peak discharge. If adequate rainfall data were available, the modified universal soil loss equation (Williams, 1975), which uses storm discharge and peak flow variables to estimate storm sediment yield, could be applied. Information on rainfall duration could be used to evaluate peak flows, and watershed characteristics applied to evaluate other parameters required for sediment yield estimation. It is possible to link a sediment yield sub-model

with the TINAU to assess sediment yield from the watershed. This information could be used to give land management practices recommendations toward minimizing soil loss from the watershed.

c) Water yield: The annual, seasonal, and monthly water yield under present land use practices are discussed above. A monthly assessment of water yield can be applied in planning, design, and operation of land treatment, vegetation management, and water utilization practices. The precipitation model could also be coupled with an irrigation model to obtain optimum scheduling of irrigation water for long-term scheduling. Coupled with a drainage model an economic and efficient drainage system could be designed for operation in the Madi plain.

d) Extreme events: The model permits estimates of extreme events and extreme sequences which are not yet recorded in the actual streamflow data. The model provides sequences of extremely low flows, which would help in designing, and planning surface irrigation system in the Butwal area, as well as operation of the hydropower plant in Butwal.

e) Small water impounding structures: One application of the model could in designing regional stock ponds. The model can be regionalized for any part of the watershed. Pond and spillway could be designed to utilize maximum flows and conserve water for dry periods.

CHAPTER SEVEN

SUMMARY AND CONCLUSION

This study presents a methodology for evaluating the hydrologic performance of small watersheds in Nepal. The methodology was intended to be used by decision makers for watershed management practices for both upstream and downstream benefits. The most practical use of the methodology could be in estimating erosion losses, water management and utilization. The model developed has particular application to the Madi, Jhumsa and Dovan valleys, and flood control measures in the Butwal area.

The methodology is in the form of a coupled stochastic and deterministic computer simulation model called TINAU. The procedure followed in the model involves:

(1) the use of the stochastic precipitation model to simulate synthetic time series of rainfall data;

(2) the use of the deterministic hydrologic model to transform the synthetic rainfall record to a long-term streamflow records; and

(3) the use of a watershed model to evaluate the hydrologic performance.

A coupled stochastic and deterministic model was utilized in this study, because stochastic simulation

creates a time series of synthetic rain fall data and allows one to view an ensemble of possible regionalization of precipitationrunoff regimes at a given site. A model that utilizes only historical data provides an unique and limited set of results. The stochastically generated time series of rainfall record include a wide spectrum of amounts and sequential patterns resulting in the probabilistic distribution range of streamflow responses.

In this study, the objective was to develop a methodology that is generally applicable, uses readily available data and operates in a simple but physically consistent manner. The TINAU model only requires daily precipitation, temperature, evaporation, and streamflow data. Data used to calibrate the parameters required by the model were obtained from the Tansen and Butwal stations.

The stochastic precipitation sub-model, which drives the watershed model, involved two main assumptions: that rainfall patterns in the monsoon season (JUN-SEP) occur in sequences; and independent thunderstorms occur in the winter and summer months. The random variables used in the model were rainfall events (groups in the monsoon season) and rainfall amount per event. The model synthesizes a time series record of daily rainfall.

The event-based, deterministic hydrologic sub-model transforms the synthetic rainfall data into daily streamflow. The hydrologic sub-model DOVAN was a generalized,

fitted parameter model. Parameters for the DOVAN model were calibrated with the daily streamflow records in the Tinau river at Butwal, and with evaporation and temperature data from Tansen and Butwal.

The watershed model TINAU was run for 100 years. The simulated mean annual runoff was 1090.6 mm with a mean maximum daily discharge of 1010.4 m³/s. High pre-monsoon rainfall produced an early commencement of baseflow in some year. Late monsoon rainfall caused long baseflow recessions throughout the summer months, March April and May. The mean consecutive days of low flow during this period was 24.42. Periods of low flow days extended as long as 157 days. April and May were the driest months with low flows falling as low as 0.10 mm. Winter flows were satisfactory with mean flow of 0.9 mm. August and September were the wettest months with an average flow of 11.6 mm.

Some limitations were observed in the watershed model. In the stochastic precipitation model, daily rainfall amount was evaluated by distributing total rainfall amount per group among the number of days in a group. So, extreme rainfall events could not be considered for the assessment of peak flows. However, the actual highest daily flow recorded was 3500 m³/s in 1981. The simulated result showed a highest daily flow of 4440 m³/s. Another limitation of the model was low baseflow production. The actual baseflow was

0.95 m³/s in 1981, but the simulated baseflow was as low as 0.253 m³/s.

The TINAU model was calibrated with only 8 years of historical data. The model could be recalibrated, tested, and validated with additional rainfall and runoff data when it becomes available. Even so, the output of the precipitation model developed closely matched the quantity and variability of the actual rainfall measured in the critical period of the monsoon season. A more reliable output would be expected if a longer historical record were available. The use of the output of the model will depend upon the objective of the user. The model may be used as a part of a comprehensive watershed model in decision making process. The model has the advantage of allowing decisions to be made based upon the risks the user is willing to take.

APPENDIX A
COMPUTER PROGRAM LISTING OF TINAU
SIMULATION MODEL, INCLUDING SUBROUTINES

```

0001 FTN7X.L
0002 *FILES9,9
0003
0004     PROGRAM TINAU
0005
0006 *****
0007 *   THIS IS A WATERSHED HYDROLOGIC MODEL DEVELOPED FROM A COUPLE *
0008 *   OF STOCHASTIC PRECIPITATION MODEL AND A DETERMINISTIC *
0009 *   STREAMFLOW MODEL. THIS MODEL IS CALIBRATED FROM CLIMATOLOGICAL *
0010 *   AND HYDROLOGICAL DATA FROM TANSEN AND BUTWAL. *
0011 * *
0012 *   STATIONS : TANSEN, and BUTWAL *
0013 *   DATA YEAR: CLIMATOLOGICAL (1972-1985) *
0014 *   HYDROLOGIC (1977-1982) *
0015 *   PROGRAM DEVELOPED BY: RABIN BOGATI *
0016 *   SCHOOL OF RENEWABLE NATURAL RESOURCES *
0017 *   THE UNIVERSITY OF ARIZONA *
0018 *   TUCSON, ARIZONA 85721 *
0019 *   1985-1986 *
0020 *****
0021 *
0022 *----- DECLARATION OF VARIABLES
0023
0024 *   REAL    FQJUN(40),FQJUL(40),FQAUG(40),FQSEP(40)
0025 *   REAL    FQOCT(40),FQNOV(40),FQDEC(40),FQJAN(40)
0026 *   REAL    FQFEB(40),FQMAR(40),FQAPR(40),FQMAY(40)
0027 *   REAL    FQMSUN(40),FQYRSUM(40),FQYRMX(75),FQYRMIN(40)
0028 *   REAL    MSOONQ, YRSUMQ, YRMXFLO, YRMINFLO
0029 *   CHARACTER CALENDAR(12)*10
0030
0031 *   COMMON /A/ III, ILEAP, PPT(366), FLOWTOT(366), RUNOFF(366), LU
0032
0033 *   COMMON /B/ FQJUN(40), FQJUL(40), FQAUG(40), FQSEP(40), FQOCT(40),
0034 *   &FQNOV(40), FQDEC(40), FQJAN(40), FQFEB(40), FQMAR(40), FQAPR(40),
0035 *   &FQMAY(40), FQMSUN(40), FQYRSUM(40), FQYRMX(40), FQYRMIN(40)
0036
0037 *   COMMON /C/ MSOONQ, YRSUMQ, YRMXFLO, YRMINFLO
0038
0039 *   DATA CALENDAR/'JUNE','JULY','AUGUST','SEPTEMBER','OCTOBER',
0040 *   &'NOVEMBER','DECEMBER','JANUARY','FEBRUARY','MARCH',
0041 *   &'APRIL','MAY'/
0042 *----- Initialize years of simulation and output
0043

```

```

0044      WRITE(1,*) 'HOW MANY YEARS YOU WANT TO BE SIMULATED?'
0045      WRITE(1,*) 'TYPE IN NUMBER AND HIT RETURN.'
0046      READ(1,*) LBARSA
0047      WRITE(1,*) 'THIS MODEL IS SIMULATING FOR:',LBARSA,'(YEARS)'
0048      WRITE(1,*) 'WHAT IS THE OUTPUT DEVICE NUMBER 1=? OR 6=?'
0049      READ(1,*) LU
0050
0051 C      OPEN(123, FILE='BARSA', STATUS = 'NEW')
0052 C      OPEN(234, FILE='BHEL', STATUS='NEW')
0053 C      OPEN(345, FILE='MINBHEL',STATUS='NEW')
0054
0055      WRITE(1,*)'***** PROGRAM IS RUNNING *****'
0056      *
0057      *----- Initialize the RANDOM NUMBER generation
0058      *                          and the addition variables.
0059      CALL SSEED(12345)
0060
0061      DO 15 N = 1, 40
0062          FQJUN(N) = 0.
0063          FQJUL(N) = 0.
0064          FQAUG(N) = 0.
0065          FQSEP(N) = 0.
0066          FQOCT(N) = 0.
0067          FQNOV(N) = 0.
0068          FQDEC(N) = 0.
0069          FQJAN(N) = 0.
0070          FQFEB(N) = 0.
0071          FQMAR(N) = 0.
0072          FQAPR(N) = 0.
0073          FQMAY(N) = 0.
0074          FQMSUN(N) = 0.
0075          FQYRSUM(N) = 0.
0076          FQYRMX(N) = 0.
0077          FQYRMIN(N) = 0.
0078      15  CONTINUE
0079
0080      *----- THE YEARLY LOOP BEGINS *****
0081
0082          DO 9999   III = 1, LBARSA
0083
0084      *                          ! Check the LEAPYEAR value
0085
0086          LEAP = MOD(III,4)
0087          IF(LEAP .EQ. 0) THEN
0088              ILEAP = 1
0089          ELSE
0090              ILEAP = 0
0091          END IF

```

```

0092
0093 * ----- Operate the Stochastic PRECIPITATION model
0094 *           to simulate one year series of rainfall data
0095
0096         CALL PANI      ! A Stochastic PRECIPITATION model
0097
0098 * ----- Operate the Streamflow generation model
0099 *           to transform one year series of rainfall to runoff
0100 *
0101 * ----- Initialize PEAK and TOTAL flows to be computed
0102         MSOONO = 0.0
0103         YRSUMQ = 0.0
0104         YRMXFLO = 0.0
0105         YRMINFLO = 2.5 ! Cubic meter per second mean daily flow
0106
0107         CALL DOVAN      ! The G.S.S.S. Hydrologic Model
0108
0109         9999 CONTINUE
0110
0111 * ----- ! Report the monthly, annual and storm flow
0112 *
0113         WRITE(LU,20) CALENDAR(1)
0114         CALL REPORT(FQJUN,0.,14.,LU,40) ! Report monthly frequency JUNE
0115         WRITE(LU,20) CALENDAR(2)
0116         CALL REPORT (FQJUL,0.,13.,LU,40) ! JULY
0117         WRITE(LU,20) CALENDAR(3)
0118         CALL REPORT (FQAUG,0.,17.,LU,40) ! AUGUST
0119         WRITE(LU,20) CALENDAR(4)
0120         CALL REPORT (FQSEP,0.,12.,LU,40) ! SEPTEMBER
0121         WRITE(LU,20) CALENDAR(5)
0122         CALL REPORT (FQOCT,0.,10.,LU,40) ! OCTOBER
0123         WRITE(LU,20) CALENDAR(6)
0124         CALL REPORT (FQNOV,0.,0.3,LU,40) ! NOVEMBER
0125         WRITE(LU,20) CALENDAR(7)
0126         CALL REPORT (FQDEC,0.,0.60,LU,40) ! DECEMBER
0127         WRITE(LU,20) CALENDAR(8)
0128         CALL REPORT (FQJAN,0.,0.3,LU,40) ! JANUARY
0129         WRITE(LU,20) CALENDAR(9)
0130         CALL REPORT (FQFEB,0.,0.25,LU,40) ! FEBRUARY
0131         WRITE(LU,20) CALENDAR(10)
0132         CALL REPORT (FQMAR,0.,0.3,LU,40) ! MARCH
0133         WRITE(LU,20) CALENDAR(11)
0134         CALL REPORT (FQAPR,0.,0.35,LU,40) ! APRIL
0135         WRITE(LU,20) CALENDAR(12)
0136         CALL REPORT (FQMAY,0.,1.5,LU,40) ! MAY
0137

```

```

00138 20  FORMAT(/////, 10X, ' MONTHLY SIMULATION RUNOFF FREQUENCY',/
00139      &,10X,' SUMMARY RESULT FOR THE MONTH - '.A10./,
00140      &10X,'          RUNOFF IN WATERSHED MM',//)
00141 * ----- Report total annual Monsoon flow
00142      WRITE(LU,30)
00143      CALL REPORT (FQMSUN, 50.0,70.,LU,40)
00144 30  FORMAT(///,12X,'SUMMARY OF TOTAL MONSOON SEASON ',/
00145
00146 *          ! Report total annual streamflow
00147      &,12X,' RUNOFF VOLUME IN MM',//)
00148      WRITE(LU,40)
00149      CALL REPORT (FDYRSUM, 100.,65.,LU,40)
00150 40  FORMAT(///,10X,'SUMMARY OF TOATL YEARLY RUNOFF VOLUME IN MM',//)
00151
00152 *          ! Report maximum annual stormflow in cubic meter per sec.
00153
00154      WRITE(LU,50)
00155
00156 50  FORMAT(///,10X,'SUMMARY OF MAXIMUM DAILY FLOW IN CU. M/SEC',//)
00157      CALL REPORT (FQYRMX,0.0 .120.,LU,40)
00158
00159 *----- ! Report minimum annual stormflow
00160      WRITE(LU,60)
00161      CALL REPORT (FQYRMIN,0.0,.07,LU,40)
00162 60  FORMAT(///,10X,' SUMMARY OF EXTREMELY LOW FLOWS IN CU.M/SEC',/,
00163      &10X,' (Flows less than 2.5 cubic meter per sec )',//)
00164      WRITE(1,*) ' PURGE *BARSA* AND *BHEL* BEFORE RERUNNING PROGRAM'
00165
00166      STOP
00167      END
00168
00169
00170
00171

```

```

0172
0173     SUBROUTINE PANI
0174     *****
0175     * A PRECIPITATION SIMULATION MODEL *
0176     * THIS SUBPROGRAM SIMULATES DAILY RAINFALL DATA FOR ONE YEAR AT *
0177     * A TIME STARTING FROM JANUARY 1 AND ENDS IN DECEMBER (JULIAN CALENDAR) *
0178     * THE RAINFALL MODEL IS A STOCHASTIC MODEL, BASED ON STATISTICAL PROBABA- *
0179     * BILITY, THAT TREATS MONTHS OF JUNE THROUGH SEPTEMBER AS A SERIES *
0180     * OF A FRONTAL STORM AND OCTOBER THROUGH MAY AS INDEPENDENT THUNDER - *
0181     * STORM RAINFALL PATTERNS. THUNDERSTORM MONTHS ARE FURTHER SUBDIVIDED *
0182     * IN TO STATISTICALLY NON DIFFERENT GROUPS. *
0183     * *
0184     *     MONSOON = JUNE, JULY, AUGUST, SEPTEMBER *
0185     * *
0186     *     Rainfall event interarrival time for OCT-MAY (IT) *
0187     *     GROUP 1 = JANUARY, FEBRUARY, MARCH (gamma) *
0188     *     GROUP 2 = APRIL, MAY (geometric) *
0189     *     GROUP 3 = OCTOBER (gamma) *
0190     *     GROUP 4 = NOVEMBER, DECEMBER (gamma) *
0191     * *
0192     *     Rainfall amount per event (PPT/EVENT) *
0193     *     GROUP 1 = JANUARY, FEBRUARY, MARCH, APRIL, MAY (gamma) *
0194     *     GROUP 2 = OCTOBER, NOVEMBER (gamma) *
0195     *     GROUP 3 = DECEMBER (gamma) *
0196     *****
0197
0198     REAL K(8), LAMDA(8), MEAN(5), PPTMON(5)
0199     REAL TEVENT(12), DAY(12), CDF(46), DBS(70), NGS(35)
0200     REAL RAIN, LNGP(25), DAYBC(10), GRAIN, CLASSRAIN
0201     INTEGER DCOUNT
0202     REAL JAL(12), MONSOON, TEMP
0203     REAL XEVENT(5)
0204
0205     COMMON /A/ III, ILEAP, PPT(366), FLOWTOT(366), RUNOFF(366), LU
0206     * THESE ARE THE STATISTICAL PARAMETER VALUES USED IN SIMULATION.
0207
0208     DATA LAMDA/.1593, .0621, .0285, .11030, .0490, .1504, .0672, .134/
0209     DATA K/1.849, .2832, .4841, .8873, .5236, 2.6308, .5648, .761/
0210     DATA MEAN/4.625, 3.730, 5.481, 1.813, 3.722/
0211     DO JL = 1, 366
0212         PPT(JL) = 0.0
0213     END DO
0214
0215
0216     JDAY = 0
0217     DCOUNT = 0
0218     RAIN = 0.0
0219

```

```

0220 ***** MODEL FOR JAN,FEB,MAR *****
0221
0222     JDAY = 90 + ILEAP
0223     DO WHILE(DCOUNT .LE. JDAY)
0224         X =URAN( )           ! a RANDOM NUMBER generation
0225         IF(X .LE. 0.4103) THEN ! shifted distribution
0226             NINT = 1
0227         ELSE
0228             NINT =INT(GAMMA(K(1),1./LAMDA(1))+1.)+1
0229         END IF
0230
0231         DCOUNT = DCOUNT + NINT
0232
0233         IF(DCOUNT .GT. JDAY) GO TO 20
0234         PPT(DCOUNT) = GAMMA(K(4),1./LAMDA(4))
0235         IF(PPT(DCOUNT) .LT. .1) PPT(DCOUNT) = .00
0236         IF(DCOUNT .GE. JDAY) GO TO 20
0237     20     END DO
0238
0239 ***** MODEL FOR THE MONTH APR-MAY *****
0240
0241     JDAY = 151 + ILEAP
0242     DO WHILE(DCOUNT .LE. JDAY)
0243         X = URAN( )
0244         IF(X .LE. 0.4766) THEN ! a shifted distribution
0245             NINT = 1
0246         ELSE
0247
0248             X = URAN( )
0249             P = 1./MEAN(1)
0250             NINT = INT((ALOG(1.-X))/(ALOG(1.-P))+1.) + 1
0251         END IF
0252
0253         DCOUNT = DCOUNT + NINT
0254         IF(DCOUNT .GT. JDAY) GO TO 30
0255
0256         IF(DCOUNT .LE. (120+ILEAP)) THEN
0257             PPT(DCOUNT) = GAMMA(K(4),1./LAMDA(4))
0258         ELSE
0259             PPT(DCOUNT) = GAMMA(K(5),1./LAMDA(5))
0260         END IF
0261         IF(PPT(DCOUNT) .LT. .10) PPT(DCOUNT) = .00
0262         PPTMON(2) = PPTMON(2) + PPT(DCOUNT)
0263         TEVENT(2) = TEVENT(2) + 1
0264
0265         IF(DCOUNT .GE. JDAY) GO TO 30
0266     30     END DO
0267

```

```

0268 ***** MONSOON SEASONAL MODEL *****
0269
0270         JDAY = 273 + ILEAP
0271         ITRANSIT = 1
0272         DO WHILE (DCOUNT .LE. JDAY)
0273             IF (ITRANSIT .EQ. 1) GO TO 31
0274 ***** Dry days between sequences (DBS)
0275
0276             X = URAN( )
0277             NINT = INT((ALOG(1.-X))*(-MEAN(5))) + 5
0278             TNINT = TNINT + 1
0279             DBS(NINT) = DBS(NINT) + 1
0280
0281             TOTDBS = TOTDBS + NINT ! between sequences.
0282
0283             DCCOUNT = DCCOUNT + NINT
0284             IF(DCCOUNT .GT. JDAY) GO TO 39
0285 ***** Groups per sequence (NGRPS)
0286     31     ITRANSIT = 0
0287
0288             NGRPS=INT(GAMMA(K(8),1./LAMDA(8)) + 1.)
0289
0290 ***** A computation within the group
0291
0292             DO 40 JJ = 1, NGRPS
0293
0294 **                 Groupduration in days (LG) ***
0295
0296                 X = URAN( )
0297                 P = 1./MEAN(2)
0298                 LG = INT((ALOG(1.-X))/(ALOG(1.-P)) + 1.)
0299                 TOTLG = TOTLG + LG ! count the group duration
0300
0301 **                 Rainfall amount per group (RP) ***
0302
0303                 CLASSRAIN = GAMMA(K(7),1./LAMDA(7)) ! simulate class
0304
0305                 GRAIN = (CLASSRAIN*10.- 5.) ! Translate the RAINFALL
0306                 MONSOON = MONSOON + GRAIN
0307                 MONEVENT= MONEVENT + 1
0308
0309             DO 45 I = 1, LG
0310                 DCCOUNT = DCCOUNT + 1
0311                 PPT(DCCOUNT) = GRAIN/LG
0312                 PPTMON(3) = PPTMON(3) + PPT(DCCOUNT)
0313                 IF(PPT(DCCOUNT) .LT. .1) PPT(DCCOUNT) = 0.0
0314     45     CONTINUE
0315
0316             IF(DCCOUNT .GE. JDAY) GO TO 39
0317

```

```

0318 **      Dry days between groups   (DBG)  ****
0319
0320 32      X = URAN( )
0321      P = 1./MEAN(4)
0322      DBG = INT (ALOG(1.-X)/(ALOG(1.-P))+1.)
0323      IF (DBG .GT. 6) GO TO 32
0324      DAYBG(DBG) = DAYBG(DBG) + 1
0325
0326      TOTDBG = TOTDBG + DBG          ! Count DBG
0327
0328      DCOUNT = DCOUNT + DBG        ! Added length of dry day too.
0329
0330      IF(DCOUNT .GE. JDAY) GO TO 39
0331
0332 40      CONTINUE
0333 39      END DO
0334
0335 ***** THE MODEL FOR OCTOBER *****
0336
0337      JDAY = 304 + ILEAP
0338      DO WHILE (DCOUNT .LE. JDAY)
0339          NINT = INT(GAMMA(K(2),1./LAMDA(2))+1)
0340          DCOUNT = DCOUNT + NINT
0341          IF(DCOUNT .GT. JDAY) GO TO 50
0342
0343          PPT(DCOUNT) = GAMMA(K(5),1./LAMDA(5))
0344          IF(PPT(DCOUNT).LT. .1) PPT(DCOUNT) = .0
0345
0346          IF(DCOUNT .GE. JDAY) GO TO 50
0347 50      END DO
0348
0349 ***** THE MODEL FOR NOV-DEC *****
0350
0351      JDAY = 365+ILEAP
0352      DO WHILE (DCOUNT .LE. JDAY)
0353          NINT = INT(GAMMA(K(3),1./LAMDA(3))+1.)
0354          DCOUNT = DCOUNT + NINT
0355          IF(DCOUNT .GT. JDAY) GO TO 60
0356
0357          IF(DCOUNT .LE.(334+ILEAP)) THEN
0358              PPT(DCOUNT) = GAMMA(K(4),1./LAMDA(4))
0359          ELSE
0360              PPT(DCOUNT) = GAMMA(K(6),1./LAMDA(6))
0361          END IF
0362          IF(PPT(DCOUNT) .LT. .1) PPT(DCOUNT) = .0
0363          IF(DCOUNT .GE. JDAY) GO TO 60
0364
0365 60      END DO
0366

```

```

0367 *----- CONVERT THE CALENDAR
0368 * JUNE 1 TO MAY LAST
0369 DO I = 1, 151+ILEAP
0370     TEMP = PPT(I)
0371     DO J = 2, 365+ILEAP
0372         PPT(J-1) = PPT(J)
0373     END DO
0374     PPT(365) = TEMP
0375 END DO
0376 *----- COUNT THE YEARLY RAINFALL
0377     TOTPPT = 0.
0378     DO M = 1, 365+ILEAP
0379         TOTPPT = TOTPPT + PPT(M)
0380     END DO
0381 C     WRITE(123,*)III, TOTPPT
0382     RETURN
0383     END
0384
0385 *****
0386 * GAMMA FUNCTION
0387 * THIS FUNCTION WAS DEVELOPED TO EVALUATE GAMMA CDF
0388 *****
0389
0390     FUNCTION GAMMA(ALPHA,BETA)
0391
0392     IF (ALPHA .GT. 1.) GO TO 222
0393     B = 1.+ALPHA/2.17828183
0394 111 U = URAN( )
0395     P = B*U
0396     U = URAN( )
0397
0398     IF (P .GT. 1.) THEN
0399         Y = -ALOG((B-P)/ALPHA)
0400         IF (U .LE. Y**(ALPHA-1.)) THEN
0401             GAMMA = Y*BETA
0402             RETURN
0403         ELSE
0404             END IF
0405     ELSE
0406         Y = P**(1./ALPHA)
0407         IF (U .LE. EXP(-Y)) THEN
0408             GAMMA = Y*BETA
0409             RETURN
0410         ELSE
0411             END IF
0412     END IF
0413

```

```
0414          GO TO 111
0415 222      A = 1./SQRT(2.*ALPHA-1.)
0416          B = ALPHA-ALOG(4.)
0417          Q = ALPHA+1./A
0418          E = 4.5
0419          D = 1.+ALOG(E)
0420 333      U1 = URAN()
0421          U2 = URAN()
0422          V = A*ALOG(U1/(1.-U1))
0423          Y = ALPHA*EXP(V)
0424          Z = U1*U1*U2
0425          W = B+Q*V-Y
0426          IF(W+D-E*Z .GE. 0. .OR. W .GE. ALOG(Z)) THEN
0427              GAMMA= Y*BETA
0428              RETURN
0429          ELSE -
0430              END IF
0431          GO TO 333
0432          END
0433
0434
0435
0436
0437
0438
```

```

0439      SUBROUTINE DOVAN
0440      *****
0441      * THIS IS A MODIFICATION OF THE GENERALIZED STREAMFLOW SIMULATION *
0442      * SYSTEM DEVELOPED BY JOINT FEDERAL-STATE RIVER FORECAST CENTER IN *
0443      * SACRAMENTO, CALIFORNIA (BURNAS, FERRAL, AND MCQUIRE, 1973). THIS *
0444      * PROGRAM UTILIZES EVAPORATION AND PRECIPITATION DATA, AND OTHER *
0445      * WATERSHED PARAMETERS TO GENERATE STREAMFLOW. PARAMETERS DEVELOPED *
0446      * ARE IN ENGLISH UNIT SO THE MODEL USES THE DATA IN ENGLISH UNIT. *
0447      * THE MODEL IS OPERATED FROM THE MONTH OF JUNE TO MAY. *
0448      * THE MODEL HAS BEEN CALIBRATED FROM CONSECUTIVE DATA ON RAINFALL *
0449      * AND RUNOFF FROM TANGEN AND TINAU RIVER AT BUTWAL RESPECTIVELY. *
0450      * THE CALIBRATION YEARS ARE 1977-1982. *
0451      * (NOTE) DAY ONE IS JUNE 1 ST AND LAST DAY OF YEAR IS MAY LAST *
0452      * *
0453      *****
0454      REAL LZTWC, LZFSC, LZFPC, LZTWM, LZFPM, LZFSK, ET(12), LZSK, LZPK,
0455      &RAINZ(11), TOTALP, TOTALQ, TOTALET, TOTALFLO, DSMOWZ(11), EVAP(366)
0456
0457      * REAL FQJUN(40), FQJUL(40), FQAUG(40), FQSEP(40), FQOCT(40),
0458      REAL MSDONQ, YRSUMQ, YRMXFLO, YRMINFLO
0459      INTEGER MON(13)
0460
0461      COMMON /A/ III, ILEAP, PPT(366), FLOWTOT(366), RUNOFF(366), LU
0462      COMMON /B/ FQJUN(40), FQJUL(40), FQAUG(40), FQSEP(40), FQOCT(40),
0463      &FQNOV(40), FQDEC(40), FQJAN(40), FQFEB(40), FQMAR(40), FQAPR(40),
0464      &FQMAY(40), FQMSUN(40), FQYRSUM(40), FQYRMX(40), FQYRMIN(40)
0465      COMMON /C/ MSDONQ, YRSUMQ, YRMXFLO, YRMINFLO
0466
0467      DATA MON/0,30,61,92,122,153,183,214,245,273,304,334,365/
0468
0469      DATA ET/2.7,2.6,2.3,1.9,1.6,1.4,1.0,1.1,1.5,2.8,3.2,3.6/
0470
0471
0472      C*****VARIABLE INITIALIZATION ROUTINE*****
0473
0474
0475      UZTWM = 5.5      ! Inches
0476      UZFWM = 2.0      !
0477      LZTWM = 7.0      !
0478      LZFSM = 5.5
0479      LZFPM = 10.0
0480
0481      *----- RATE OF DRAINAGE TO INTERFLOW AND BASEFLOW
0482      UZK = 0.19
0483      LZSK = 0.02
0484      LZPK = 0.01
0485

```

```

0486 *----- PERCOLATION, FRACT, and PFREE
0487     ZPERC = 25.
0488     REXP = 1.9
0489     FRACT = 1.0
0490     PFREE = 0.3
0491
0492 *----- INITIAL WATER CONTENT AT DIFFERENT LAYERS
0493     UZTWC = 4.9
0494     UZFWC = 0.5
0495     LZTWC = 0.09
0496     LZFSC = 0.05
0497     LZFPC = 1.50
0498
0499     ADIMP = 0.1
0500     PCTIM = 0.0005
0501     SIDE = 5.
0502     ADIMC = 0.0
0503     PBASE = 0.3
0504
0505     SSOUT=0.
0506     SARVA=0.
0507     RSERV=0.4
0508     IMPRT=0.
0509
0510
0511 C *****VARIABLE*INITIALIZATION*DONE,**READ*DATA*FILES*****
0512
0513     TOTALP=0.
0514     TOTALQ=0.
0515     TOTALET=0.
0516     FLOWTOT=0.
0517     TOTALM =0.
0518     TOTALPL=0.
0519     PAV=0.
0520
0521
0522 *----- ANNUAL LOOP BEGINS
0523
0524
0525     DO 999, I = 1, 365+ILEAP
0526
0527     PLIQ = PPT(I)/25.4 ! Assign today's precip in English unit
0528 *
0529 * COMPUTE TRANSPIRATION LOSS FROM UPPER ZONE TENSION WATER.
0530 * THE RAINY DAY NORMAL IS ASSUMED TO BE ONE INCH OF RAIN. THIS IS PDNOR.
0531     PDNOR=1.
0532     PDN20=PDNOR*.20
0533     E2=0.
0534 C *****ALTERED GSSS TO USE EVAP ARRAY WITHOUT MULTIPLYING FRACT****
0535

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0536      CALL CONVERT (1,MAHINA,ILEAP)
0537      EDMND = ET (MAHINA)/25.4 ! Assign evapotranspiration data
0538
0539 C*****
0540
0541      E1 = EDMND * UZTWC/UZTWM
0542      RED = EDMND - E1
0543      UZTWC = UZTWC - E1
0544      IF(UZTWC)100,155,155
0545 100 E1 = EDMND + UZTWC
0546      UZTWC = 0.
0547      RED = EDMND - E1
0548
0549 ***Compute transpiration loss from UPPER ZONE FREE WATER
0550
0551      IF (UZFWC - RED) 130,140,140
0552 130 E2 = UZFWC
0553      UZFWC = 0.
0554      GO TO 170
0555 140 E2 = RED
0556      UZFWC = UZFWC - E2
0557
0558 *** If upper zone free water exceeds upper zone tension
0559 *   water content, then transfer free water into tension.
0560
0561 155 CONTINUE
0562      A = UZTWC/UZTWM
0563      B = UZFWC/UZFWM
0564      IF (A - B) 160,170,170
0565 160 A = (UZTWC + UZFWC)/(UZTWM + UZFWM)
0566      UZTWC = UZTWM * A
0567      UZFWC = UZFWM * A
0568 170 CONTINUE
0569
0570 *** Evaporation from ADIMP area
0571      E5 = E1+RED*(ADIMC-E1-UZTWC)/(UZTWM+LZTWM)
0572 C COMPUTE TRANSPIRATION LOSS FROM LOWER ZONE TENSION WATER.
0573      E3=(RED-E2)*LZTWC/(UZTWM+LZTWM)
0574      LZTWC=LZTWC-E3
0575      IF(LZTWC)180,185,185
0576 180 E3=E3+LZTWC
0577      LZTWC=0.
0578
0579 C RESUPPLY LOWER ZONE TENSION WATER FROM LOWER ZONE FREE IF MORE
0580 C WATER AVAILABLE THERE.
0581
0582 185 CONTINUE
0583      A = LZTWC/LZTWM
0584      B=(LZFPC+LZFSC-SAVED+LZTWC)/(LZFPM+LZFSM-SAVED+LZTWM)
0585      IF(A-B)190,210,210
0586 190 DEL = (B-A) * LZTWM

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0587
0588 C TRANSFER WATER FROM LOWER ZONE SECONDARY FREE WATER TO LOWER
0589 C ZONE TENSION
0590     LZTWC=LZTWC+DEL
0591     LZFSC=LZFSC-DEL
0592     IF(LZFSC)200,210,210
0593 C TRANSFER PRIMARY FREE WATER IF SECONDARY FREE WATER INADEQUATE.
0594
0595     200 LZFPC=LZFPC+LZFSC
0596     LZFSC=0.
0597     210 CONTINUE
0598 C RUNOFF FROM IMPERVIOUS OR WATER-COVERED AREA
0599     ROIMP=PLIQ*PCTIM
0600
0601 C ADJUST ADIMC STORAGE FOR EVAPORATION
0602     ADIMC=ADIMC-ES
0603     IF(ADIMC)211,212,212.
0604     211 ES=ES+ADIMC
0605     ADIMC=0
0606     212 ES=ES*ADIMP.
0607
0608 C REDUCE RAIN BY AMOUNT OF UPPER ZONE TENSION WATER DEFICIENCY.
0609     PAV=PLIQ+UZTWC-UZTWM
0610     IF(PAV)220,230,230
0611
0612 C FILL UPPER ZONE TENSION WATER AS MUCH AS RAIN PERMITS
0613
0614     220 UZTWC=UZTWC+PLIQ
0615     PAV=0.
0616     GO TO 240
0617     230 UZTWC=UZTWM
0618
0619 C DETERMINE NUMBER OF INCREMENTS.
0620
0621     240 NINC=1.+5.*(UZFWC*FRACT+PAV)
0622     ADIMC=ADIMC+PLIQ-PAV
0623     DINC=NINC
0624     DINC=1./DINC
0625     PINC=PAV*DINC
0626     FLOBF=0.
0627     FLOSF=0.
0628     FLOIN=0.
0629     DINC=DINC*FRACT
0630 C MODIFICATIONS THIS DATE TO LUMP RAINFALL IN LESS THAN THE 7-14-71
0631 C TOTAL PERIOD TO APPROXIMATE SOMEWHAT THE INTENSITY VARIATION EFFECT.
0632

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0633      DUZ=UZK
0634      DLZP=LZPK
0635      DLZS=LZSK
0636      ITIME=2
0637      ADJ=1.
0638      IF(PAV-PDN20)246,246,242
0639      242 IF(PAV-PDNOR)243,244,244
0640
0641      C THE EFFECTIVE LENGTH OF RAIN IN A PERIOD IS ASSUMED TO BE HALF
0642      C PERIOD LENGTH FOR RAIN EQUAL TO THE RAINY PERIOD NORMAL.
0643
0644      243 ADJ=.5*SQRT(PAV/PDNOR)
0645      GO TO 245
0646      244 ADJ=1.-.5*PDNOR/PAV
0647      245 CONTINUE
0648      ITIME=0
0649      C SECTIONAL ANALYSIS*****
0650
0651      400 ITIME=ITIME+1
0652      NINC=1.+S.*(UZFWC*FRACT*ADJ+PAV)
0653      DINC=NINC
0654      DINC=1./DINC
0655      PINC=PAV*DINC
0656      DINC=DINC*FRACT*ADJ
0657      C ADJUSTMENT OF FREE-WATER STORAGE DEPLETION COEFFICIENTS FOR
0658      C PERIODS DIFFERENT FROM ONE DAY WHEN COEFFICIENTS WERE DERIVED FOR
0659      C DAILY DISCHARGE. DC(X DAYS)=1.-(1.-DC(ONE DAY))**X
0660
0661      GO TO 247
0662      246 IF(DINC-1.)247,248,248
0663      247 CONTINUE
0664      DUZ=1.-(1.-UZK)**DINC
0665      DLZP=1.-(1.-LZPK)**DINC
0666      DLZS=1.-(1.-LZSK)**DINC
0667      248 CONTINUE
0668
0669      C BEGIN DRAINAGE AND PERCOLATION LOOP
0670
0671      DO 385 INC=1,NINC
0672      PAV=PINC
0673      C IMPERVIOUS RUNOFF MODIFIED FOR CHANGE IN IMPERVIOUS AND WATER-COVERED
0674      C AREA PERCENT.
0675
0676      RATIO=(ADINC-UZTWC)/LZTWM
0677      ADDRO=PINC*RATIO*RATIO
0678
0679      C COMPUTE BASEFLOW FROM LOWER ZONE.
0680

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0681      BF=LZFPC*DLZP
0682      FLOBF=FLOBF+BF
0683      LZFPF=LZFPF-BF
0684      BF=LZFSC*DLZS
0685      LZFSC=LZFSC-BF
0686      FLOBF=FLOBF+BF
0687
0688      C COMPUTE PERCOLATION, BUT NO CALCULATIONS IF NO WATER TO PERCOLATE.
0689          IF(PINC+UZFWC-.01)380,380.250
0690
0691      C COMPUTE PERCOLATION FROM UPPER ZONE TO LOWER. PERC=PBASE IN LITERATURE.
0692          250 PERCM=LZPK*LZFPF+LZSK*LZFSM
0693          PERC=PERCM*DINC
0694          PERC=PERC*UZFWC/UZFWM
0695      C MODIFICATION FOR FASTER PERCOLATION INTO MOISTURE-DEFICIENT SOIL
0696      C IN LOWER LAYER.
0697          PERC=PERC*(1.+(2PERC*(1.-(LZFPF+LZFSC+LZTWC)/(LZFPF+LZFSM+LZTWM))
0698          ***REXP))
0699          IF(PERC-UZFWC)270,270,260
0700          260 PERC=UZFWC
0701          UZFWC=0
0702          GO TO 300
0703          270 UZFWC=UZFWC-PERC
0704          CHECK=LZFPF+LZFSC+LZTWC+PERC-LZFPF-LZFSM-LZTWM
0705          IF(CHECK)290,290.280
0706          280 PERC=PERC-CHECK
0707          UZFWC=UZFWC+CHECK
0708      C COMPUTE INTERFLOW
0709          290 DEL=DUZ*UZFWC
0710          FLOIN=FLOIN+DEL
0711          UZFWC=UZFWC-DEL
0712
0713      C DISTRIBUTE PERCOLATED WATER,CHECK IF COMPUTED PERCOLATION EXCEEDS
0714      C AVAILABLE TENSION CAPACITY IN LOWER LEVEL. REPLACE TENSION WATER FIRST.
0715          300 CONTINUE
0716      C SAVE PERCOLATED VOLUME FOR PFREE AREA
0717          SPERC=PERC
0718          PERC=PERC*(1.-PFREE)
0719          IF(PERC-LZTWM+LZTWC)310,310,320
0720          310 LZTWC=LZTWC+PERC
0721          PERC=0.
0722          GO TO 321
0723          320 PERC=PERC-LZTWM+LZTWC
0724          LZTWC=LZTWM
0725
0726      C DISTRIBUTE PERIOD PERCOLATION IN EXCESS OF LOWER ZONE TENSION REQUIREMENT.
0727          321 PERC=PERC+SPERC*PFREE
0728          IF(PERC)340,340,323
0729          323 CONTINUE

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```

0730      HPL=LZFP/ (LZFP+LZFS)
0731      RATL=LZFP/LZFP
0732      PERCS=PERC-PERC*(HPL*2.*(1.-RATL)/((1.-RATL)+(1.-LZFS/LZFS)))
0733      LZFS=LZFS+PERCS
0734      IF(LZFS-LZFS)330,330,322
0735      322 PERCS=PERCS-LZFS+LZFS
0736      LZFS=LZFS
0737      330 CONTINUE
0738      LZFP=LZFP+(PERC-PERCS)
0739      340 IF(PAV)380,380,350
0740      350 CONTINUE
0741
0742      C DISTRIBUTE PERIOD RAIN IN EXCESS OF UPPER ZONE TENSION REQUIREMENTS
0743      C TO UPPER ZONE FREEWATER AND SURFACE RUNOFF.
0744      C CHECK WHETHER RESIDUAL RAIN EXCEEDS AVAILABLE UPPER LEVEL FREE WATER MAX.
0745
0746      IF(PAV-UZFW+UZFW)360,360,370
0747      360 UZFW=UZFW+PAV
0748      GO TO 380
0749      370 PAV=PAV-UZFW+UZFW
0750      UZFW=UZFW
0751      FLOS=FLOS+PAV
0752      ADDR=ADDR+PAV*(1.-ADDR/PINC)
0753      380 CONTINUE
0754      ADIM=ADIM+PINC-ADDR
0755      ROIMP=ROIMP+ADDR*ADIM
0756      C RECYCLE POINT FOR THE TIME INCREMENT ANALYSIS.
0757      385 CONTINUE
0758      ADJ=1.-ADJ
0759      PAV=0.
0760      390 CONTINUE
0761
0762      C SECTIONAL RECYCLE POINT (1ST PASS=WET,2ND PASS=DRY)
0763      IF(ITIME-1)400,400,410
0764      410 EUSED=E1+E2+E3
0765      FLOS=FLOS*(1.-PCTIM-ADIM)
0766      FLOIN=FLOIN*(1.-PCTIM-ADIM)
0767      FLOBF=FLOBF*(1.-PCTIM-ADIM)
0768      C OUTPUT COMPUTED,WRITTEN TO FILE AND SCREEN,THEN RE-RUN OPTION EXECUTED.
0769      FLOWTOT(I)=FLOS+FLOIN+FLOBF+ROIMP
0770      SOILWATER=UZTW+UZFW+LZTW+LZFS+LZFP
0771
0772      C*****ADJUSTMENT FOR RIDICULOUSLY PROLONGED LOW FLOWS*****
0773      IF(FLOWTOT.LE.0.00014)FLOWTOT=0.
0774
0775      FLOWTOT(I) = FLOWTOT(I)*25.4 ! CONVERT TO METRIC MM
0776      RUNOFF(I) = FLOWTOT(I)/.152 ! CONVERT DAILY MEAN FLOW C.M/S
0777
0778      ***** STATISTICAL CALULATION OF FLOWS
0779

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0780
0781     IF( MAHINA .EQ. 1) THEN
0782         CALL UPDATE(FLOWTOT(I), FQJUN, 0.,17.,40) ! JUNE
0783     ELSE IF (MAHINA .EQ. 2) THEN
0784         CALL UPDATE(FLOWTOT(I), FQJUL, 0.,13.,40) ! JULY
0785     ELSE IF (MAHINA .EQ. 3) THEN
0786         CALL UPDATE(FLOWTOT(I), FQAUG, 0.,17.,40) ! AUGUST
0787     ELSE IF (MAHINA .EQ. 4) THEN
0788         CALL UPDATE(FLOWTOT(I), FQSEP, 0.,12.,40) ! SEPTEMBER
0789     ELSE IF (MAHINA .EQ. 5) THEN
0790         CALL UPDATE(FLOWTOT(I), FQOCT, 0.,10.,40) ! OCTOBER
0791     ELSE IF (MAHINA .EQ. 6) THEN
0792         CALL UPDATE(FLOWTOT(I), FQNOV, 0.,25,40) ! NOVEMBER
0793     ELSE IF (MAHINA .EQ. 7) THEN
0794         CALL UPDATE(FLOWTOT(I), FQDEC, 0.,0.2,40) ! DECEMBER
0795     ELSE IF (MAHINA .EQ. 8) THEN
0796         CALL UPDATE(FLOWTOT(I), FQJAN, 0.,0.15,40) ! JANUARY
0797     ELSE IF (MAHINA .EQ. 9) THEN
0798         CALL UPDATE(FLOWTOT(I), FQFEB, 0.,.20,40) ! FEBRUARY
0799     ELSE IF (MAHINA .EQ. 10) THEN
0800         CALL UPDATE(FLOWTOT(I), FQMAR, 0.,0.2,40) ! MARCH
0801     ELSE IF (MAHINA .EQ. 11) THEN
0802         CALL UPDATE(FLOWTOT(I), FQAPR, 0.,.10,40) ! APRIL
0803     ELSE IF (MAHINA .EQ. 12) THEN
0804         CALL UPDATE(FLOWTOT(I), FQMAY, 0.,0.3,40) ! MAY
0805     ELSE
0806     END IF
0807 ****
0808 * ----- COUNT THE MONSOON SEASONAL AND ANNUAL TOTAL RUNOFF
0809     IF(MAHINA .EQ. 1 .OR. MAHINA .EQ. 2 .OR. MAHINA .EQ. 3 .OR.
0810     &MAHINA .EQ. 4) THEN
0811         HSDONG = MSDONG + FLOWTOT(I)
0812     ELSE
0813     END IF
0814 ***
0815     YRSUMQ = YRSUMQ + FLOWTOT(I) ! ANNUAL TOTAL FLOW
0816
0817 *** Sorting the peak discharge and minimum flow of a year.
0818
0819     IF(RUNOFF(I) .GT. YRMXFLO) YRMXFLO = RUNOFF(I)
0820     IF(RUNOFF(I) .LT. YRMINFLO) YRMINFLO = RUNOFF(I)
0821
0822 ** A count for the consecutive days of low baseflow.
0823 ** A baseflow less than or equal to 1.5 Cu. M /S is considered.
0824 *     IF (RUNOFF(I) .LE. 1.5) THEN
0825 *         WRITE(345,*) III, I, RUNOFF(I)
0826 *     ELSE
0827 *     END IF
0828     999 CONTINUE

```

```

00829 ***** END OF DAILY LOOP
00830 * COUNT YEARLY FREQUENCIES.
00831     CALL UPDATE (MSOONO, FOMSUN, 50.0,70.,40)
00832     CALL UPDATE (YRSUMQ, FOYRSUM.100.,65.,40)
00833     CALL UPDATE (YRMXFLO,FOYRMX,0.0 ,120.,40)
00834     CALL UPDATE (YRMINFLO, FOYRMIN, 0.0,0.07,40)
00835 ***** END OF ANNUAL LOOP
00836
00837 C     WRITE(234,*) III, YRSUMQ
00838
00839     RETURN
00840 1000 STOP
00841     END
00842
00843
00844     SUBROUTINE CONVERT (I, MAHINA, ILEAP)
00845 *****
00846 * A SUBPROGRAM TO CONVERT THE DAILY DATA TO MONTHLY SET
00847     INTEGER I, J, MAHINA, NDAY, LAST(12)
00848
00849
00850     LAST(1) = 30
00851     LAST(2) = 61
00852     LAST(3) = 92
00853     LAST(4) =122
00854     LAST(5) =153
00855     LAST(6) =183
00856     LAST(7) =214
00857     LAST(8) =245
00858     LAST(9) =273+ILEAP
00859     LAST(10)=304+ILEAP
00860     LAST(11)=334+ILEAP
00861     LAST(12)=365+ILEAP
00862
00863
00864     DO 600 J = 1, 12
00865
00866         IF( I .LE. LAST(J)) THEN
00867             MAHINA = J
00868             RETURN
00869         ELSE
00870             END IF
00871
00872 600 CONTINUE
00873
00874     RETURN
00875     END
00876
00877

```

```
0878
0879     SUBROUTINE UPDATE( FLOW, FREQ, CLS, DELTA, LG)
0880     *****
0881 *     A SUBPROGRAM TO COUNT THE FREQUENCY OF THE FLOW
0882
0883     REAL FREQ(40)
0884
0885     DO 400 J = 1, LG
0886         RLOW = FLOAT(J-1)*DELTA + CLS
0887         RUPR = FLOAT(J)*DELTA + CLS
0888         IF (FLOW .LE. RLOW .OR. FLOW .GT. RUPR) GO TO 400
0889         FREQ(J) = FREQ(J) + 1
0891
0891     RETURN
0892 400 CONTINUE
0893
0894     RETURN
0895     END
0896
0897
0898
```

```

0899      SUBROUTINE REPORT (FREQ, CLS, CLASS, LU, NN)
0900      *****
0901      * A SUBPROGRAM TO REPORT THE MONTHLY, SEASONAL AND YEARLY FREQUENCIES
0902
0903      REAL    CLS, PDF, CDF, CLASS
0904      REAL    FREQ(40), TOTAL
0905
0906      WRITE(LU,10)
0907      10  FORMAT(3X,'      CLASS          FREQUENCY      PDF      CDF',/)
0908
0909      TOTAL = 0
0910      X = 0.
0911      XX = 0.
0912      DO 100 J = 1, NN
0913          TOTAL = TOTAL + FREQ(J)
0914      100  CONTINUE
0915      IF (TOTAL.LE.0) THEN TOTAL = 1
0916      CDF = 0.
0917      DO 200 I = 1, NN
0918          FREQI = FREQ(I)
0919          CDF = CDF + FREQI
0920          PDF = FREQI/TOTAL
0921          CDF1 = CDF/TOTAL
0922          XL = CLASS*FLOAT(I-1) + CLS
0923          XU = CLASS*FLOAT(I) + CLS
0924          X = FREQI * 0.5*(XL+XU) + X
0925          XX = XX + FREQI*.25*(XL+XU)*(XL+XU)
0926          WRITE(LU,20)XL,XU,FREQI,PDF,CDF1
0927
0928      200  CONTINUE
0929      20  FORMAT(3X,F7.2,'-',F7.2,4X,F8.1,4X,F9.7,2X,F9.7)
0930
0931          X = X/TOTAL
0932          XX = SQRT(XX/TOTAL - X*X)
0933          WRITE(LU,30) TOTAL, X, XX
0934      30  FORMAT(/,10X,'      TOTAL (EVENTS) = ',F10.1,/
0935      &10X,'      MEAN (FLOW) = ', F10.1,/,
0936      &10X,'      STD. DEV (FLOW) =',F10.2,/)
0937      RETURN
0938      END
0939

```

APPENDIX B

LISTING OF PARAMETER VALUES FOR
THE STREAM FLOW GENERATION MODEL

Parameter	Description	Value
UZTWM	Upper zone tension water capacity	147.30
UZFWM	Upper zone free water capacity	53.30
LZTWM	Lower zone tension water capacity	180.30
LZFSM	Lower zone free water supplemental capacity	139.70
LZFPM	Lower zone free water primary capacity	254.00
UZK	Fraction of upper zone free water intering interflow in one day	0.19
LZSK	Fraction of lower zone supplemental free water drained in one day	0.02
LZPK	Fraction of lower zone primary free water drained in one day	0.01
ZPERC	Percolation factor	25.00
REXP	An exponent governing rate of change of percolation rate with changing lower zone water contents	1.90
SIDE	A ratio of non-channel baseflow to channel baseflow	5.00
SSOUT	Surface channel flow required before channel flow becomes visible at the gaging station	0.0

Parameter	Descriptoin	Value
PCTIM	Watershed fraction contributing to the instantaneous runoff	0.0005
SARVA	Watershed fraction covered by streams, lakes, and riparian vegetations	0.0
RSERV	Fraction of lower zone free water not available for lower zone tension water	0.4
UZTWC	Upper zone tension initial water content	124.5
UZFWC	Upper zone initial free water content	12.70
LZTWC	Lower zone initial tension water content	2.30
LZFSC	Lower zone initial free supplemental water content	1.3
LZFPC	Lower zone initial free primary water content	38.1
ADIMP	Additional fraction of impervious area which develops as tension water requirements are met	0.10
PFREE	Fraction of percolated water directly claimed by lower zone free water storages	0.30

APPENDIX C

SAMPLE OUTPUT OF TINAU MODEL

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - JANUARY
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .30	2.0	.0006452	.0006452
.30- .60	43.0	.0138710	.0145161
.60- .90	99.0	.0319355	.0464516
.90- 1.20	346.0	.1116129	.1580645
1.20- 1.50	592.0	.1909677	.3490323
1.50- 1.80	719.0	.2319355	.5809678
1.80- 2.10	614.0	.1980645	.7790322
2.10- 2.40	387.0	.1248387	.9038709
2.40- 2.70	181.0	.0583871	.9622581
2.70- 3.00	63.0	.0203226	.9825807
3.00- 3.30	24.0	.0077419	.9903226
3.30- 3.60	2.0	.0006452	.9909678
3.60- 3.90	3.0	.0009677	.9919355
3.90- 4.20	2.0	.0006452	.9925807
4.20- 4.50	1.0	.0003226	.9929032
4.50- 4.80	0.0	0.0000000	.9929032
4.80- 5.10	4.0	.0012903	.9941936
5.10- 5.40	2.0	.0006452	.9948387
5.40- 5.70	4.0	.0012903	.9961290
5.70- 6.00	2.0	.0006452	.9967742
6.00- 6.30	1.0	.0003226	.9970968
6.30- 6.60	3.0	.0009677	.9980645
6.60- 6.90	1.0	.0003226	.9983871
6.90- 7.20	1.0	.0003226	.9987097
7.20- 7.50	1.0	.0003226	.9990323
7.50- 7.80	1.0	.0003226	.9993548
7.80- 8.10	0.0	0.0000000	.9993548
8.10- 8.40	1.0	.0003226	.9996774
8.40- 8.70	0.0	0.0000000	.9996774
8.70- 9.00	0.0	0.0000000	.9996774
9.00- 9.30	0.0	0.0000000	.9996774
9.30- 9.60	0.0	0.0000000	.9996774
9.60- 9.90	0.0	0.0000000	.9996774
9.90- 10.20	1.0	.0003226	1.0000000
10.20- 10.50	0.0	0.0000000	1.0000000
10.50- 10.80	0.0	0.0000000	1.0000000
10.80- 11.10	0.0	0.0000000	1.0000000
11.10- 11.40	0.0	0.0000000	1.0000000
11.40- 11.70	0.0	0.0000000	1.0000000
11.70- 12.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
 MEAN (FLOW) = 1.7
 STD. DEV (FLOW) = .65

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - FEBRUARY
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .25	35.0	.0123894	.0123894
.25- .50	240.0	.0849558	.0973451
.50- .75	1060.0	.3752213	.4725664
.75- 1.00	926.0	.3277876	.8003540
1.00- 1.25	392.0	.1387611	.9391150
1.25- 1.50	127.0	.0449558	.9840708
1.50- 1.75	15.0	.0053097	.9893805
1.75- 2.00	4.0	.0014159	.9907964
2.00- 2.25	2.0	.0007080	.9915044
2.25- 2.50	4.0	.0014159	.9929204
2.50- 2.75	2.0	.0007080	.9936283
2.75- 3.00	3.0	.0010619	.9946903
3.00- 3.25	2.0	.0007080	.9953983
3.25- 3.50	1.0	.0003540	.9957522
3.50- 3.75	0.0	0.0000000	.9957522
3.75- 4.00	2.0	.0007080	.9964602
4.00- 4.25	4.0	.0014159	.9978760
4.25- 4.50	1.0	.0003540	.9982301
4.50- 4.75	1.0	.0003540	.9985840
4.75- 5.00	0.0	0.0000000	.9985840
5.00- 5.25	1.0	.0003540	.9989381
5.25- 5.50	0.0	0.0000000	.9989381
5.50- 5.75	2.0	.0007080	.9996461
5.75- 6.00	0.0	0.0000000	.9996461
6.00- 6.25	0.0	0.0000000	.9996461
6.25- 6.50	0.0	0.0000000	.9996461
6.50- 6.75	0.0	0.0000000	.9996461
6.75- 7.00	0.0	0.0000000	.9996461
7.00- 7.25	0.0	0.0000000	.9996461
7.25- 7.50	0.0	0.0000000	.9996461
7.50- 7.75	0.0	0.0000000	.9996461
7.75- 8.00	0.0	0.0000000	.9996461
8.00- 8.25	0.0	0.0000000	.9996461
8.25- 8.50	0.0	0.0000000	.9996461
8.50- 8.75	0.0	0.0000000	.9996461
8.75- 9.00	1.0	.0003540	1.0000000
9.00- 9.25	0.0	0.0000000	1.0000000
9.25- 9.50	0.0	0.0000000	1.0000000
9.50- 9.75	0.0	0.0000000	1.0000000
9.75- 10.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 2825.0
 MEAN (FLOW) = .8
 STD. DEV (FLOW) = .41

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - MARCH
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .30	77.0	.0248387	.0248387
.30- .60	1035.0	.3338709	.3587097
.60- .90	1421.0	.4583871	.8170968
.90- 1.20	446.0	.1438710	.9609678
1.20- 1.50	101.0	.0325806	.9935484
1.50- 1.80	11.0	.0035484	.9970968
1.80- 2.10	2.0	.0006452	.9977419
2.10- 2.40	1.0	.0003226	.9980645
2.40- 2.70	1.0	.0003226	.9983871
2.70- 3.00	0.0	0.0000000	.9983871
3.00- 3.30	0.0	0.0000000	.9983871
3.30- 3.60	1.0	.0003226	.9987097
3.60- 3.90	0.0	0.0000000	.9987097
3.90- 4.20	0.0	0.0000000	.9987097
4.20- 4.50	0.0	0.0000000	.9987097
4.50- 4.80	2.0	.0006452	.9993548
4.80- 5.10	1.0	.0003226	.9996774
5.10- 5.40	0.0	0.0000000	.9996774
5.40- 5.70	0.0	0.0000000	.9996774
5.70- 6.00	0.0	0.0000000	.9996774
6.00- 6.30	0.0	0.0000000	.9996774
6.30- 6.60	0.0	0.0000000	.9996774
6.60- 6.90	0.0	0.0000000	.9996774
6.90- 7.20	0.0	0.0000000	.9996774
7.20- 7.50	0.0	0.0000000	.9996774
7.50- 7.80	0.0	0.0000000	.9996774
7.80- 8.10	0.0	0.0000000	.9996774
8.10- 8.40	0.0	0.0000000	.9996774
8.40- 8.70	0.0	0.0000000	.9996774
8.70- 9.00	0.0	0.0000000	.9996774
9.00- 9.30	0.0	0.0000000	.9996774
9.30- 9.60	0.0	0.0000000	.9996774
9.60- 9.90	1.0	.0003226	1.0000000
9.90- 10.20	0.0	0.0000000	1.0000000
10.20- 10.50	0.0	0.0000000	1.0000000
10.50- 10.80	0.0	0.0000000	1.0000000
10.80- 11.10	0.0	0.0000000	1.0000000
11.10- 11.40	0.0	0.0000000	1.0000000
11.40- 11.70	0.0	0.0000000	1.0000000
11.70- 12.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
 MEAN (FLOW) = .7
 STD. DEV (FLOW) = .33

MONTHLY SIMULATION RUNOFF FREQUENCY
SUMMARY RESULT FOR THE MONTH - APRIL
RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .35	33.0	.0110000	.0110000
.35- .70	210.0	.0700000	.0810000
.70- 1.05	957.0	.3190000	.4000000
1.05- 1.40	1035.0	.3450000	.7450000
1.40- 1.75	509.0	.1696667	.9146667
1.75- 2.10	178.0	.0593333	.9740000
2.10- 2.45	52.0	.0173333	.9913334
2.45- 2.80	16.0	.0053333	.9966667
2.80- 3.15	0.0	0.0000000	.9966667
3.15- 3.50	0.0	0.0000000	.9966667
3.50- 3.85	0.0	0.0000000	.9966667
3.85- 4.20	1.0	.0003333	.9970000
4.20- 4.55	0.0	0.0000000	.9970000
4.55- 4.90	0.0	0.0000000	.9970000
4.90- 5.25	1.0	.0003333	.9973333
5.25- 5.60	1.0	.0003333	.9976667
5.60- 5.95	1.0	.0003333	.9980000
5.95- 6.30	3.0	.0010000	.9990000
6.30- 6.65	0.0	0.0000000	.9990000
6.65- 7.00	0.0	0.0000000	.9990000
7.00- 7.35	1.0	.0003333	.9993334
7.35- 7.70	0.0	0.0000000	.9993334
7.70- 8.05	1.0	.0003333	.9996667
8.05- 8.40	0.0	0.0000000	.9996667
8.40- 8.75	0.0	0.0000000	.9996667
8.75- 9.10	0.0	0.0000000	.9996667
9.10- 9.45	0.0	0.0000000	.9996667
9.45- 9.80	0.0	0.0000000	.9996667
9.80- 10.15	1.0	.0003333	1.0000000
10.15- 10.50	0.0	0.0000000	1.0000000
10.50- 10.85	0.0	0.0000000	1.0000000
10.85- 11.20	0.0	0.0000000	1.0000000
11.20- 11.55	0.0	0.0000000	1.0000000
11.55- 11.90	0.0	0.0000000	1.0000000
11.90- 12.25	0.0	0.0000000	1.0000000
12.25- 12.60	0.0	0.0000000	1.0000000
12.60- 12.95	0.0	0.0000000	1.0000000
12.95- 13.30	0.0	0.0000000	1.0000000
13.30- 13.65	0.0	0.0000000	1.0000000
13.65- 14.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3000.0
MEAN (FLOW) = 1.2
STD. DEV (FLOW) = .51

MONTHLY SIMULATION RUNOFF FREQUENCY
SUMMARY RESULT FOR THE MONTH - MAY
RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 1.50	2087.0	.6732258	.6732258
1.50- 3.00	888.0	.2864516	.9596775
3.00- 4.50	37.0	.0119355	.9716129
4.50- 6.00	24.0	.0077419	.9793549
6.00- 7.50	5.0	.0016129	.9809678
7.50- 9.00	11.0	.0035484	.9845161
9.00- 10.50	9.0	.0029032	.9874194
10.50- 12.00	9.0	.0029032	.9903226
12.00- 13.50	3.0	.0009677	.9912903
13.50- 15.00	4.0	.0012903	.9925807
15.00- 16.50	3.0	.0009677	.9935484
16.50- 18.00	1.0	.0003226	.9938710
18.00- 19.50	5.0	.0016129	.9954839
19.50- 21.00	0.0	0.0000000	.9954839
21.00- 22.50	3.0	.0009677	.9964516
22.50- 24.00	0.0	0.0000000	.9964516
24.00- 25.50	1.0	.0003226	.9967742
25.50- 27.00	0.0	0.0000000	.9967742
27.00- 28.50	0.0	0.0000000	.9967742
28.50- 30.00	1.0	.0003226	.9970968
30.00- 31.50	2.0	.0006452	.9977419
31.50- 33.00	3.0	.0009677	.9987097
33.00- 34.50	1.0	.0003226	.9990323
34.50- 36.00	0.0	0.0000000	.9990323
36.00- 37.50	0.0	0.0000000	.9990323
37.50- 39.00	0.0	0.0000000	.9990323
39.00- 40.50	0.0	0.0000000	.9990323
40.50- 42.00	0.0	0.0000000	.9990323
42.00- 43.50	0.0	0.0000000	.9990323
43.50- 45.00	1.0	.0003226	.9993548
45.00- 46.50	0.0	0.0000000	.9993548
46.50- 48.00	0.0	0.0000000	.9993548
48.00- 49.50	0.0	0.0000000	.9993548
49.50- 51.00	0.0	0.0000000	.9993548
51.00- 52.50	0.0	0.0000000	.9993548
52.50- 54.00	0.0	0.0000000	.9993548
54.00- 55.50	0.0	0.0000000	.9993548
55.50- 57.00	1.0	.0003226	.9996774
57.00- 58.50	1.0	.0003226	1.0000000
58.50- 60.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
MEAN (FLOW) = 1.6
STD. DEV (FLOW) = 2.77

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - JUNE
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 14.00	2928.0	.9760000	.9760000
14.00- 28.00	27.0	.0090000	.9850000
28.00- 42.00	16.0	.0053333	.9903333
42.00- 56.00	4.0	.0013333	.9916667
56.00- 70.00	6.0	.0020000	.9936666
70.00- 84.00	5.0	.0016667	.9953333
84.00- 98.00	5.0	.0016667	.9970000
98.00- 112.00	4.0	.0013333	.9983333
112.00- 126.00	0.0	0.0000000	.9983333
126.00- 140.00	0.0	0.0000000	.9983333
140.00- 154.00	0.0	0.0000000	.9983333
154.00- 168.00	2.0	.0006667	.9990000
168.00- 182.00	1.0	.0003333	.9993334
182.00- 196.00	0.0	0.0000000	.9993334
196.00- 210.00	0.0	0.0000000	.9993334
210.00- 224.00	0.0	0.0000000	.9993334
224.00- 238.00	0.0	0.0000000	.9993334
238.00- 252.00	1.0	.0003333	.9996667
252.00- 266.00	0.0	0.0000000	.9996667
266.00- 280.00	0.0	0.0000000	.9996667
280.00- 294.00	0.0	0.0000000	.9996667
294.00- 308.00	0.0	0.0000000	.9996667
308.00- 322.00	0.0	0.0000000	.9996667
322.00- 336.00	0.0	0.0000000	.9996667
336.00- 350.00	0.0	0.0000000	.9996667
350.00- 364.00	0.0	0.0000000	.9996667
364.00- 378.00	0.0	0.0000000	.9996667
378.00- 392.00	0.0	0.0000000	.9996667
392.00- 406.00	0.0	0.0000000	.9996667
406.00- 420.00	0.0	0.0000000	.9996667
420.00- 434.00	0.0	0.0000000	.9996667
434.00- 448.00	0.0	0.0000000	.9996667
448.00- 462.00	0.0	0.0000000	.9996667
462.00- 476.00	0.0	0.0000000	.9996667
476.00- 490.00	0.0	0.0000000	.9996667
490.00- 504.00	0.0	0.0000000	.9996667
504.00- 518.00	1.0	.0003333	1.0000000
518.00- 532.00	0.0	0.0000000	1.0000000
532.00- 546.00	0.0	0.0000000	1.0000000
546.00- 560.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3000.0
 MEAN (FLOW) = 8.2
 STD. DEV (FLOW) = 13.21

MONTHLY SIMULATION RUNOFF FREQUENCY
SUMMARY RESULT FOR THE MONTH - JULY
RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 13.00	2928.0	.9445162	.9445162
13.00- 26.00	40.0	.0129032	.9574194
26.00- 39.00	28.0	.0090323	.9664516
39.00- 52.00	30.0	.0096774	.9761291
52.00- 65.00	14.0	.0045161	.9806452
65.00- 78.00	9.0	.0029032	.9835484
78.00- 91.00	9.0	.0029032	.9864516
91.00- 104.00	9.0	.0029032	.9893548
104.00- 117.00	5.0	.0016129	.9909678
117.00- 130.00	3.0	.0009677	.9919355
130.00- 143.00	3.0	.0009677	.9929032
143.00- 156.00	3.0	.0009677	.9938710
156.00- 169.00	4.0	.0012903	.9951613
169.00- 182.00	2.0	.0006452	.9958065
182.00- 195.00	3.0	.0009677	.9967742
195.00- 208.00	0.0	0.0000000	.9967742
208.00- 221.00	0.0	0.0000000	.9967742
221.00- 234.00	1.0	.0003226	.9970968
234.00- 247.00	3.0	.0009677	.9980645
247.00- 260.00	0.0	0.0000000	.9980645
260.00- 273.00	0.0	0.0000000	.9980645
273.00- 286.00	0.0	0.0000000	.9980645
286.00- 299.00	1.0	.0003226	.9983871
299.00- 312.00	0.0	0.0000000	.9983871
312.00- 325.00	2.0	.0006452	.9990323
325.00- 338.00	0.0	0.0000000	.9990323
338.00- 351.00	0.0	0.0000000	.9990323
351.00- 364.00	2.0	.0006452	.9996774
364.00- 377.00	0.0	0.0000000	.9996774
377.00- 390.00	0.0	0.0000000	.9996774
390.00- 403.00	0.0	0.0000000	.9996774
403.00- 416.00	0.0	0.0000000	.9996774
416.00- 429.00	1.0	.0003226	1.0000000
429.00- 442.00	0.0	0.0000000	1.0000000
442.00- 455.00	0.0	0.0000000	1.0000000
455.00- 468.00	0.0	0.0000000	1.0000000
468.00- 481.00	0.0	0.0000000	1.0000000
481.00- 494.00	0.0	0.0000000	1.0000000
494.00- 507.00	0.0	0.0000000	1.0000000
507.00- 520.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
MEAN (FLOW) = 10.2
STD. DEV (FLOW) = 22.75

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - AUGUST
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 17.00	2907.0	.9377420	.9377420
17.00- 34.00	67.0	.0216129	.9593549
34.00- 51.00	44.0	.0141935	.9735484
51.00- 68.00	29.0	.0093548	.9829032
68.00- 85.00	15.0	.0048387	.9877419
85.00- 102.00	10.0	.0032258	.9909678
102.00- 119.00	5.0	.0016129	.9925807
119.00- 136.00	4.0	.0012903	.9938710
136.00- 153.00	4.0	.0012903	.9951613
153.00- 170.00	2.0	.0006452	.9958065
170.00- 187.00	1.0	.0003226	.9961290
187.00- 204.00	1.0	.0003226	.9964516
204.00- 221.00	4.0	.0012903	.9977419
221.00- 238.00	1.0	.0003226	.9980645
238.00- 255.00	1.0	.0003226	.9983871
255.00- 272.00	1.0	.0003226	.9987097
272.00- 289.00	0.0	0.0000000	.9987097
289.00- 306.00	1.0	.0003226	.9990323
306.00- 323.00	0.0	0.0000000	.9990323
323.00- 340.00	0.0	0.0000000	.9990323
340.00- 357.00	1.0	.0003226	.9993548
357.00- 374.00	1.0	.0003226	.9996774
374.00- 391.00	0.0	0.0000000	.9996774
391.00- 408.00	0.0	0.0000000	.9996774
408.00- 425.00	0.0	0.0000000	.9996774
425.00- 442.00	0.0	0.0000000	.9996774
442.00- 459.00	0.0	0.0000000	.9996774
459.00- 476.00	0.0	0.0000000	.9996774
476.00- 493.00	0.0	0.0000000	.9996774
493.00- 510.00	0.0	0.0000000	.9996774
510.00- 527.00	0.0	0.0000000	.9996774
527.00- 544.00	0.0	0.0000000	.9996774
544.00- 561.00	0.0	0.0000000	.9996774
561.00- 578.00	0.0	0.0000000	.9996774
578.00- 595.00	0.0	0.0000000	.9996774
595.00- 612.00	0.0	0.0000000	.9996774
612.00- 629.00	0.0	0.0000000	.9996774
629.00- 646.00	0.0	0.0000000	.9996774
646.00- 663.00	1.0	.0003226	1.0000000
663.00- 680.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
 MEAN (FLOW) = 12.2
 STD. DEV (FLOW) = 22.56

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - SEPTEMBER
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 12.00	2686.0	.8953333	.8953333
12.00- 24.00	104.0	.0346667	.9299999
24.00- 36.00	59.0	.0196667	.9496666
36.00- 48.00	39.0	.0130000	.9626666
48.00- 60.00	33.0	.0110000	.9736667
60.00- 72.00	18.0	.0060000	.9796667
72.00- 84.00	10.0	.0033333	.9830000
84.00- 96.00	6.0	.0020000	.9850000
96.00- 108.00	6.0	.0020000	.9870000
108.00- 120.00	5.0	.0016667	.9886667
120.00- 132.00	6.0	.0020000	.9906666
132.00- 144.00	6.0	.0020000	.9926667
144.00- 156.00	5.0	.0016667	.9943334
156.00- 168.00	2.0	.0006667	.9950000
168.00- 180.00	2.0	.0006667	.9956666
180.00- 192.00	7.0	.0023333	.9980000
192.00- 204.00	2.0	.0006667	.9986666
204.00- 216.00	0.0	0.0000000	.9986666
216.00- 228.00	0.0	0.0000000	.9986666
228.00- 240.00	1.0	.0003333	.9990000
240.00- 252.00	0.0	0.0000000	.9990000
252.00- 264.00	0.0	0.0000000	.9990000
264.00- 276.00	1.0	.0003333	.9993334
276.00- 288.00	0.0	0.0000000	.9993334
288.00- 300.00	0.0	0.0000000	.9993334
300.00- 312.00	1.0	.0003333	.9996667
312.00- 324.00	0.0	0.0000000	.9996667
324.00- 336.00	0.0	0.0000000	.9996667
336.00- 348.00	0.0	0.0000000	.9996667
348.00- 360.00	0.0	0.0000000	.9996667
360.00- 372.00	0.0	0.0000000	.9996667
372.00- 384.00	0.0	0.0000000	.9996667
384.00- 396.00	0.0	0.0000000	.9996667
396.00- 408.00	0.0	0.0000000	.9996667
408.00- 420.00	0.0	0.0000000	.9996667
420.00- 432.00	0.0	0.0000000	.9996667
432.00- 444.00	0.0	0.0000000	.9996667
444.00- 456.00	1.0	.0003333	1.0000000
456.00- 468.00	0.0	0.0000000	1.0000000
468.00- 480.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3000.0
 MEAN (FLOW) = 10.9
 STD. DEV (FLOW) = 22.24

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - OCTOBER
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- 10.00	3001.0	.9680645	.9680645
10.00- 20.00	71.0	.0229032	.9909678
20.00- 30.00	20.0	.0064516	.9974194
30.00- 40.00	1.0	.0003226	.9977419
40.00- 50.00	3.0	.0009677	.9987097
50.00- 60.00	2.0	.0006452	.9993548
60.00- 70.00	0.0	0.0000000	.9993548
70.00- 80.00	1.0	.0003226	.9996774
80.00- 90.00	0.0	0.0000000	.9996774
90.00- 100.00	0.0	0.0000000	.9996774
100.00- 110.00	0.0	0.0000000	.9996774
110.00- 120.00	0.0	0.0000000	.9996774
120.00- 130.00	0.0	0.0000000	.9996774
130.00- 140.00	0.0	0.0000000	.9996774
140.00- 150.00	0.0	0.0000000	.9996774
150.00- 160.00	0.0	0.0000000	.9996774
160.00- 170.00	0.0	0.0000000	.9996774
170.00- 180.00	0.0	0.0000000	.9996774
180.00- 190.00	0.0	0.0000000	.9996774
190.00- 200.00	0.0	0.0000000	.9996774
200.00- 210.00	0.0	0.0000000	.9996774
210.00- 220.00	0.0	0.0000000	.9996774
220.00- 230.00	0.0	0.0000000	.9996774
230.00- 240.00	0.0	0.0000000	.9996774
240.00- 250.00	0.0	0.0000000	.9996774
250.00- 260.00	0.0	0.0000000	.9996774
260.00- 270.00	0.0	0.0000000	.9996774
270.00- 280.00	0.0	0.0000000	.9996774
280.00- 290.00	0.0	0.0000000	.9996774
290.00- 300.00	0.0	0.0000000	.9996774
300.00- 310.00	0.0	0.0000000	.9996774
310.00- 320.00	0.0	0.0000000	.9996774
320.00- 330.00	0.0	0.0000000	.9996774
330.00- 340.00	0.0	0.0000000	.9996774
340.00- 350.00	0.0	0.0000000	.9996774
350.00- 360.00	0.0	0.0000000	.9996774
360.00- 370.00	1.0	.0003226	1.0000000
370.00- 380.00	0.0	0.0000000	1.0000000
380.00- 390.00	0.0	0.0000000	1.0000000
390.00- 400.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
 MEAN (FLOW) = 5.6
 STD. DEV (FLOW) = 7.17

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - NOVEMBER
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .30	0.0	0.0000000	0.0000000
.30- .60	30.0	.0100000	.0100000
.60- .90	40.0	.0133333	.0233333
.90- 1.20	152.0	.0506667	.0740000
1.20- 1.50	304.0	.1013333	.1753333
1.50- 1.80	494.0	.1646667	.3400000
1.80- 2.10	584.0	.1946667	.5346667
2.10- 2.40	595.0	.1983333	.7330000
2.40- 2.70	405.0	.1350000	.8680000
2.70- 3.00	243.0	.0810000	.9490000
3.00- 3.30	83.0	.0276667	.9766667
3.30- 3.60	30.0	.0100000	.9866667
3.60- 3.90	11.0	.0036667	.9903333
3.90- 4.20	7.0	.0023333	.9926667
4.20- 4.50	6.0	.0020000	.9946667
4.50- 4.80	4.0	.0013333	.9960001
4.80- 5.10	3.0	.0010000	.9970000
5.10- 5.40	1.0	.0003333	.9973333
5.40- 5.70	2.0	.0006667	.9980000
5.70- 6.00	0.0	0.0000000	.9980000
6.00- 6.30	2.0	.0006667	.9986666
6.30- 6.60	1.0	.0003333	.9990000
6.60- 6.90	0.0	0.0000000	.9990000
6.90- 7.20	1.0	.0003333	.9993334
7.20- 7.50	0.0	0.0000000	.9993334
7.50- 7.80	0.0	0.0000000	.9993334
7.80- 8.10	1.0	.0003333	.9996667
8.10- 8.40	0.0	0.0000000	.9996667
8.40- 8.70	0.0	0.0000000	.9996667
8.70- 9.00	0.0	0.0000000	.9996667
9.00- 9.30	0.0	0.0000000	.9996667
9.30- 9.60	0.0	0.0000000	.9996667
9.60- 9.90	0.0	0.0000000	.9996667
9.90- 10.20	0.0	0.0000000	.9996667
10.20- 10.50	1.0	.0003333	1.0000000
10.50- 10.80	0.0	0.0000000	1.0000000
10.80- 11.10	0.0	0.0000000	1.0000000
11.10- 11.40	0.0	0.0000000	1.0000000
11.40- 11.70	0.0	0.0000000	1.0000000
11.70- 12.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3000.0
 MEAN (FLOW) = 2.1
 STD. DEV (FLOW) = .67

MONTHLY SIMULATION RUNOFF FREQUENCY
 SUMMARY RESULT FOR THE MONTH - DECEMBER
 RUNOFF IN WATERSHED MM

CLASS	FREQUENCY	PDF	CDF
0.00- .60	0.0	0.0000000	0.0000000
.60- 1.20	39.0	.0125806	.0125806
1.20- 1.80	69.0	.0222581	.0348387
1.80- 2.40	279.0	.0900000	.1248387
2.40- 3.00	480.0	.1548387	.2796774
3.00- 3.60	692.0	.2232258	.5029032
3.60- 4.20	666.0	.2148387	.7177420
4.20- 4.80	487.0	.1570968	.8748387
4.80- 5.40	249.0	.0803226	.9551613
5.40- 6.00	69.0	.0222581	.9774194
6.00- 6.60	23.0	.0074194	.9848387
6.60- 7.20	4.0	.0012903	.9861290
7.20- 7.80	7.0	.0022581	.9883871
7.80- 8.40	4.0	.0012903	.9896774
8.40- 9.00	3.0	.0009677	.9906452
9.00- 9.60	4.0	.0012903	.9919355
9.60- 10.20	4.0	.0012903	.9932258
10.20- 10.80	3.0	.0009677	.9941936
10.80- 11.40	0.0	0.0000000	.9941936
11.40- 12.00	2.0	.0006452	.9948387
12.00- 12.60	1.0	.0003226	.9951613
12.60- 13.20	3.0	.0009677	.9961290
13.20- 13.80	1.0	.0003226	.9964516
13.80- 14.40	3.0	.0009677	.9974194
14.40- 15.00	0.0	0.0000000	.9974194
15.00- 15.60	1.0	.0003226	.9977419
15.60- 16.20	2.0	.0006452	.9983871
16.20- 16.80	0.0	0.0000000	.9983871
16.80- 17.40	0.0	0.0000000	.9983871
17.40- 18.00	0.0	0.0000000	.9983871
18.00- 18.60	1.0	.0003226	.9987097
18.60- 19.20	1.0	.0003226	.9990323
19.20- 19.80	3.0	.0009677	1.0000000
19.80- 20.40	0.0	0.0000000	1.0000000
20.40- 21.00	0.0	0.0000000	1.0000000
21.00- 21.60	0.0	0.0000000	1.0000000
21.60- 22.20	0.0	0.0000000	1.0000000
22.20- 22.80	0.0	0.0000000	1.0000000
22.80- 23.40	0.0	0.0000000	1.0000000
23.40- 24.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 3100.0
 MEAN (FLOW) = 3.7
 STD. DEV (FLOW) = 1.44

SUMMARY OF TOTAL MONSOON SEASON
RUNOFF VOLUME IN MM

CLASS	FREQUENCY	PDF	CDF
50.00- 120.00	3.0	.0300000	.0300000
120.00- 190.00	2.0	.0200000	.0500000
190.00- 260.00	2.0	.0200000	.0700000
260.00- 330.00	6.0	.0600000	.1300000
330.00- 400.00	3.0	.0300000	.1600000
400.00- 470.00	7.0	.0700000	.2300000
470.00- 540.00	5.0	.0500000	.2800000
540.00- 610.00	8.0	.0800000	.3600000
610.00- 680.00	5.0	.0500000	.4100000
680.00- 750.00	10.0	.1000000	.5100000
750.00- 820.00	6.0	.0600000	.5700001
820.00- 890.00	3.0	.0300000	.6000000
890.00- 960.00	6.0	.0600000	.6600000
960.00-1030.00	5.0	.0500000	.7100000
1030.00-1100.00	5.0	.0500000	.7600000
1100.00-1170.00	4.0	.0400000	.8000000
1170.00-1240.00	6.0	.0600000	.8600000
1240.00-1310.00	2.0	.0200000	.8800000
1310.00-1380.00	4.0	.0400000	.9200000
1380.00-1450.00	1.0	.0100000	.9299999
1450.00-1520.00	1.0	.0100000	.9400001
1520.00-1590.00	0.0	0.0000000	.9400001
1590.00-1660.00	0.0	0.0000000	.9400001
1660.00-1730.00	2.0	.0200000	.9600000
1730.00-1800.00	1.0	.0100000	.9700000
1800.00-1870.00	0.0	0.0000000	.9700000
1870.00-1940.00	2.0	.0200000	.9900000
1940.00-2010.00	0.0	0.0000000	.9900000
2010.00-2080.00	0.0	0.0000000	.9900000
2080.00-2150.00	1.0	.0100000	1.0000000
2150.00-2220.00	0.0	0.0000000	1.0000000
2220.00-2290.00	0.0	0.0000000	1.0000000
2290.00-2360.00	0.0	0.0000000	1.0000000
2360.00-2430.00	0.0	0.0000000	1.0000000
2430.00-2500.00	0.0	0.0000000	1.0000000
2500.00-2570.00	0.0	0.0000000	1.0000000
2570.00-2640.00	0.0	0.0000000	1.0000000
2640.00-2710.00	0.0	0.0000000	1.0000000
2710.00-2780.00	0.0	0.0000000	1.0000000
2780.00-2850.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 100.0
MEAN (FLOW) = 812.3
STD. DEV (FLOW) = 428.31

SUMMARY OF TOATL YEARLY RUNOFF VOLUME IN MM

CLASS	FREQUENCY	PDF	CDF
100.00- 165.00	1.0	.0100000	.0100000
165.00- 230.00	1.0	.0100000	.0200000
230.00- 295.00	2.0	.0200000	.0400000
295.00- 360.00	1.0	.0100000	.0500000
360.00- 425.00	2.0	.0200000	.0700000
425.00- 490.00	4.0	.0400000	.1100000
490.00- 555.00	2.0	.0200000	.1300000
555.00- 620.00	2.0	.0200000	.1500000
620.00- 685.00	2.0	.0200000	.1700000
685.00- 750.00	8.0	.0800000	.2500000
750.00- 815.00	3.0	.0300000	.2800000
815.00- 880.00	7.0	.0700000	.3500000
880.00- 945.00	4.0	.0400000	.3900000
945.00-1010.00	9.0	.0900000	.4800000
1010.00-1075.00	4.0	.0400000	.5200000
1075.00-1140.00	5.0	.0500000	.5700001
1140.00-1205.00	5.0	.0500000	.6200000
1205.00-1270.00	5.0	.0500000	.6700000
1270.00-1335.00	6.0	.0600000	.7300000
1335.00-1400.00	4.0	.0400000	.7700000
1400.00-1465.00	1.0	.0100000	.7800000
1465.00-1530.00	5.0	.0500000	.8300000
1530.00-1595.00	5.0	.0500000	.8800000
1595.00-1660.00	2.0	.0200000	.9000000
1660.00-1725.00	0.0	0.0000000	.9000000
1725.00-1790.00	2.0	.0200000	.9200000
1790.00-1855.00	1.0	.0100000	.9299999
1855.00-1920.00	0.0	0.0000000	.9299999
1920.00-1985.00	2.0	.0200000	.9500000
1985.00-2050.00	1.0	.0100000	.9600000
2050.00-2115.00	1.0	.0100000	.9700000
2115.00-2180.00	0.0	0.0000000	.9700000
2180.00-2245.00	0.0	0.0000000	.9700000
2245.00-2310.00	2.0	.0200000	.9900000
2310.00-2375.00	1.0	.0100000	1.0000000
2375.00-2440.00	0.0	0.0000000	1.0000000
2440.00-2505.00	0.0	0.0000000	1.0000000
2505.00-2570.00	0.0	0.0000000	1.0000000
2570.00-2635.00	0.0	0.0000000	1.0000000
2635.00-2700.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 100.0
 MEAN (FLOW) = 1090.6
 STD. DEV (FLOW) = 472.64

SUMMARY OF MAXIMUM DAILY FLOW IN CU. M/SEC

CLASS	FREQUENCY	PDF	CDF
0.00- 120.00	9.0	.0900000	.0900000
120.00- 240.00	5.0	.0500000	.1400000
240.00- 360.00	6.0	.0600000	.2000000
360.00- 480.00	8.0	.0800000	.2800000
480.00- 600.00	9.0	.0900000	.3700000
600.00- 720.00	8.0	.0800000	.4500000
720.00- 840.00	5.0	.0500000	.5000000
840.00- 960.00	6.0	.0600000	.5599999
960.00-1080.00	5.0	.0500000	.6100000
1080.00-1200.00	4.0	.0400000	.6500000
1200.00-1320.00	7.0	.0700000	.7400000
1320.00-1440.00	4.0	.0400000	.7800000
1440.00-1560.00	3.0	.0300000	.8099999
1560.00-1680.00	4.0	.0400000	.8500000
1680.00-1800.00	2.0	.0200000	.8700000
1800.00-1920.00	1.0	.0100000	.8800000
1920.00-2040.00	2.0	.0200000	.9000000
2040.00-2160.00	1.0	.0100000	.9100000
2160.00-2280.00	1.0	.0100000	.9200000
2280.00-2400.00	2.0	.0200000	.9400000
2400.00-2520.00	2.0	.0200000	.9600000
2520.00-2640.00	0.0	0.0000000	.9600000
2640.00-2760.00	0.0	0.0000000	.9600000
2760.00-2880.00	1.0	.0100000	.9700000
2880.00-3000.00	1.0	.0100000	.9800000
3000.00-3120.00	0.0	0.0000000	.9800000
3120.00-3240.00	0.0	0.0000000	.9800000
3240.00-3360.00	0.0	0.0000000	.9800000
3360.00-3480.00	0.0	0.0000000	.9800000
3480.00-3600.00	0.0	0.0000000	.9800000
3600.00-3720.00	0.0	0.0000000	.9800000
3720.00-3840.00	0.0	0.0000000	.9800000
3840.00-3960.00	0.0	0.0000000	.9800000
3960.00-4080.00	0.0	0.0000000	.9800000
4080.00-4200.00	1.0	.0100000	.9900000
4200.00-4320.00	0.0	0.0000000	.9900000
4320.00-4440.00	1.0	.0100000	1.0000000
4440.00-4560.00	0.0	0.0000000	1.0000000
4560.00-4680.00	0.0	0.0000000	1.0000000
4680.00-4800.00	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 100.0
 MEAN (FLOW) = 1010.4
 STD. DEV (FLOW) = 814.18

SUMMARY OF EXTREMELY LOW FLOWS IN CU.M/SEC
 (Flows less than 2.5 cubic meter per sec)

CLASS	FREQUENCY	PDF	CDF
0.00- .07	0.0	0.0000000	0.0000000
.07- .14	0.0	0.0000000	0.0000000
.14- .21	0.0	0.0000000	0.0000000
.21- .28	1.0	.0100000	.0100000
.28- .35	0.0	0.0000000	.0100000
.35- .42	0.0	0.0000000	.0100000
.42- .49	1.0	.0100000	.0200000
.49- .56	0.0	0.0000000	.0200000
.56- .63	1.0	.0100000	.0300000
.63- .70	1.0	.0100000	.0400000
.70- .77	2.0	.0200000	.0600000
.77- .84	1.0	.0100000	.0700000
.84- .91	3.0	.0300000	.1000000
.91- .98	4.0	.0400000	.1400000
.98- 1.05	3.0	.0300000	.1700000
1.05- 1.12	9.0	.0900000	.2600000
1.12- 1.19	6.0	.0600000	.3200000
1.19- 1.26	5.0	.0500000	.3700000
1.26- 1.33	6.0	.0600000	.4300000
1.33- 1.40	10.0	.1000000	.5300000
1.40- 1.47	4.0	.0400000	.5700001
1.47- 1.54	7.0	.0700000	.6400000
1.54- 1.61	2.0	.0200000	.6600000
1.61- 1.68	9.0	.0900000	.7500000
1.68- 1.75	4.0	.0400000	.7900000
1.75- 1.82	6.0	.0600000	.8500000
1.82- 1.89	6.0	.0600000	.9100000
1.89- 1.96	4.0	.0400000	.9500000
1.96- 2.03	2.0	.0200000	.9700000
2.03- 2.10	0.0	0.0000000	.9700000
2.10- 2.17	0.0	0.0000000	.9700000
2.17- 2.24	1.0	.0100000	.9800000
2.24- 2.31	1.0	.0100000	.9900000
2.31- 2.38	0.0	0.0000000	.9900000
2.38- 2.45	1.0	.0100000	1.0000000
2.45- 2.52	0.0	0.0000000	1.0000000
2.52- 2.59	0.0	0.0000000	1.0000000
2.59- 2.66	0.0	0.0000000	1.0000000
2.66- 2.73	0.0	0.0000000	1.0000000
2.73- 2.80	0.0	0.0000000	1.0000000

TOTAL (EVENTS) = 100.0
 MEAN (FLOW) = 1.4
 STD. DEV (FLOW) = .40

SUMMARY OF CONSECUTIVE DAYS OF FLOWS LESS THAN 1.5 C.M/S

CONSECUTIVE DAYS	FREQUENCY	PDF	CDF
2	5	0.06329	0.06329
3	6	0.07595	0.13924
4	5	0.06329	0.20253
5	2	0.02532	0.22785
6	1	0.01266	0.24051
7	2	0.02532	0.26582
8	2	0.02532	0.29114
9	4	0.05406	0.34177
10	1	0.01266	0.35443
11	3	0.03797	0.39241
12	2	0.02532	0.41772
13	1	0.01266	0.43038
14	3	0.03797	0.46835
16	2	0.02532	0.49367
17	2	0.02532	0.51899
18	3	0.03797	0.55696
19	1	0.01266	0.56962
20	1	0.01266	0.58228
21	1	0.01266	0.59494
22	1	0.01266	0.60759
23	1	0.01266	0.62025
24	2	0.02532	0.64557
25	1	0.01266	0.65823
26	3	0.03797	0.69620
30	2	0.02532	0.72152
31	3	0.03797	0.75949
32	1	0.01266	0.77215
33	1	0.01266	0.78481
34	2	0.02532	0.81013
37	1	0.01266	0.82278
43	2	0.02532	0.84810
45	1	0.01266	0.86076
49	1	0.01266	0.87342
50	1	0.01266	0.88608
51	1	0.01266	0.89873
62	1	0.01266	0.91139
66	1	0.01266	0.92405
67	1	0.01266	0.93671
71	1	0.01266	0.94937
77	1	0.01266	0.96203
86	1	0.01266	0.97469
111	1	0.01266	0.98734
157	1	0.01266	1.00000

MEAN CONSECUTIVE DAYS = 24.42
STANDARD DEVIATION (DAYS) = 26.55

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