WATER BALANCE CALCULATIONS, WATER USE EFFICIENCY, AND ABOVEGROUND NET PRODUCTION
L. J. Lane and J. J. Stone
The authors are Hydrologists, USDA-ARS, Tucson, Arizona 85705

Abstract
A discrete form of the water balance equation is used to illustrate the interaction among precipitation, runoff, percolation below the root zone, bare soil evaporation, plant transpiration, and plant available soil moisture. Under rangeland conditions, water availability is often the limiting factor in plant survival and growth. Therefore, the water balance equation is used, together with soils data and water use efficiency factors, to estimate annual aboveground net primary production of perennial vegetation.

INTRODUCTION
Arid and semiarid rangelands are extensive (e.g., Branson et al., 1981), and in these regions, water availability is most often the limiting factor in plant production. Water availability is also the most important environmental factor controlling survival of range plants (e.g., Brown, 1977). Therefore, water balance calculations are essential in soil-water-plant relationships studies. Plant water use, as a component of total evapotranspiration (ET), affects the water balance and soil moisture content, and thus, infiltration and runoff. As will be discussed later, soil moisture often limits the rate of ET so that consideration of feedback is necessary in water balance calculations.

If we assume no net subsurface water movement in the horizontal direction and a limited rooting depth well above the permanent water table, then a discrete form of the water balance equation in terms of unit area is:

\[ \Delta S = P - Q - ET - L \]  \hspace{1cm} (1)

where:
- \( S \) = soil water content (mm representing units of volume per unit area),
- \( \Delta t \) = time for the calculation period (e.g., hr, day, month, etc.),
- \( P \) = rainfall depth for the time interval (mm),
- \( Q \) = net runoff for the area (mm),
- \( ET \) = combined evaporation and plant transpiration (mm), and
- \( L \) = seepage or percolation below the root zone (mm).

The rate of ET in Eq. (1) depends upon atmospheric demand (a potential rate), soil texture, and vegetation characteristics (e.g., the leaf area index (LAI)) when soil moisture (SM) is nonlimiting, and upon soil water content, soil texture, atmospheric demand, and vegetation characteristics when water is limiting. Veihmeyer and Hendrickson (1955) suggested that the ratio of actual evapotranspiration (AET) to potential evapotranspiration (PET) stayed about constant as the soil moisture decreased from field capacity to near the wilting point. Thornthwaite and Mather (1955) suggested a linear relationship between the ratio AET/PET and available soil water. Other investigators have proposed nonlinear relationships, as summarized by Hanson (1973, Fig. 3, p. 16). This relationship can be expressed as:

\[
\begin{align*}
\text{PET} & \quad \text{SM} \geq \text{SM}_1 \\
\text{AET} = f(\text{SM}, \text{PET}) & \quad \text{SM}_0 < \text{SM} < \text{SM}_1 \\
0 & \quad \text{SM} \leq \text{SM}_0
\end{align*}
\]  \hspace{1cm} (2)

where \( \text{SM}_1 \) is a limiting soil water content (between wilting point and field capacity) where soil water content above \( \text{SM}_1 \) does not influence AET, and \( \text{SM}_0 \) is the permanent wilting point where plant available soil water is exhausted. The function \( f \) controls the ratio of AET to PET when soil moisture is between the limits \( \text{SM}_1 \) and \( \text{SM}_0 \). The parameters \( \text{SM}_1 \) and \( \text{SM}_0 \) are often expressed in terms of the soil water content at -1/3 and -15 bars, respectively.

Leaf area index (LAI) is the leaf area per unit area of land surface (e.g., \( \text{m}^2/\text{m}^2 \)), and is often used in models to relate evapotranspiration and plant growth (e.g., Chang, 1968; Stern, 1965; Penman, 1963; and Brougham, 1956). Hanson (1973, Fig. 2, p. 15) illustrates two relationships between LAI and AET/PET. For LAI = 0 (bare soil), the ratio AET/PET varied from zero to about 0.26, depending upon the
relationships assumed and on the potential rate, PET. For LAI = 1.0, the ratio was about 0.5, and for LAI > 2.5, the ratio approached 1.0. Therefore, LAI can be used to estimate the magnitude of AET and to partition total AET into bare soil evaporation, E, and plant transpiration, T (e.g., Parton et al., 1978; Ritchie, 1972; Ritchie and Burnett, 1971; Ritchie et al., 1976; and Smith and Williams, 1980).

Ritchie (1972) describes the two stages of bare soil evaporation: stage 1 (constant rate limited by supply of energy to the surface) and stage 2 (falling rate stage where water movement to evaporation sites at the surface depends upon soil properties). The volume of evaporation during stage 1, \( u \), is:

\[
u = a (c - b) d\tag{3}
\]

where \( u \) is in mm, \( a, b, \) and \( d \) are fitting coefficients (e.g., \( a = 9, b = 3, \) and \( d = 0.42 \) fit Ritchie's data), and \( c \) is a parameter dependent upon soil texture. Stage 2 evaporation rate is:

\[
E_{S2} = \frac{1}{2} c t^{-1/2}\tag{4}
\]

where \( E_{S2} \) is evaporation in mm/day, \( t \) is time in days since initiation of stage 2, and \( c \) is the soil texture parameter. The total soil evaporation during a drying period, \( t \), can then be approximated as:

\[
E_t = u + \int E_{S2} dt\tag{5}
\]

which, as a first order approximation for a finite time period, can be represented as:

\[
E_t = u + Au = (1 + A)u\tag{6}
\]

where \( A \) is the ratio of stage 2 to stage 1 evaporation volume.

Now, if this volume is assumed to represent water extracted when soil moisture is between field capacity, FC, and wilting point, WP, then by continuity, we can write:

\[
y(FC - WP) = (1 + A)u\tag{7}
\]

where \( y \) is the effective depth influenced by bare soil evaporation. Values of \( c \), in Eq. (3) and (4), are reported by Ritchie (1972) and by Jackson et al. (1976). The value of \( E_t \), in Eq. (5), was compared with data from Ritchie (1972), Hillel (1976), and Jackson et al. (1976) to estimate the value of \( A \) in Eq. (6). Data presented by Jackson et al. (1973, Fig. 8) show the locus of the zero flux surface (vertical direction) in a soil profile (16 to 38 cm with a mean of about 20 cm) which provided experimental data to compare with \( y \) values calculated from Eq. (7). Finally, SCS (1982) presents additional data on values of \( c \) in Eq. (4) vs. soil texture.

Rawls et al. (1982) analyzed data from over 1300 soils in 32 states to generalize soil water properties as a function of textural class. Interpolated values from the texture triangle (from Rawls et al. (1982) and from SCS (1982)) were combined to produce the gross soil properties summarized in Tables 1 and 2. Table 1 lists 12 texture classes, representative proportions of clay, silt, and sand, mean rates of saturated hydraulic conductivity, and soil evaporation parameters. Notice that the data labeled "avg" represent a central, or representative, value for mean hydraulic conductivity and \( c \) in Table 1, and for mean water holding capacities in Table 2. The columns labeled "low" and "high" refer to low and high estimates on the mean, not to the maximum expected range for the given parameter. Moreover, the parameters in Tables 1 and 2 were derived predominately from agricultural soils, and probably do not represent desert soils with high gravel content.

<table>
<thead>
<tr>
<th>Soil texture class</th>
<th>Representative composition</th>
<th>Saturated hydraulic conductivity (cm/hr)</th>
<th>Bare soil evaporation parameter c(mm/day 1/2)</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Clay</td>
<td>Silt</td>
<td>Sand</td>
</tr>
<tr>
<td>Sand</td>
<td>3</td>
<td>7</td>
<td>90</td>
</tr>
<tr>
<td>Loamy sand</td>
<td>5</td>
<td>15</td>
<td>80</td>
</tr>
<tr>
<td>Sandy loam</td>
<td>10</td>
<td>20</td>
<td>70</td>
</tr>
<tr>
<td>Loam</td>
<td>20</td>
<td>40</td>
<td>40</td>
</tr>
<tr>
<td>Silt loam</td>
<td>15</td>
<td>65</td>
<td>20</td>
</tr>
<tr>
<td>Silt</td>
<td>5</td>
<td>87</td>
<td>8</td>
</tr>
<tr>
<td>Sandy clay loam</td>
<td>30</td>
<td>10</td>
<td>60</td>
</tr>
<tr>
<td>Clay loam</td>
<td>35</td>
<td>35</td>
<td>30</td>
</tr>
<tr>
<td>Silty clay loam</td>
<td>35</td>
<td>55</td>
<td>10</td>
</tr>
<tr>
<td>Sandy clay</td>
<td>45</td>
<td>5</td>
<td>50</td>
</tr>
<tr>
<td>Silty clay</td>
<td>45</td>
<td>50</td>
<td>5</td>
</tr>
<tr>
<td>Clay</td>
<td>65</td>
<td>20</td>
<td>15</td>
</tr>
</tbody>
</table>
At each time step, the sum of soil evaporation, \( E \) from Eq. (3), where \( T_0 \) is computed from Eq. (9), SM is soil moisture < 0.25 FC, and computed as:

\[
T = \left( \frac{T_0}{(SM)} \right) / (0.25 \text{ FC})
\]

(10)

where \( T_0 \) is computed from Eq. (9), SM is soil moisture ≤ 0.25 FC, and FC is field capacity of the soil. At each time step, the sum of soil evaporation, \( E \) from Eq. (3), (4), and (8), and transpiration, \( T \) from Eq. (9) and (10), cannot exceed \( E_0 \) calculated from temperature and radiation. Therefore, total AET is partitioned into soil evaporation and plant transpiration according to the value of LAI. This value of LAI varies throughout the season to reflect crop growth stage or plant phenological stages.

Therefore, to estimate AET and plant water use using a model such as Ritchie's equations in the CREAMS model, it is necessary to know soil properties (Tables 1 and 2), plant rooting depth, the seasonal LAI, and the other data used to solve the water balance equation. The CREAMS model is typical of many water balance-climate models in that it uses LAI to partition AET into \( E \) and \( T \), and AET is limited by available soil moisture, as described by Eq. (2). More complex models (e.g., the ELM model, Innis (1978) and the EPIC model, Williams et al. (1983)) explicitly include plant growth components and direct feedback with the water balance components. In the absence of sufficient information to estimate the parameters of these complex models, or if the objective is to obtain first-order approximations for plant production when water is limiting, then simple water balance-plant production models (e.g., Wight and Hanks (1981) or Lane et al. (1983)) can be used to relate the water balance and plant production.
Selected WUE data, from greenhouse or carefully controlled plot studies, are listed in Table 3. The first column of Table 3 lists the vegetation types, the second column lists WUE in units of g of above-ground dry matter production per kg of water use, which is the same as g d.w./m² over mm H₂O/m², or grams dry weight per square meter per mm of plant transpiration per square meter. The third column of Table 3 provides comments, and the fourth column lists the appropriate reference. The references were selected to be representative of results obtained from the literature for arid and semiarid rangelands. Data collected by McGinnies and Arnold (1939) were from 75-liter galvanized Witt cans filled with 125 to 145 kg of soil. Because of the rather large-sized soil containers and duration of the experiment (1932-1936), these data are probably among the most reliable of those shown in Table 3. These data represent annual means, with summer WUE being from near to one and a half times as large as the annual values. The mean annual WUE for the six desert grasses was 1.8 g d.w./kg H₂O and ranged from 1.7 to over 1.9. In contrast, the mean annual WUE for the desert shrubs and trees was 0.57 g d.w./kg H₂O and ranged from 0.42 to 1.30. Therefore, these data suggest that, on the average, the selected desert grasses have a WUE some three times higher than the selected trees and shrubs (the range of the ratio of WUE grass to WUE trees and shrubs was 1.3 to 4.6). Dwyer and DeGarmo (1970) used pots (18 cm diameter by 25 cm deep) to study the WUE of four desert grasses and four desert shrubs. Means for the desert grasses ranged from 1.0 to 1.09, with an overall mean of 1.03 g d.w./kg H₂O. The means for Dwyer and DeGarmo’s four shrubs ranged from 0.23 to 0.57, with an overall mean of 0.28 g d.w./kg H₂O. These data suggest that, on the average, the desert grasses have a WUE some 3.7 times higher than the selected desert shrubs.

<table>
<thead>
<tr>
<th>Vegetation</th>
<th>WUE</th>
<th>Comments</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Desert grasslands</td>
<td>1.70 - 1.92</td>
<td>Annual values for 6 species</td>
<td>McGinnies and Arnold (1939)</td>
</tr>
<tr>
<td>Plains grasslands</td>
<td>1.42 - 2.09</td>
<td>Annual values for 3 gramma grasses</td>
<td>McGinnies and Arnold (1939)</td>
</tr>
<tr>
<td>Southern tall grass</td>
<td>1.41 - 2.12</td>
<td>Annual values for 3 species</td>
<td>McGinnies and Arnold (1939)</td>
</tr>
<tr>
<td>Desert shrubs and trees</td>
<td>0.42 - 1.30</td>
<td>Annual values for 6 species</td>
<td>McGinnies and Arnold (1939)</td>
</tr>
<tr>
<td>Desert grasses</td>
<td>1.00 - 1.09</td>
<td>Greenhouse study of 4 species, 120 days</td>
<td>Dwyer and DeGarmo (1970)</td>
</tr>
<tr>
<td>Desert shrubs</td>
<td>0.23 - 0.57</td>
<td>Greenhouse study of 4 species, 120 days</td>
<td>Dwyer and DeGarmo (1970)</td>
</tr>
<tr>
<td>Native and seeded within the pinyon-Juniper type</td>
<td>Brief survey of values reported in the literature</td>
<td>Branson et al. (1981) from Gifford (1976)</td>
<td></td>
</tr>
<tr>
<td>Selected grasses</td>
<td>0.31 - 2.95</td>
<td>Brief survey of values reported in the literature</td>
<td>Teare (1977)</td>
</tr>
</tbody>
</table>

The ratio of mean WUE for the desert grasses, in McGinnies and Arnold’s study to Dwyer and DeGarmo’s, was 1.81/1.03 = 1.75, and, for the trees and shrubs, a similar ratio was 0.57/0.28 = 2.04. The reasons for these differences between the two experiments are not clear. One possible reason might be distortions due to root bound plants in the small pots. This situation, if it did in fact occur, would tend to lower the mean water use efficiency. Finally, Branson et al. (1981) and Teare (1977) present brief summaries of water use efficiency data selected from a variety of sources. Data, such as those presented in Table 3, may be used to estimate aboveground production. However, the magnitude of the variability in WUE values, as discussed above, should be kept in mind.

**ANNUAL ABOVEGROUND NET PRIMARY PRODUCTION**

Annual aboveground net primary production (ANPP) is the total aboveground dry weight (d.w.) biomass produced per unit area in a growing season. Net refers to the gross production minus respiration per unit time of an individual plant, and primary is the sum of the net productions of all individual plants in a unit area (Whittaker and Macks, 1975). Total above and below ground net primary production of general ecosystem types is summarized by Whittaker and Likens (1975). They suggest values for desert and semidesert scrub of 10 - 250 g/m²/yr, with a mean value of 90. Corresponding estimates for woodland and shrubland range from 250 - 1200, with a mean of 700, while values for temperate grassland vary from 200 - 1500, with a mean of 600. These values are for total annual net primary productivity so that the above-ground values would be significantly less, depending on the root-shoot ratios. Lieth (1975) lists similar estimates for selected climatic regions, and Murphey (1975) lists estimates for tropical terrestrial ecosystems. Webb et al. (1975) present productivity data, including grasslands and desert ecosystems,
and Szarek (1979) summarizes data for four North American deserts.

Generalized relationships relating annual actual evapotranspiration, AET, and annual aboveground net primary production, ANPP, have been developed by several investigators (e.g., Rosenzweig, 1968 and Webb et al., 1978). Data on ANPP and AET for three North American deserts (Szarek, 1979 and Lane et al. 1983) and for a site in the Tunisian steppes (Floret et al., 1982) are shown in Fig. 1. Notice that, for the data shown, Rosenzweig's equation tends to provide an upper limit on ANPP, and that the Webb et al. equation seems to fit the overall trend, but underestimates for low values of AET < 120 mm.

![Graph showing the relationship between AET and ANPP for selected data from desert and semidesert areas and proposed general relationships.](image)

**Figure 1.** Relation between AET and ANPP for selected data from desert and semidesert areas and proposed general relationships.

An expanded plot of the Rock Valley, Nevada data are shown in Fig. 2. The curve labeled 1 is the equation from Webb et al. (1978), and the curve labeled 2 is a linear regression between ANPP and AET. The least squares equation derived for this particular data set provides a better fit than the generalized equation which is intended for broader application over wider ranges of AET. For comparison, annual actual transpiration, AT, is plotted against ANPP, as shown in the left portion of Fig. 2, identified as curve 3. If seasonal estimates of transpiration from a continuous simulation model are used, then the unexplained variance in ANPP is further reduced ($R^2 = .90$), as described by Lane et al. (1983). The slope of the line relating AT and ANPP can be used as an approximate estimate of the water use efficiency, WUE, provided the intercept is near zero. From curve 3, in Fig. 2, the WUE for Rock Valley data is about 0.7, which is well within the range of values suggested in Table 3 (WUE values of 0.42 to 1.30).

If the product of AT and WUE provides a first order approximation of ANPP (curve 3 in Fig. 2), then relationships, such as curves 1 and 2 in Fig. 2 and the curves shown in Fig. 1, provide a somewhat coarser approximation using AET as an approximate predictor for actual plant water use, AT. The relationship between ANPP and AET is inherently weaker than the one between ANPP and AT, because AET includes bare soil evaporation, and is thus an imperfect predictor for AT or plant water use. However, less information is required to estimate AET than is required to estimate AT, so that the choice of approximating relationships depends upon the degree of accuracy required, the amount of information available, and the intended scale of application.
SUMMARY

Water balance calculations are important in soil-water-plant relationships studies and essential to predict net primary production when water availability is the limiting factor. These circumstances are often realized on arid and semiarid rangelands so that the water balance equation can be used together with soils data and water use efficiency factors to estimate net primary production. Equations (1) and (2) describe a simple water balance in a discrete form, selected soil properties are summarized in Tables 1 and 2, and water use efficiency factors are presented in Table 3. Figures 1 and 2 illustrate relationships between actual evapotranspiration, actual transpiration, and annual aboveground net primary production. Depending upon scale of application, available information, and desired accuracy, relationships such as these can be used to estimate annual aboveground net primary production of perennial vegetation.

References Cited


