

Networking our science to characterize the state, vulnerabilities, and management opportunities of soil organic matter

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Abstract

Soil organic matter supports the Earth's ability to sustain terrestrial ecosystems, provide food and fiber, and retain the largest pool of actively cycling carbon (C). Over 75% of the soil organic carbon (SOC) in the top meter of soil is directly affected by human land use. Large land areas have lost SOC as a result of land use practices, yet there are compensatory opportunities to enhance land productivity and SOC storage in degraded lands through improved management practices. Large areas with and without intentional management are also being subjected to rapid changes in climate, making many SOC stocks vulnerable to losses by decomposition or disturbance. In order to quantify potential SOC losses or sequestration at field, regional, and global scales, measurements for detecting changes in SOC are needed. Such measurements and soil-management best practices should be based on well-established and emerging scientific understanding of processes of C stabilization and destabilization over various timescales, soil types, and spatial scales. As newly engaged members of the International Soil Carbon Network, we have identified gaps in data, modeling, and communication that underscore the need for an open, shared network to frame and guide the study of soil organic matter and C and their management for sustained production and climate regulation.

49 1. Introduction

50 Soil organic matter (SOM) governs many physical and chemical characteristics of soils, and is
51 one determinant of a soil's capacity for fertility, ecosystem productivity, and CO₂ sequestration.
52 Thus SOM, and its main constituent soil organic carbon (SOC), interacts with several aspects of
53 the Earth system and its services to society (Banwart et al., 2014), including food, fiber, water,
54 energy, cycling of C and nutrients, and biodiversity. Historically, estimates of global SOC stocks
55 ranged from 500 to over 3000 Pg, with recent estimates of ca. 2000 Pg to a depth of two meters
56 (Scharlemann et al. 2014; Batjes, 2016). Large land areas (up to 6 billion ha) are estimated to be in
57 some state of soil degradation (Gibbs and Salmon, 2015), associated in many cases with deficient
58 stocks of SOM. Increasing SOM content, and thus SOC storage, can improve the state of soil and
59 ecological sustainability, and can also contribute to climate change mitigation by capturing
60 atmospheric CO₂.

61 SOM research has traditionally been dominated by two scientific communities that have
62 been publishing in rather disparate types of journals (SM2a), one focused on soil science/soil
63 health and the other focused on the terrestrial C cycle /biogeochemistry. Soil health or quality is a
64 concept formalized in the 1990s to describe soil management practices that enhance the biological,
65 chemical, and physical processes of soil. Terrestrial carbon cycling refers to the exchange of land-
66 based C with atmospheric CO₂ and CH₄ . Owing to the very large and dynamic stocks of soil C
67 globally, the role of soils in climate regulation has been increasingly studied in context of
68 ecological and geological perspectives that link organic matter processing to carbon, nutrients,
69 productivity, hydrology, and landscape dynamics. As a result, conceptual frameworks and
70 computer models have become quite elegant and sophisticated in representing both site-based,
71 land management options and global scale syntheses. The goals of these communities are
72 converging and should not be pursued in isolation from each other. Together, the science
73 communities working on agricultural soils and those working on soil C cycling have an
74 opportunity to combine and transform our knowledge, databases, and mathematical frameworks
75 for the benefit of environmental health and humanity.

76 At the global scale, SOM is one of the largest and actively cycling C reservoirs (Cias et al.,
77 2013; Jackson et al. 2017) and direct human activities impact over 70% of C stocks in the upper
78 meter of soil. Globally soils store 1,300-1,500 Pg of C in the top meter (Fig. 1a; Batjes, 2016).
79 Much of this SOM is in lands impacted directly by cropping, grazing, and forestry practices, with
80 30% residing in lands only indirectly impacted by human activities such as peatlands and
81 permafrost soils (Hugelius et al., 2014; Köchy et al., 2015; Loisel et al., 2017). The distribution of
82 soils in managed lands follows the distribution of human land use (Fig. 1b, c) and overlaying the
83 estimated SOC stocks with human land-use data shows that the majority of near-surface SOC
84 stocks are directly affected by human activities today (Fig. 1c).

85 Efforts such as the ‘4-per-1000’ program, a global initiative to reduce atmospheric CO₂
86 through soil C sequestration (Minasny et al., 2017), demonstrate that many soils in managed
87 systems could offer an opportunity for climate regulation. While uncertainties are very large, it is
88 evident that land management practices can lead to C gains from 0.01 kg C yr⁻¹ up to 0.07 kg C yr⁻¹
89 (Minasny et al., 2017). Other additional estimates are in the same order of magnitude (Paustian et
90 al. 2016, Smith et al. 2007). If these numbers are applied across all Earth’s managed lands, there is
91 an opportunity to sequester several Pg C yr⁻¹ globally (fig. 1d). While not all lands are likely to be
92 managed consistently, this maximum estimate could potentially offset future C emissions from
93 permafrost (Koven et al., 2015) or the combined projected emissions from land use change and
94 agricultural management (Pugh et al., 2015; projected emissions in fig. 1d).

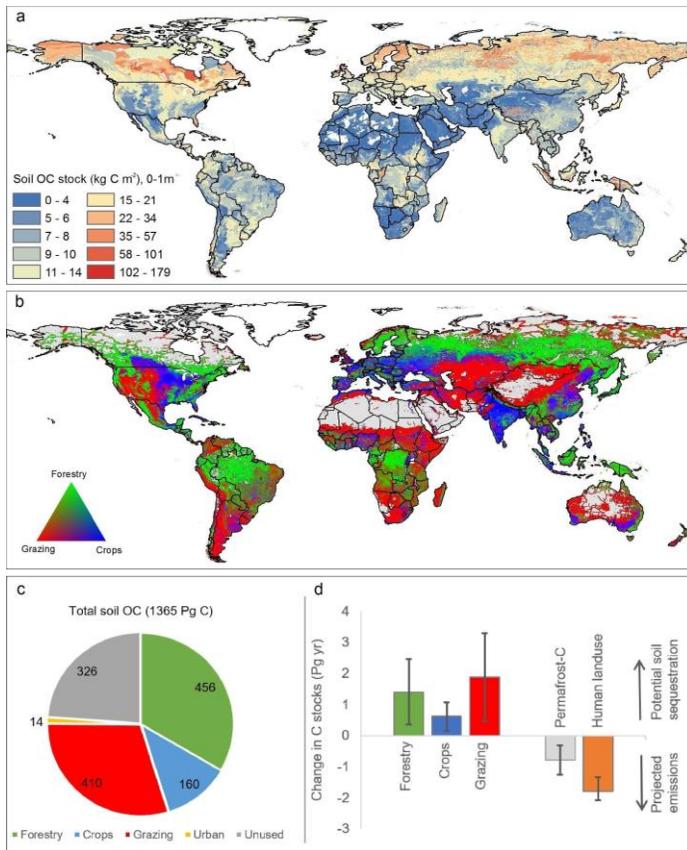
95 The ability to detect shifts in SOC and to potentially increase SOC storage is increasingly
96 important for scientific and societal challenges in the face of rapidly changing terrestrial
97 landscapes, yet detecting changes in SOM are as problematic to measure and predict as they are
98 important for climate and landuse planning. For example, estimates of future SOM dynamics
99 range widely, and recent compilations of soil radiocarbon demonstrate that global models
100 underestimate the transit time of C in soil, biasing estimates for soil C sequestration in future years
101 (He et al, 2016). Meanwhile conceptual frameworks for SOM stabilization are also changing,
102 challenging the science community to shift methods and measurements to test alternative models.

103 For example, paradigms and metrics of SOM_C (de)stabilization and storage have been shifting
104 (Schmidt et al., 2011 ; Lehmann and Kleber, 2015). Emerging paradigms de-emphasize the
105 chemical properties of SOM itself and focus more on mechanisms that isolate or stabilize C, such
106 as sorption of biopolymers and their decomposition products on mineral surfaces and the
107 entrapment of organic matter in aggregates. These and other recent developments call for model
108 development and new datasets to address aggregate dynamics, carbon use efficiency of microbial
109 organisms, the role of dissolved organic matter, priming to enhance SOM decomposition, and
110 mineral protection of organic matter.

111 In this article we posit that there is a need and an opportunity for the scientific community
112 to: 1) better identify datasets to characterize ecosystem and landscape properties, processes, and
113 mechanisms that dictate SOC storage and stabilization and their vulnerabilities to change; 2)
114 identify, rescue, and disseminate existing datasets; 3) develop platforms for sharing data, models,
115 and management practices for SOC science; and 4) improve the connection between global C
116 cycle and soil management research communities. The International Soil Carbon Network (ISCN)
117 is a community devoted to open and shared high-quality science for characterizing the state,
118 vulnerabilities, and opportunities for managing SOM. To this end, the ISCN supports cross-
119 disciplinary collaborations that target actionable SOM-related science questions, which are
120 outlined in section 2. Challenges and strategies for the ISCN to function as a community platform
121 for communication, modeling, and data sharing, as well as to increase interoperability among
122 SOM-relevant networks, are outlined in section 3.

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125
 126 Figure 1. Soil organic carbon stocks and areas currently under land use practices. (a) Spatial
 127 variability of soil SOC stocks in the upper meter of soil (where 1 kgC/m² = 10 MgC/ha) , based
 128 on the WISE 3.1. database (Batjes, 2016). (b) Fractional human use of the land surface through
 129 forestry, grazing and agricultural crops (Erb et al., 2007); grey areas represent.... (c) Global SOC
 130 stocks (0-1 m) distributed under different land-use categories. (d) Potential opportunities for soil C
 131 sequestration in presently managed forest, crop, and grazing lands (assuming average management
 132 C gains of 0.04 kg C yr⁻¹ with error bars showing the range of 0.01-0.07 kg C yr⁻¹; Minasny et al.,
 133 2017) could compensate for total emission projections from permafrost-C due to the climate
 134 feedback (Koven et al., 2015; mean and range of projection until 2100 under RCP8.5) and the
 135 projected impact of “human land use”, defined as land use change, agricultural representation,
 136 crop harvest, and management (Pugh et al., 2015; mean and ensemble range of projection until
 137 2100 under RCP8.5). Note that harvest from forestry is not included in this last projection.

138
 139 **2. Challenges for characterizing the state, vulnerabilities, and management opportunities of**
 140 **soil organic matter**

141 Most needed from our community is detection of changes in SOM and SOC, yet such changes
142 vary spatially and temporally because of the many processes that are linked to variations and
143 changes in climate, land use, vegetation, topographic, and geologic factors. Broad-scale ecosystem
144 models generally build upon mechanistic understanding originating from much finer temporal or
145 spatial scales. Upscaling – the scaling or application of knowledge and data from finer to broader
146 scales or from shorter to longer timescales – requires insight, data, and models at various scales,
147 types and complexity because the responses of soil processes to forcing factors are typically
148 different on different spatial scales (O’Rourke et al, 2015). At fine scales, the response might be
149 related to a specific landscape or climate attribute. When aggregating over broad spatial scales,
150 however, information on the relationship between the driver and the response may be lost or
151 obscured. One such example is the apparent dominating influence of temperature, rather than over
152 precipitation, over tropical and global ecosystem carbon fluxes (Cox *et al.*, 2013; Wang *et al.*,
153 2013; Wang *et al.*, 2014). The smaller apparent role of precipitation globally or across the tropics
154 results from large spatial heterogeneity in precipitation. Unusually dry and wet regions cancel each
155 other out when averaged globally, which can obscure an often stronger local/regional precipitation
156 control of ecosystem fluxes (Ahlström *et al.*, 2015; Jung *et al.*, 2017).

157 Long-term changes in SOM_C are particularly difficult to capture with measurements.
158 Fluxes of heterotrophic respiration, for example, can be measured only at fine spatial and temporal
159 scales (Bond-Lamberty et al., 2016) whereas observing short-term changes in SOM_C pools is
160 reduced to detecting small changes relative to a large pool of bulk SOM_C (Stockmann et al.,
161 2013). While radiocarbon measurements suggest that the majority of bulk SOM_C is much older
162 (He et al., 2016), and hence not very active, long-term changes in SOM_C storage could be
163 governed by processes other than those that determine short term fluxes. It is increasing clear that
164 understanding changes and variations in SOM_C requires a robust understanding of processes and
165 mechanisms that underlie stabilization, protection, and destabilization of SOM_C.

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167

168 **2.1 Understanding mechanisms underlying storage and (de)stabilization of SOC**

169 Changes in SOM and SOC are generally based on assessments of stocks and on turnover,
170 residence times, and transit times (Sierra et al. 2017 ; He et al, 2016). Assessments of SOC stocks
171 and transit times remain a critical constraint on the ability of models to predict CO₂ exchanges and
172 their responses to environmental and land use pressures (Todd-Brown et al, 2013). Advancements
173 in measurements and numerical models must be grounded in our best understanding of the
174 processes controlling SOM dynamics across scales (Hinckley et al, 2014).

175 Mechanisms of C stabilization and destabilization (herein referred to as (de)stabilization)
176 are of particular importance for establishing a predictive understanding of SOM dynamics because
177 these same mechanisms presumably drive vulnerabilities (to emission) and opportunities
178 (accumulation or sequestration) under changes in climate, management, or other disturbances.

179 Currently, most global model frameworks rely on state-factor theory (Campbell and
180 Paustian, 2015), where soil properties are the product of a suite of factors such as climate, biota,
181 topography, parent material, and stage or age of pedogenesis (Jenny,1941), superimposed with
182 major land uses such as deforestation or agriculture (Amundson and Jenny, 1991). Under this
183 framework, global-scale spatial heterogeneity of SOC is a direct reflection of variation within
184 these factors and, accordingly, will vary with climate and land use change. A quantitative and
185 predictive understanding of how soil and ecosystem properties interact to regulate SOC remains
186 elusive due to interactions and interdependencies of the state variables and small, local-scale
187 physico-chemical, and biological processes and mechanisms that also influence the
188 (de)stabilization of SOC.

189 A quantitative understanding of SOC pool dynamics requires a quantitative understanding
190 of both processes and mechanisms leading to C (de)stabilization. A process represents a
191 fundamental sequence of actions or steps that lead to a particular outcome, whereas a mechanism
192 reflects the combined interaction of processes (Fig. 2). Processes are often more directly
193 measureable than mechanisms and, therefore, a more fundamental construct for incorporation into
194 models. We tend to classify mechanisms of soil (de)stabilization as being primarily biological,

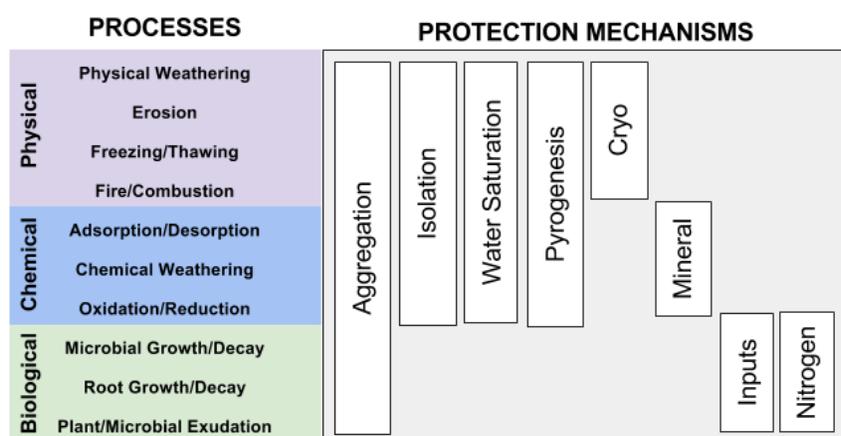
195 physical, or chemical (Six et al., 2002; Fig. 2), but many mechanisms cross these boundaries due
196 to interactions among many processes. The past two decades brought substantial advances in our
197 conceptual understanding of mechanisms of SOM (de)stabilization (Schmidt et al., 2011 and
198 Lehmann and Kleber 2015). Yet, quantitative representations of these concepts in global and
199 regional models lags, due in part to a lack of balance among theory-model-data synthesis and lack
200 of incorporation of local-scale understanding of SOM dynamics.

201 Understanding the mechanisms of SOM (de)stabilization, the underlying driving soil
202 processes, and the relationships between processes and drivers at various spatial scales is needed
203 to evaluate the potential for SOC to change. To address this need, an emerging priority is the
204 execution and synthesis of manipulative field and lab experiments that specifically target
205 processes and drivers at a variety of spatial and temporal scales (see section 2.2. and Fig. 3).

206 Processes can be observed and often measured as rates of change, either as direct flux
207 measurements over short timescales or changes in stocks over moderate to long timescales.
208 However, identifying drivers often requires manipulation or big data synthesis to infer causation
209 Examples include networks of experimental manipulations that target specific processes, such as
210 the Detritus Input and Removal Treatments (DIRT) that manipulates inputs to soil (e.g., Lajtha et
211 al. 2014), the international Soil Experimental Network (iSEN) that warms deep soil (Torn et al.
212 2015), Drought-Net that manipulates precipitation, and temperature and moisture gradient studies
213 (Giardina et al. 2014, reference for moisture). By coupling broadly distributed and comparable
214 data synthesis efforts with process-based models, we have the opportunity to capture mechanistic
215 understanding and to constrain the SOC storage and its sensitivity to disturbance.

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218
 219 Figure 2. Processes controlling SOM_C pools and the mechanisms involved in stabilizing SOC.
 220 Isolation = physical disconnection (e.g. Schimel & Schaeffer 2012); Cryo = cryopreservation;
 221 Pyrogenesis = fire residues; Mineral = mineral interaction; Inputs = microbial and plant residues
 222 that influence desirability or access to microbes (e.g. Kallenbach *et al.* 2016 *Nature*
 223 *Communications*); Nitrogen = nitrogen or other nutrient limitations (e.g. Averill *et al.* 2014
 224 *Nature*)

225

226 2.2. Prioritizing soil data to empower our science

227 There are many types of data, beyond SOC stock data, used to investigate C dynamics at different
 228 spatial and temporal scales (Fig. 3). Data consolidation and archiving efforts so far have focused
 229 principally on SOC stocks (e.g. Batjes *et al.* 2016; Scharlemann *et al.* 2014), but SOC stocks
 230 typically change slowly over timescales of decades to millennia, providing limited sensitivity for
 231 investigating shorter-term processes such as land use and climate impacts (Jastrow *et al.* 2005;
 232 Kravchenko and Robertson 2011; Phillips *et al.* 2015). At the same time, technique advancements
 233 over the last several decades have seen an escalation in methods pertinent for investigating SOC
 234 change at shorter timescales (Fig. 3). For instance, utilization of the enriched atmospheric ¹⁴C
 235 signal (“bomb C”) has allowed tracing and dating of SOC at annual timescales (Trumbore 2000).
 236 Density and size fractionation techniques have helped to distinguish more rapidly cycled SOC
 237 from protected, less rapidly-cycled C (Jastrow 1996, Kong *et al.* 2005, Gregorich and Janzen
 238 1996). More recently, *in situ* chemistry techniques have been used to investigate SOC

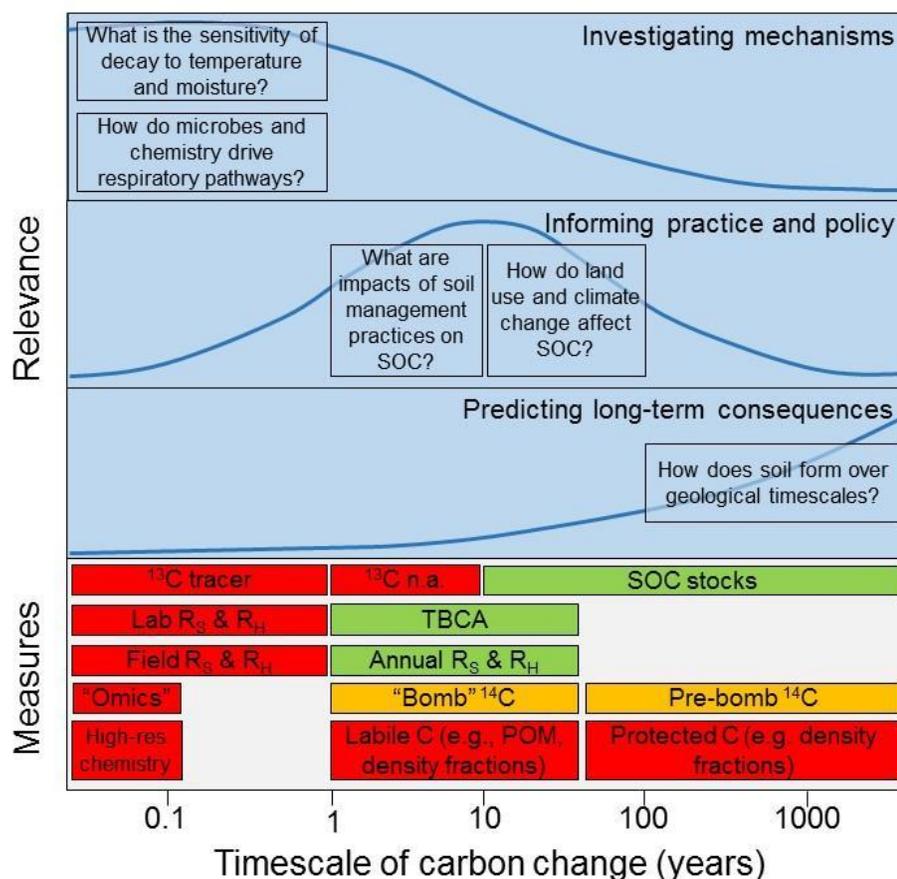
239 transformation over timescales of hours to days (Mackelprang et al. 2016; Haggerty et al. 2014).
240 The data types that are most relevant for measuring SOC change at experimental timescales,
241 however, have not been consolidated and archived, thus impeding two of the more important lines
242 of inquiry in SOC science, namely 1) the biochemical mechanisms of SOC stabilization and
243 destabilization, and 2) the anticipated impacts of climate change and land use (see top panels of
244 Fig. 3).

245 Part of the challenge in assembling and archiving diverse SOC data types is social--they
246 are collected by different sub-communities of soil science and microbiology, and part is logistical--
247 the data have different structures and storage formats (see Supp. Material SM4). Nevertheless,
248 some of these data types have been widely collected, and archiving efforts could open several
249 novel research opportunities. For instance, the soil-to-atmosphere CO₂ flux (soil respiration or R_S)
250 is one data type that has been measured extensively, both in laboratory incubations and *in situ* in
251 field studies, and offers many possibilities for more extensive use. R_S data provides an
252 instantaneous measurement of soil metabolism, and thus offers a unique window into terrestrial
253 carbon dynamics at fine temporal and spatial resolution where questions about temperature,
254 moisture sensitivity and respiratory pathways are addressed (Fig. 3). While a considerable effort
255 has been made to synthesize seasonal and annual averages for field-based R_S fluxes (e.g., Bond-
256 Lamberty and Thomson, 2010a), flux datasets including isotopic measurements (isofluxes), time
257 series and experimental manipulations that include soil moisture, and laboratory-based incubation
258 data have only sparingly archived in centralized repositories (e.g.; Kim et al., 2012). R_S data have
259 been used only sparsely for soil C model validation (Wang et al., 2014) or model benchmarking
260 (Shao et al., 2013) despite having characteristics ideal for these purposes; they reflect fundamental
261 metabolic processes, are geographically widespread, and do not require extensive post-
262 observational processing. High-temporal-resolution R_S data may also present unique possibilities
263 for constraining and validating fluxes inferred from eddy covariance (Phillips et al., 2016) and
264 spatiotemporal analyses (Lavoie et al. 2014; Leon et al. 2014). Finally, because soil-to-atmosphere
265 C fluxes (in particular soil heterotrophic respiration) cannot be directly measured at scales larger

266 than $\sim 1 \text{ m}^2$ (Bond-Lamberty et al., 2016), data compilations have enormous value for upscaling
 267 and for synthesizing our understanding of soil metabolism. While R_s is but one example of data
 268 that will help meet challenges for characterizing SOM and SOC, their relevance to mechanistic
 269 questions of SOC (de)stabilization has the potential to address higher level questions related to
 270 landuse practices, policy, and longterm consequences of change (Fig. 3).

271

272



273 Figure 3. Example research questions and datasets useful for investigating SOC change at different
 274 timescales. Blue lines indicate relevance of the topic and question to the timescale of
 275 measurement. Colors for measures indicate status of data archiving efforts. Measurements can be
 276 well aggregated in centralized repositories (green), have had limited compilation (yellow), or have
 277 had very limited compilation (red). R_s , soil respiration; R_H , heterotrophic respiration; ^{13}C n.a., ^{13}C
 278 natural abundance; TBCA, Total belowground total allocation; POM, particulate organic matter.
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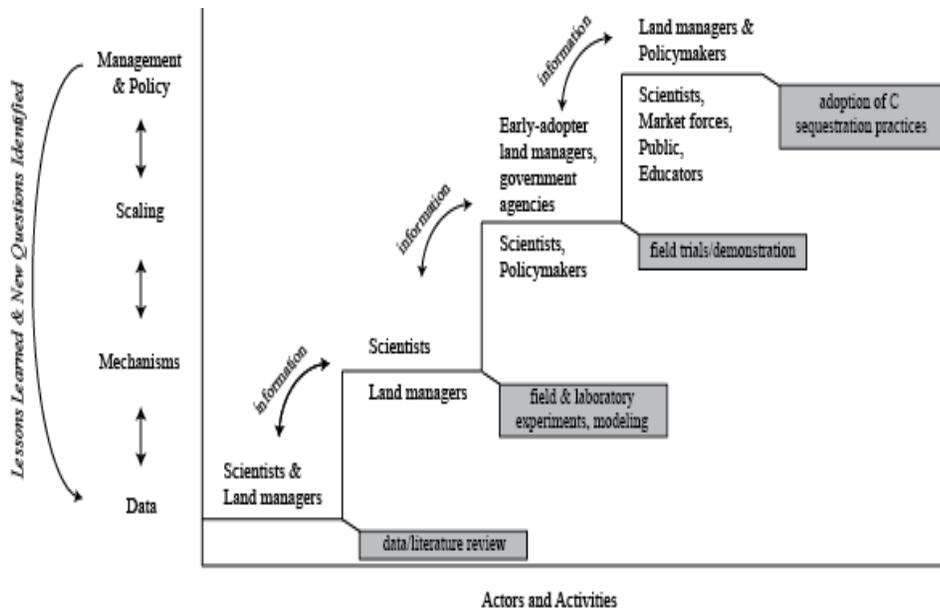
281 2.3 Land management and its potential to increase SOM_C: an emerging priority

282 Increases in SOM_C play a key role in climate regulation through sequestration of CO_2 , but there
 283 also co-benefits relevant to land managers through increased land yield, soil water retention,

284 resilience to extreme weather and nutrient retention. Land managers are primary agents governing
285 changes to SOM and SOM_C stocks, thus in order for scientists to help shape and drive more
286 successful and scalable practices, it is important to view SOC research as a social enterprise as
287 well as a scientific enterprise.

288 Successful management of SOM requires collaboration among scientists, land managers,
289 landowners, and policy makers. A science-land manager-policy partnership can be initiated at any
290 stage of a problem, for example as a science question or a land management challenge. One
291 example (Fig. 4) starts with research question and tethers field/lab experiments to ecological and
292 social issues important to land managers. Seeking feedback from stakeholders at each phase of
293 inquiry also generates new inquiries, which can be visualized in Fig. 4 as movement from right to
294 left on the research-to-policy progression. A cooperative research approach introduces more
295 sources of feedback and points of iteration than an isolated scientific process, but is instrumental
296 for influencing SOM management practices.

297



298
 299 Figure 4. Creating conditions to optimize the effectiveness of land use to sequester SOM_C .
 300 Actors involved in managing lands for soil C sequestration change in response to the scale and
 301 level of information needed. Evaluating and implementing practices (Y axis) starts with scientists
 302 working with land managers and propagates through broader spatial scales and policies as goals
 303 are defined, communicated, and met. Major actors can vary with each step, with activities shown
 304 in the gray boxes. Arrows represent flows of information. In this example, the step-wise progress
 305 from local to more regional scales represent the increasing opportunity to impact both productivity
 306 and CO₂ sequestration through soil C sequestration.

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 309
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 311 **Grazinglands** (rangelands) represent a largely untapped global potential for SOM_C
 312 sequestration as they occur across a wide range of bioclimatic conditions, cover ca. 40% of ice-
 313 free land and store ca. 30% of the terrestrial SOC pool to 1 m depth (Fig. 1) The global potential
 314 for rangeland C sequestration has been estimated to range from 0.3 to as much as 1.6 Pg CO₂-eq
 315 yr⁻¹ (Paustian et al, 2016). Many grazing lands have degraded SOM_C stocks due to historic, poor
 316 management practices and changes in land use intensity. Stocks of SOM_C in grazing lands are
 317 vulnerable to losses through erosion, compaction, and reductions in plant C inputs from plant
 318 community shifts or overgrazing. Improved grazing, irrigation, plant species management, and the
 319 use of organic or inorganic fertilizers of these lands can significantly increase soil C stocks
 320 (Conant et al. 2017). Application of composted organic waste streams has been demonstrated to be
 321 an economic and beneficial proactive that contributes to both rangeland productivity and climate
 322 regulation (Ryals et al 2013, DeLonge et al. 2013; see SM5). Lifecycle assessments, in which

323 broader implications for land management are tracked, (e.g., the waste management and energy
324 systems; DeLonge et al. 2013) and other ecosystem services and values (e.g. biodiversity or
325 endemic plant impacts) are also important issues that drive land management choices.

326 **Forest** SOM management often focuses on minimizing losses to erosion and disturbance
327 and less on building SOM through residue and vegetation management, as is common in
328 grazinglands and croplands (Binkley and Fisher 2013). While there are robust, broadly consistent
329 methods for accounting for and predicting future C stocks in forest aboveground biomass, there is
330 less consensus on methods for assessing belowground SOM and SOM_C. Long-term monitoring
331 (Johnson and Todd 1998; McLaughlin and Phillips 2006), experimental manipulation (Edwards
332 and Ross-Todd 1983; Gundale et al. 2005), expert review (Jandl et al. 2007; Lal 2005),
333 quantitative synthesis (Laganiere et al. 2010; Nave et al. 2010), and ecosystem modeling (Kurz et
334 al. 2009; Scheller et al. 2011) have all produced valuable insights into forest management impacts
335 on SOM. At the same time, the many conflicting results of these studies raise the question of
336 whether responses of SOM to forest management can be generalized across soil and ecosystem
337 types. In addition, the lack of spatially explicit assessments (e.g., maps, geostatistical models) of
338 forest management impacts on SOM highlights our challenge to quantify SOM_C stocks and the
339 complex spatiotemporal processes involved in scaling. Given these limitations, methods of
340 quantifying the spatial distribution and controls on forest SOM across scales are needed for forest
341 practices. These applications may be aided by promising advances in digital soil mapping
342 (Mansuy et al. 2014; Mishra and Riley 2015) and spatially explicit soil carbon assessments
343 (Domke et al. 2017; Soil Survey Staff 2013).

344 **Croplands** have been managed for more than two decades in ways that benefit soil
345 conditions and reduce greenhouse gas emissions (e.g., Smith et al. 2007, Paustian et al.
346 2016). There are many practices influencing SOM_C storage in croplands. These include tillage
347 management (in some cases, Powlson et al. 2014); crop rotations and cover crops (Poeplau and
348 Don 2015); improving crop production through fertilization and irrigation management; selection
349 of high residue-yielding crops; crop intensification by removing bare-fallow management in a

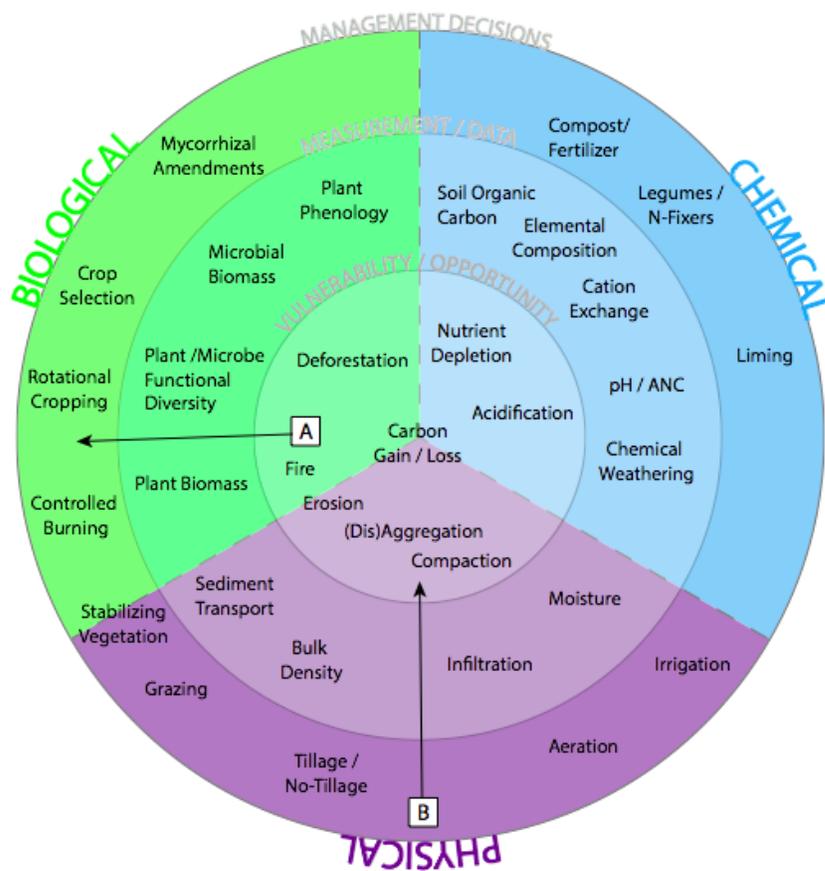
350 cropping system; application of silica residues to reduce greenhouse gas emissions (Gutekunst et
351 al. 2017), and application of organic amendments with manure or biochar.

352 Despite existing knowledge, there is a limited ability to accurately estimate the changes in
353 SOM_C , particularly at smaller scales (Ogle et al. 2010, Paustian et al. 2016) . For example,
354 mechanistic understanding such as the effect of tillage management on aggregate dynamics (Six et
355 al. 2000), has not been effectively incorporated into modelling frameworks. Biochar amendments
356 have emerged as one of the most promising practices for sequestering C in agricultural soils
357 (Lehmann, 2007), but there are still questions about the impact of biochar on SOM dynamics
358 (Knicker, 2011). Efforts to incorporate agricultural SOM_C sequestration into policy programs
359 have been plagued by lack of understanding about the longer term impacts of pervasive warming
360 on SOM_C pools (Conant et al. 2011), which could vary widely depending on the response of
361 microbial communities (Wieder et al. 2015).

362

363 **3. The ISCN as a platform for communication, modeling, and data**

364 While science communities targeting soil health or climate regulation are making great
365 strides in the science of SOM, a combined and coordinated effort could take advantage of
366 technological and communication advances to meet challenges discussed in section 2. The
367 International Soil Carbon Network or ISCN along with partnering entities seeks to establish the
368 basis (platforms) by which we share openly our means of communication, modeling, and data
369 sharing.



370
 371 Fig 5. An approach for applying management options to the science of SOC (de)stabilization.
 372 Three general classes of soil carbon (de)stabilization processes (biological, chemical and physical)
 373 are fundamental to understanding the susceptibility of soils to disturbance (e.g., compaction and
 374 erosion, etc). As such, knowledge of the relevant mechanism at play for a given soil can inform
 375 key measurements needed (e.g., soil infiltration and sediment transport) and effective management
 376 strategies (e.g., diversify vegetation/minimize use and plant stabilizing vegetation/control runoff).
 377
 378 **Communication** of our science starts with restructuring and broadening the soil data that are
 379 shared within and by ISCN, allowing for different types of data, and discovering new ways to
 380 share data without compromising its attribution and credits. To increase the potential impact of
 381 SOM science and to better impact land management practices, it also is beneficial to frame and
 382 disseminate our information in the context of both soil health and climate regulation. For example,
 383 given some knowledge of the dominant processes leading to C stability in a given soil (path A,
 384 Fig. 5), one may evaluate which disturbances may release SOM and SOC and what measurements
 385 would mitigate SOM losses. Conversely, we may apply this framework in the reverse direction.
 386 Given some ongoing or historical management practices (path B, Fig. 5), we can work inward and
 387 to assess what processes could be most affected. Carbon cycle science can also be reframed from

388 the biological, chemical and physical processes paradigm presented in Fig. 2 to a land
 389 management perspective (Table 1). See supplementary materials SM2 for more precise definitions
 390 and references.

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 393 Table 1. Linkages between soil health indicators and SOC. Soil health indicators are readily-
 394 measured soil properties that are used to diagnose the ability of soil to provide services such as
 395 nutrient cycling, erosion mitigation, water storage, or microbial activity. Many of these soil health
 396 indicators relate directly to SOC content, and many can be ameliorated through restorative
 397 practices that increase SOM. For all examples listed, the practices that enhance soil health also
 398 restore (and enhance) SOM and SOC, thus what is good for the goose (soil) is good for the gander
 399 (atmosphere). Based on these example, scientists and land managers can readily agree that
 400 management practices that protect, promote, and conserve soil carbon are practices that prevent
 401 erosion, provide and preserve water and nutrient capacity.

	HEALTH INDICATORS	FUNCTIONAL PROBLEMS	EXPLANATORY C VARIABLES	RESTORATIVE PRACTICES
Physical	Macroaggregate Stability	Erosion, compaction	Root growth, fungal biomass, biological crusts	Conservation tillage, "no-till"
	Water Infiltration Rate	Low infiltration, erosion	SOM content	High residue inputs, cover crops, conservation tillage,
	Available Water Capacity	Arid region water management	SOM content	OM additions
Chemical	Potentially Mineralizable N	Poor fertility	Potentially mineralizable C	Fertility management
	Soil test P	Poor fertility	Applied organic matter	Fertility and pH management
Biological	Microbial Biomass Carbon	Limited soil life	Applied organic matter, root biomass	High residue inputs, cover crops, conservation tillage, "no-till", OM additions

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 408 **Modeling** with computational and conceptual paradigms is an integral part of the scientific
 409 process that greatly enhances our ability to understand variations among spatial and temporal
 410 scales and to precisely and accurately estimate and predict changes in SOC. Conceptual paradigms
 411 that form the scientific basis for our computational models were initially based on "humification"
 412 processes (Hedges, 1988; RothC model (Jenkinson and Rayner, 1977) and Century model (Parton
 413 et al., 1987). The community is increasingly recognizing the role of microbial access to SOC and
 414 its stabilization involving specific mechanisms described in Fig. 2 (Jastrow et al. 1996, Six et al.
 415 2000, Kaiser and Kalbitz 2012, Averill et al. 2014, Keiluweit et al. 2015, Cotrufo et al. 2015,

416 Lehmann and Kleber 2015). Measurements used to drive and test these models vary and are often
417 not structured experimentally to test one model over another. As discussed above in section 2,
418 issues of **spatiotemporal scaling** must address whether mechanisms and functions change or vary
419 between spatial and/or temporal scales. As discussed above, models are increasingly incorporating
420 these new ideas into mathematical frameworks to construct new paradigms (e.g., Allison et al.
421 2010, Wieder et al. 2013, Sulman et al. 2014). Moreover, datasets and databases are needed to
422 evaluate models. The ISCN strives to openly share modeling code and specific parts of models
423 along with data used to drive and test model performance.

424
425 In addition to simply sharing model codes, it is also becoming clear that a **community-based**
426 **model** could emerge from the soils community. In particular modular frameworks with water,
427 temperature, and plant production modules would allow for “plug and play” with new SOM
428 modules that are under development. These other modules would not likely be the focus of
429 development, but are needed to realistically simulate SOM dynamics from experiments and
430 regional analyses. The design of such supporting modules could be informed by or rely on recent
431 progress with frameworks (e.g. PeCAN project <http://pecanproject.github.io/index.html>). As new
432 models are published and shown to work better than the existing SOM community model, the
433 community model would be replaced with improved mathematical frameworks for SOM
434 dynamics. In turn, scientists and investigators evaluating SOM dynamics could incorporate the
435 latest science embodied in the SOM community model housed on the platform into their
436 assessments. ISCN would effectively encourage the use of the latest science in national
437 assessments such as evaluating climate change impacts, greenhouse gas emissions and soil health
438 (e.g. Ogle et al., 2014).

439 As soil or belowground datasets are collected, compiled, and assembled for specific
440 questions or assessments there is an emerging opportunity for **Big soil data** to be designed as a
441 searchable database for soil properties. Empirical models could be structured from a searchable,
442 robust database, but we could also challenge our conceptual and computational models with robust

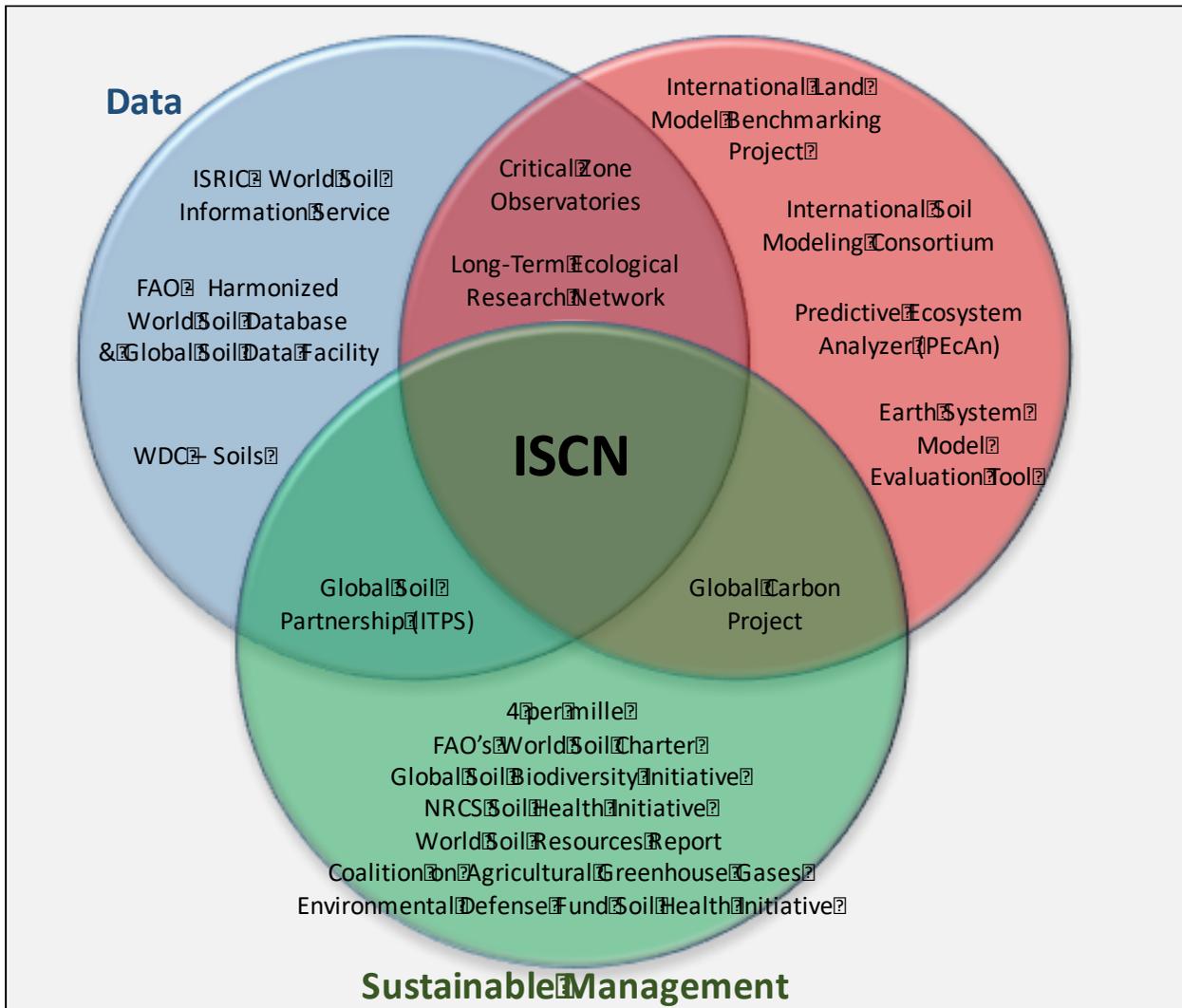
443 data. The ISCN network database (<http://doi.org/10.17040/ISCN/1305039> or <http://ameriflux->
444 data.lbl.gov:8080/ISCN/DOI.html.) afforded early opportunities to design common data templates,
445 promote data synthesis, and generate publications. The ISCN-gen3 database (<http://ameriflux->
446 data.lbl.gov:8080/ISCN/DOI.html) is poised to move beyond observational soil point data and
447 associated drivers, and well into the realm of process-level attributes such as soil fractions and
448 spectral data. These data types have been envisioned and piloted since its earliest generations, but
449 have only recently gained wider attention and use among the wider community of scientists
450 interested in soil carbon.

451 Currently the ISCN database has a mix of overlapping and unique data as compared with
452 other databases (Supplemental Material SM1). For example, most closely aligned are the World
453 Soil Information Service (WoSIS) and ISCN, both of which report soil profile data but for
454 different attributes: The ISCN reports over 100 (carbon plus other attributes) soil
455 properties/attributes for ~70,000 profiles and their constituent layers, whereas WoSIS reports 12
456 properties for over 150,000 profiles. ISCN currently hosts solid phase attributes for soil, and the
457 data are structured in a way compatible with ecosystem CO₂ –land-atmosphere flux data served by
458 FLUXNET and AmeriFlux networks . Despite the large number of soil profiles included in both
459 WoSIS and ISCN, however, there remains an enormous amount of un-archived soil data.
460 Compiling and harmonizing these data could help answer questions of C turnover; soil properties
461 related to mechanisms controlling SOM_C (de)stabilization; soil respiration fluxes in context of
462 soil and environmental measurements; and metrics of pools or forms of bioavailable vs non-
463 available SOM_C.

464 This so-called ‘long tail’ of data has been identified in other fields (Dietze 2013) and
465 represents data that have been collected but, for one reason or another, is not easily available for
466 re-analysis. A comparison of literature and data repository records suggested that process and
467 biological data are underrepresented in repositories, relative to descriptive, chemical and physical
468 data (Suppl. Materials Figure SF1, methods in supplementary materials SM3). Comparison of the
469 top keywords in SOM and SOC literature to data repositories suggested that other data types ripe

470 for synthesis in context of SOM_C include soil incubation and temperature sensitivity, soil
471 chronosequence studies, wildfire emissions/retention, nitrogen and phosphorus cycling, root and
472 fungal dynamics, and soil microbiology. For example, a soil carbon-related data repository search
473 suggests that only 1% of the entries in the broader literature have been archived in data
474 repositories (Suppl. Materials SM3). **Therefore rescuing data from the literature and making them**
475 **accessible and compatible with other contributions and databases is a high priority.**

476 Harmonizing disparate datasets poses unique challenges due to the diversity in types of
477 measurements and their associated methods, unlike larger national and regional survey campaigns
478 which operate under a single protocol. For example, the Biomass And Allometry Database
479 (BAAD) (Falster et al., 2015) has been a unique highly successful example of a community-based
480 data aggregation effort. Public repositories, including Dryad, FigShare, and ORNL DAAC, have
481 emerged and enjoyed enthusiastic support. As these data repositories have grown, issues around
482 discoverability have emerged such as getting people in a common community to agree on a
483 common technical vocabulary has been challenging. Many efforts (e.g. DataONE) have focused
484 on semantics and linked many of these repositories in a unified search framework. Finally, data
485 harmonization is required not only for typical data cleaning operations like correcting unit mis-
486 matches and transparent reproducibility, but also to reconcile different methods and evaluate
487 reliability. This final step requires not only computational skills but also domain expertise.



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 493 Fig 6: Examples of organizations, groups or entities addressing data, modeling and management
 494 relevance of soil carbon. These currently disparate niches need bridging to address complex
 495 problems in soil C science. The soil community is data- and knowledge-ready for a platform like
 496 ISCN that can bridge data, tools, best management practices and outreach. We propose a way
 497 forward to improve soil C data curation with a focus on process variables, which can be applied
 498 into a community model framework and actionable science that harnesses mechanistic
 499 understanding to address questions on soil health management.

500
 501 While these international efforts of the ISCN gain momentum, there are parallel requirements to
 502 coordinate and share technology, data, protocols, and experiences to maximize resources and
 503 generate knowledge. Arguably, this can only be achieved by increasing interoperability within
 504 ISCN and among partner networks, organizations, and members. Interoperability is broadly
 505 defined as the ability of a system to work with or use the parts of another system (Chen et al.

506 2008; Vargas et al. 2017).

507 Challenges related to conceptual barriers include syntactic and semantic differences in the
508 types of information (Madin et al. 2008); technological barriers such as incompatibility of
509 information technologies (e.g. methods to acquire, process, store, exchange, and communicate
510 data; Peters et al. 2014); organizational barriers related to current institutional responsibility and
511 authority such as with institutions, networks, or governments (Supp. Table ST3) ; and cultural
512 barriers that can be country-specific but must be considered to increase interoperability of
513 networks (Vargas et al. 2017).

514 **4.0 Conclusions and Implications**

515 Soils have entered an ‘anthropogenic state’, with most of the global surface area either
516 directly managed by humans or influenced by human activities. As a result, soils globally have
517 lost SOM since at least the Industrial Revolution, with direct impacts on climate, ecosystem
518 productivity and resilience to disturbance. There is a crucial need to improve our science and to
519 communicate our findings . In this paper, we identified the following goals: (1) identify key
520 datasets needed to improve our detection of broad-scale soil C trends and understanding of SOM-
521 C stabilization and destabilization mechanisms, (2) set up infrastructure to rescue, centralize, and
522 disseminate currently disparate soil datasets relevant to critical soil processes, (3) develop a robust
523 and modular modeling platform for developing process-based models that would move field data
524 and localized experiments into a larger Earth systems framework and (4) improve the connection
525 between soil C-cycle science and land management practices. These goals can be achieved as the
526 ISCN improves the exchange of ideas, data, modeling tools, and as we share information and
527 support other networks, organizations, and institutions.

528 Processes that influence changes in SOM and SOC have been defined and quantified over
529 the past several decades, and metrics for soil health, degradation, and storage are beginning reflect
530 the interdisciplinary science needed to link soil/land/ecosystem/crop productivity to CO₂ budgets
531 at various scales. Growing populations, increased land use, and intensified land us compel us to
532 merge the sciences of soil health to those of C cycling. The current state of our soils and

533 opportunities and vulnerabilities that result from different land management practices are of
534 particular importance. In addition, quantifying the optimal SOM_C storage capacity of soils would
535 provide a benchmark to further assess human impact on soils and help quantify future/potential
536 benefits of altered soil management practices. More importantly, these science-based estimates
537 would inform soil valuation by economists, both as resource and service providers, for our
538 societies and ecosystems while also improving assessments of soil C and its exchange with the
539 atmosphere.

540

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550 References

551 Allison, SD, Wallenstein, MD, and Bradford, MA (2010): Soil-carbon response to warming dependent on
552 microbial physiology, *Nat. Geosci*, 3, 336–340.

553

554 Amundson, R, and Jenny, H (1991). A place of humans in the state factor theory of ecosystems and their
555 soils. *Soil Sci.* 151:99–109. doi:10.1097/00010694-199101000-00012

556

557

558 Averill C, Turner BL, Finzi AC. (2014). Mycorrhiza-mediated competition between plants and
559 decomposers drives soil carbon storage. *Nature* 505:543-545.

560

561 Banwart, Steve, et al. "Benefits of soil carbon: report on the outcomes of an international scientific
562 committee on problems of the environment rapid assessment workshop." *Carbon Management* 5.2 (2014):
563 185-192.

564

565 Batjes, NH, Ribeiro, E, van Oostrum, A, Leenaars, J, and Jesus de Mendes, J (2016): Standardised soil
566 profile data for the world (WoSIS, July 2016 snapshot), doi:10.727/isric-wdcsoils.2016003.

567

568 Binkley, D and Fisher, RF (2013): Nutrition Management, in *Ecology and management of forest soils*, pp.
569 254–275, Wiley, Hoboken, NJ

570

571 Bond-Lamberty, B and Thomson, AM (2010a). A global database of soil respiration data, *Biogeosciences*,
572 7, 1915–1926

573

574 Bond-Lamberty, B and Thomson, AM (2010b) Temperature-associated increases in the global soil
575 respiration record, *Nature*, 464(7288), 579–582
576

577 Bond-Lamberty, B, Epron, D, Harden, JW, Harmon, ME, Hoffman, FM, Kumar, J, McGuire, AD and
578 Vargas, R (2016) Estimating heterotrophic respiration at large scales: challenges, approaches, and next
579 steps, *Ecosphere*, 7(6), d01380
580

581 Campbell EE and Paustian K (2015) Current developments in soil organic matter modeling and the
582 expansion of model applications: a review. *Environ. Res. Lett.* 10 123004.
583

584 Chen, D, Doumeingts, G and Vernadat, F (2008) Architectures for enterprise integration and
585 interoperability: Past, present and future. *Computers in Industry* 59:647-659.
586

587

588 Conant, RT (2011) Temperature and soil organic matter decomposition rates – synthesis of current
589 knowledge and a way forward. *Global Change Biology* 17(11): 3392–3404.
590

591 Conant, RT, Cerri, CEP, Osborne, BB and Paustian, K (2017): Grassland management impacts on soil
592 carbon stocks: a new synthesis, *Ecol. Appl.*, 27(2), 662–668, doi:10.1002/eap.1473.

593

594 Cotrufo MF, Soong JL, Horton AJ, Campbell EE, Haddix MH, Wall DL, Parton WJ (2015) Soil organic
595 matter formation from biochemical and physical pathways of litter mass loss. *Nature*
596 *Geosciences*, doi:10.1038/ngeo2520.
597

598 Ciais, P, Gasser, T, Paris, JD, Caldeira, K, Raupach, MR, Canadell, JG, Patwardhan, A, Friedlingstein, P, Piao,
599 SL, and Gitz, V (2013): "Attributing the increase in atmospheric CO₂ to emitters and absorbers." *Nature Climate*
600 *Change* 3, no. 10 926.
601

602 DeLonge, MS, Ryals, R, Silver, WL (2013) A lifecycle model to evaluate carbon sequestration potential
603 and greenhouse gas dynamics of managed grasslands. *Ecosystems*, 16 (2013), pp. 962–979
604

605 Dietze, MC, Serbin, SP, Davidson, C, Desai, AR, Feng, X, Kelly, R, Kooper, R, LeBauer, DS, Mantooth, J,
606 McHenry, K and Wang, D (2014) A quantitative assessment of a terrestrial biosphere model's data needs
607 across North American biomes, *Journal of Geophysical Research-Biogeosciences*, 119(3), 286–300.
608

609 Domke GM., Perry C., Walter B., Nave LE, Woodall CW, Swanston CW (In Press) Toward inventory-
610 based estimates of soil organic carbon in forests of the United States. *Ecological Applications*.
611

612 Edwards, N.T., Ross-Todd, B.M., 1983. Soil carbon dynamics in a mixed deciduous forest following clear
613 cutting with and without residue. *Soil Science Society of America Journal* 47, 1014–1021.
614

615 Erb KH., Gaube V, Krausmann F, Plutzer C, Bondeau A and Haberl H (2007) A comprehensive global 5
616 min resolution land-use data set for the year 2000 consistent with national census data, *Journal of Land Use*
617 *Science*, 2:3, 191-224, DOI: 10.1080/17474230701622981
618

619 Falster, D. S., Duursma, R. A., Ishihara, M. I., Barneche, D. R., FitzJohn, R. G., Vårhammar, A., Aiba, M.,
620 Ando, M., Anten, N., Aspinwall, M. J., Baltzer, J. L., Baraloto, C., Battaglia, M., Battles, J. J., Bond-
621 Lamberty, B., van Breugel, M., Camac, J., Claveau, Y., Coll, L., Dannoura, M., Delagrangé, S., Domec, J.-
622 C., Fatemi, F., Feng, W., Gargaglione, V., Goto, Y., Hagihara, A., Hall, J. S., Hamilton, S., Harja, D.,
623 Hiura, T., Holdaway, R., Hutley, L. S., Ichie, T., Jokela, E. J., Kantola, A., Kelly, J. W. G., Kenzo, T.,
624 King, D., Kloeppel, B. D., Kohyama, T., Komiyama, A., Laclau, J.-P., Lusk, C. H., Maguire, D. A., le

625 Maire, G., Mäkelä, A., Markesteijn, L., Marshall, J., McCulloh, K., Miyata, I., Mokany, K., Mori, S.,
626 Myster, R. W., Nagano, M., Naidu, S. L., Nouvellon, Y., O’Grady, A. P., O’Hara, K. L., Ohtsuka, T.,
627 Osada, N., Osunkoya, O. O., Peri, P. L., Petritan, A. M., Poorter, L., Portsmouth, A., Potvin, C., Ransijn, J.,
628 Reid, D., Ribeiro, S. C., Roberts, S. D., Rodríguez, R., Saldaña-Acosta, A., Santa-Regina, I., Sasa, K.,
629 Selaya, N. G., Sillett, S. C., Sterck, F., Takagi, K., Tange, T., Tanouchi, H., Tissue, D., Umehara, T.,
630 Utsugi, H., Vadeboncoeur, M. A., Valladares, F., Vanninen, P., Wang, J. R., Wenk, E., Williams, R., de
631 Aquino Ximenes, F., Yamaba, A., Yamada, T., Yamakura, T., Yanai, R. D. and York, R. A (2015). BAAD:
632 a Biomass And Allometry Database for woody plants, *Ecology*, 96(5), 1445–1445
633

634 Gundale, MJ, DeLuca, TH, Fiedler, CE, Ramsey, PW, Harrington, MG, Gannon, JE (2005) Restoration
635 treatments in a Montana ponderosa pine forest: effects on soil physical, chemical and biological properties.
636 *Forest Ecology and Management* 213, 25–38.
637

638 Gutekunst, MY, Vargas, R., and Seyfferth ,AL (2017) Impacts of soil incorporation of pre-incubated silica-
639 rich rice residue on soil biogeochemistry and greenhouse gas fluxes under flooding and drying. *Science of*
640 *the Total Environment*:134-143.
641

642 Giardina, CP , Litton, CM , Crow, SE, Asner, GP (2014) Warming-related increases in soil CO2 efflux are
643 explained by increased below-ground carbon flux. *Nat. Clim. Change*, pp. 1758–6798.
644

645 Gibbs, HK and Salmon, JM (2015) Mapping the world’s degraded lands. *Applied Geography* 57, 12–21.
646 doi:10.1016/j.apgeog.2014.11.024
647

648 Gregorich, EG and Janzen, HH (1996) Storage of Soil Carbon in the Light Fraction and Macroorganic
649 Matter, in *Structure and organic matter storage in agricultural soils*, edited by M. R. Carter and B. A.
650 Stewart, Lewis Publishers, Boca Raton, FL.
651

652 Hagerty, SB, van Groenigen, KJ, Allison, SD, Hungate, BA, Schwartz, E, Koch, GW, Kolka, RK and
653 Dijkstra, P(2014) Accelerated microbial turnover but constant growth efficiency with warming in soil, *Nat.*
654 *Clim. Change*, 4(10), 903–906, doi:10.1038/nclimate2361.
655

656 He, Y, Trumbore, SE, Torn, MS, Harden, JW, Vaughn, LJS, Allison, SD, and Randerson, JT
657 (2016)Radiocarbon constraints imply reduced carbon uptake by soils during the 21st century, *Science*, 353.
658 1419-1424.
659

660 Hedges, JI (1988) in *Humic Substances and their Role in the Environment* (eds Frimmel, F. H. &
661 Christman, R. F.) 45–58 (John Wiley & Sons)
662

663 Hinckley, E-L S, Wieder, W, Fierer, N and Paul, E (2014) Digging Into the World Beneath Our Feet:
664 Bridging Across Scales in the Age of Global Change, *Eos Trans. Am. Geophys. Union*, 95(11), 96–97,
665 doi:10.1002/2014EO110004.
666

666 Hugelius, G, Strauss, ., Zubrzycki, S, Harden, JW, Schuur, EAG, Ping, C-, Schirrmeister,L, Grosse, G,
667 Michaelson, GJ, Koven, C, O’Donnell, J, Elberling, B, Mishra, U, Camill,P, Yu, Z, Palmtag, J, Kuhry, P
668 (2014) Estimated stocks of circumpolar permafrost carbon with quantified uncertainty ranges and identified
669 data gaps, *Biogeosciences*, Volume 11, pages 6573-6593
670

671 Jackson, RB, Lajtha, K, Crow, SE, Hugelius, G, Kramer, MG, and Piñeiro G (2017) The ecology of soil
672 carbon: pools, vulnerabilities, and biotic and abiotic controls. *Annual Review of Ecology, Evolution, and*
673 *Systematics*, in press.
674

675 Jandl R, Lindner M, Vesterdal L, Bauwensd B, Baritze R, Hagedorn F, Johnson D, Minkkinen K, Byrne K.

676 (2007) How strongly can forest management influence soil carbon sequestration? *Geoderma* 137: 253-268.
677

678 Jastrow, JD (1996) Soil aggregate formation and the accrual of particulate and mineral-associated organic
679 matter, *Soil Biol. Biochem.*, 28(4–5), 665–676, doi:10.1016/0038-0717(95)00159-X.

680

681 Jastrow, JD, Michael Miller, R, Matamala, R, Norby, RJ, Boutton, TW, Rice, CW, and Owensby, CE
682 (2005.) Elevated atmospheric carbon dioxide increases soil carbon, *Glob. Change Biol.*, 11(12), 2057–2064,
683 doi:10.1111/j.1365-2486.2005.01077.x

684 Jenny, Hans 1941. *Factors of Soil Formation*.

685

686 Jenkinson, DS and Rayner, JH: (1977) The Turnover of Soil Organic Matter in Some of the Rothamsted
687 Classical Experiments, *Soil Sci.*, 123(5)

688

689 Johnson, D, Todd, DE (1998) The effects of harvesting on long-term changes in nutrient pools in a mixed
690 oak forest. *Soil Science Society of America Journal* 62, 1725–1735.

691

692

693 Kaiser, K and Kalbitz, K (2012) Cycling downwards – dissolved organic matter in soils, *Soil Biol.*
694 *Biochem.*, 52, 29–32, doi:10.1016/j.soilbio.2012.04.002.

695

696 Kallenbach, CM, Frey, SD, and Grandy, AS (2016) Direct evidence for microbial-derived soil organic
697 matter formation and its ecophysiological controls, *Nature Comm.*, 7, 13630.

698

699 Keiluweit M, Bougoure JJ, Nico PS, Pett-Ridge J, Weber PK and Kleber M (2015) Mineral protection of
700 soil carbon counteracted by root exudates. *Nature Climate Change*, 1–8.

701

702 Kravchenko, AN and Robertson, GP (2011) Whole-Profile Soil Carbon Stocks: The Danger of Assuming
703 Too Much from Analyses of Too Little, *Soil Sci. Soc. Am. J.*, 75(1), 235, doi:10.2136/sssaj2010.0076

704

705 Kim, D-G., Vargas, R, Bond-Lamberty, B and Turetsky, MR (2012) Effect of rewetting and thawing on
706 soil gas fluxes: a review of current literature and suggestions for future research, *Biogeosciences*, 9, 2459–
707 2483

708

709 Knicker, H (2011) Pyrogenic organic matter in soil: Its origin and occurrence, its chemistry and survival in
710 soil environments. *Quaternary International* 243:251-263

711

712 Köchy, M, Hiederer, R, & Freibauer, A (2015) Global distribution of soil organic carbon–Part 1: Masses
713 and frequency distributions of SOC stocks for the tropics, permafrost regions, wetlands, and the world.
714 *Soil*, 1(1), 351-365.

715

716 Kong, AY, Six, J, Bryant, DC, Denison, RF and van Kessel, C (2005) The Relationship between Carbon
717 Input, Aggregation, and Soil Organic Carbon Stabilization in Sustainable Cropping Systems, *Soil Sci. Soc.*
718 *Am. J.*, 69(4), 1078, doi:10.2136/sssaj2004.0215.

719

720 Koven CD, Schuur EAG, Schädel C, Bohn TJ, Burke EJ, Chen G, Chen X, Ciais P, Grosse G, Harden JW,
721 Hayes DJ, Hugelius G, Jafarov EE, Krinner G, Kuhry P, Lawrence DM, Macdougall AH, Marchenko SS,
722 McGuire AD, Natali SM, Nicolsky DJ, Olefeldt D, Peng S, Romanovsky VE, Schaefer KM, Strauss J, Treat
723 CC, Turetsky M. (2015) A simplified, data-constrained approach to estimate the permafrost carbon–climate
724 feedback. *Philosophical Transactions of the Royal Society of London A: Mathematical, Physical and*

725 Engineering Sciences, 373, DOI: 10.1098/rsta.2014.0423
726
727 Kravchenko, AN and Robertson, GP (2011) Whole-Profile Soil Carbon Stocks: The Danger of Assuming
728 Too Much from Analyses of Too Little, *Soil Sci. Soc. Am. J.*, 75(1), 235, doi:10.2136/sssaj2010.0076,
729
730 Kurz, WA, Dymond, CC, White, TM, Stinson, G, Shaw, CH, Rampley, GJ, Smyth, C, Simpson, BN,
731 Neilson, ET, Trofymow, JA, Metsaranta, J, Apps, MJ. (2009) CBM-CFS3: A model of carbon-dynamics in
732 forestry and land-use change implementing IPCC standards. *Ecological Modelling* 220: 480-504.
733
734 Laganriere J, Angers DA, Pare D. (2010) Carbon accumulation in agricultural soils after afforestation: a
735 meta-analysis. *Global Change Biology* 16: 439-453.
736
737 Lajtha, K, Bowden, RD and Nadelhoffer, K (2014) Litter and Root Manipulations Provide Insights into
738 Soil Organic Matter Dynamics and Stability, *Soil Sci. Soc. Am. J.*, 78(S1), S261,
739 doi:10.2136/sssaj2013.08.0370nafsc
740
741 Lal R (2005) Forest soils and carbon sequestration. *Forest Ecology and Management* 220: 242-258.
742
743 Lavoie, M, Phillips, CL, and Risk, D (2014) A practical approach for uncertainty quantification of high
744 frequency soil respiration using Forced Diffusion chambers, *J. Geophys. Res.-Biogeosciences*, 120,
745 doi:10.1002/2014JG002773, 2014.
746
747 Leon, E, Vargas, R, Bullock, S, Lopez, E, Panosso, AR, and La Scala, N (2014) Hot spots, hot moments,
748 and spatio-temporal controls on soil CO₂ efflux in a water-limited ecosystem, *Soil Biol. Biochem.*,
749 doi:10.1016/j.soilbio.2014.05.029
750
751 Lehmann, J (2007) A handful of carbon. *Nature* 447(7141): 143-144.
752
753 Lehmann, J and Kleber, M (2015) The contentious nature of soil organic matter, *Nature* 528, 60-68,
754 doi:10.1038/nature16069.
755
756 Loisel, Julie , van Bellen, Simon , Pelletier, Luc , Talbot, Julie, Hugelius, Gustaf , Karran, Daniel, Yu,
757 Zicheng, Nichols, Jonathan, Holmquist, James (2017), Insights and issues with estimating northern peatland
758 carbon stocks and fluxes since the Last Glacial Maximum. *Earth Science Reviews*, volume 165, pp 59-80.
759 <http://doi.org/10.1016/j.earscirev.2016.12.001>
760
761 Mackelprang R, Saleska SR, Jacobsen CS, Jansson JK, Taş N (2016) Permafrost Meta-Omics and
762 Climate Change. *Annual Review of Earth and Planetary Sciences*, **44**, 439–462.
763
764 Madin, JS., Bowers, S, Schildhauer, MP, and Jones, MB (2008) Advancing ecological research with
765 ontologies. *Trends in Ecology & Evolution* 23:159-168.
766
767 Mansuy N, Thiffault E, Pare D, Bernier P, Guindon L, Villemaire P, Poirier V, Beaudoin A (2014) Digital
768 mapping of soil properties in Canadian managed forests at 250 m of resolution using the k-nearest neighbor
769 method. *Geoderma* 235:59-73.
770
771 McLaughlin, JW, Phillips, SA (2006). Soil carbon, nitrogen, and base cation cycling 17 years after whole-
772 tree harvesting in a low-elevation red spruce (*Picea rubens*)-balsam fir (*Abies balsamea*) forested watershed
773 in central Maine, USA. *Forest Ecology and Management* 222, 234–253.
774
775 Minasny, B, Malone, BP, McBratney, AB, Angers, DA, Arrouays, D, Chambers, A, Chaplot, V, Chen, ZS,

775 Cheng, K, Das, BS and Field, DJ (2017) Soil carbon 4 per mille. *Geoderma*, 292, pp.59-86.
776

777 Mishra U, Riley WJ. (2015) Scaling impacts on environmental controls and spatial heterogeneity of soil
778 organic carbon stocks. *Biogeosciences* 12: 3993-4004.
779

780 Nave L.E., Vance E.D., Swanston C.W., Curtis P.S. (2010) Harvest impacts on soil carbon storage in
781 temperate forests. *Forest Ecology and Management* 259: 857-866.
782

783 Ogle, S.M., F.J. Breidt, M. Easter, S. Williams, K. Killian, and K. Paustian. (2010) Scale and uncertainty in
784 modeled soil organic carbon stock changes for US croplands using a process-based model. *Global Change*
785 *Biology* 16:810-820.
786

787 Ogle, S.M., L. Olander, L. Wollenberg, T. Rosenstock, F. Tubiello, K. Paustian, L. Buendia, A. Nihart, and
788 P. Smith. (2014) Reducing agricultural greenhouse gas emissions in developing countries: providing the
789 basis for action. *Global Change Biology* 20:1-6
790

791 O'Rourke SM, Angers DA, Holden NM, McBratney AB. (2015) Soil organic carbon across scales. *Global*
792 *Change Biology* 21: 3561–3574.
793

794 Parton W. J., Schimel D. S., Cole C. V. and Ojima D. S. 1987. Analysis of factors controlling soil organic-
795 matter levels in Great-Plains grasslands. **51**, 1173–1179.
796

797

798 Paustian, K., J. Lehmann, S. Ogle, D. Reay, G.P. Robertson, P. Smith. (2016) Climate smart soils. *Nature*:
799 532:49-57.
800

801 Peters, D. P. C., H. W. Loescher, M. D. SanClements, and K. M. Havstad. (2014) Taking the pulse of a
802 continent: expanding site-based research infrastructure for regional- to continental-scale ecology. *Ecosphere* **5**:1-
803 23.
804

805 Phillips, C. L., Murphey, V., Lajtha, K. and Gregg, J. W. (2016) Asymmetric and symmetric warming
806 increases turnover of litter and unprotected soil C in grassland mesocosms. *Biogeochemistry*, 128(1–2),
807 217–231, doi:10.1007/s10533-016-0204-x

808

809 Phillips, C. L., Bond-Lamberty, B., Desai, A. R., Lavoie, M., Risk, D., Tang, J., Todd-Brown, K. E. O. and
810 Vargas, R., (2016) The value of soil respiration measurements for interpreting and modeling terrestrial
811 carbon cycling, *Plant Soil*, in press, doi:10.1007/s11104-016-3084-x
812

813 Poepflau, C., Don, A.(2015) Carbon sequestration in agricultural soils via cultivation of cover crops – A
814 meta-analysis. *Agriculture, Ecosystems & Environment* 200, 33–41. doi:10.1016/j.agee.2014.10.024
815

816 Powlson, D.S., Stirling, C.M., Jat, M.L., Gerard, B.G., Palm, C.A., Sanchez, P.A., Cassman, K.G.
817 (2014)Limited potential of no-till agriculture for climate change mitigation. *Nature Climate Change* 4, 678–
818 683. doi:10.1038/nclimate2292
819

820 Pugh ,T.A.M., Arneth, A., Olin, S., Ahlström, A., Bayer, A.D., Klein, Goldewijk K., Lindeskog, M. and
821 Schurgers, G. (2015) Simulated carbon emissions from land-use change are substantially enhanced by
822 accounting for agricultural management, *Environmental Research Letters*, 10, 124008, doi:10.1088/1748-
823 9326/10/12/124008
824

825 Ryals, R., W.L. Silver. (2013) Effects of organic matter amendments on net primary productivity and
826 greenhouse gas emissions in annual grasslands, *Ecological Applications*, 23(1), 46-59.
827

828 Scharlemann, J. P., Tanner, E. V., Hiederer, R., & Kapos, V. (2014) Global soil carbon: understanding and
829 managing the largest terrestrial carbon pool. *Carbon Management*, 5(1), 81-91.
830

831 Scheller RM, Hua D, Bolstad PV, Birdsey RA, Mladenoff DJ. 2011. The effects of forest harvest intensity
832 in combination with wind disturbance on carbon dynamics in Lake States Mesic Forests. *ECOLOGICAL*
833 *MODELLING* 222: 144-153.
834

835 Schimel J. and Schaeffer S. M. 2012. Microbial control over carbon cycling in soil. *Front. Microbio.* **3**, 1–
836 11. Available at: <http://journal.frontiersin.org/Journal/10.3389/fmicb.2012.00348/full>.
837

838

839 Schmidt, M. W., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., et al. (2011). Persistence of soil
840 organic matter as an ecosystem property. *Nature*, 478(7367), 49-56.
841

842 Shao, P., Zeng, X., Moore, D. J. P. and Zeng, X. (2013) Soil microbial respiration from observations and
843 Earth System Models, *Environ. Res. Lett.*, 8(3), 034034.
844

845 Sierra, C. A., Müller, M., Metzler, H., Manzoni, S. and Trumbore, S. E. (2017) The muddle of ages,
846 turnover, transit, and residence times in the carbon cycle, *Glob. Change Biol.*, 23(5), 1763–1773,
847 doi:10.1111/gcb.13556

848 Six J., Elliott E. T. and Paustian K. (2000) Soil macroaggregate turnover and microaggregate formation: a
849 mechanism for C sequestration under no-tillage agriculture. *Soil Biol Biochem* **32**, 2099–2103.
850

851 Six, J., Conant, R.T., Paul, E.A., and Paustian, K. 2002. Stabilization mechanisms of soil organic matter:
852 implications for C-saturation of soils, *Plant and Soil*, 241, 155-176
853

854 Smith, P. Martino, D., Cai, Z., Gwary, D., Janzen, H., Pushpam, K., McCari, B., Ogle, S., O'Mara, F.,
855 Rice, C., Scholes, B., Sirotenko, O., Howden, M., McAllister, T., Pan, G., Romaenkov, V., Schneider, W.,
856 and Towparyoon, (2007) S. Policy and technological constraints to implementation of greenhouse gas
857 mitigation options in agriculture, - *Agriculture, Ecosystems & Environment*, V. 18, 1, 6-28.
858

859 Smith, P., D. Martino, Z. Cai, D. Gwary, H. Janzen, P. Kumar, B. McCarl, S. Ogle, F. O'Mara, C. Rice, B.
860 Scholes, O. Sirotenko, M. Howden, T. McAllister, G. Pan, V. Romanenkov, U. Schneider, S. Towprayoon,
861 M. Wattenbach, and J. Smith. (2007) Greenhouse gas mitigation in agriculture. *Phil. Transactions of Royal*
862 *Society B*, DOI: 10.1098/rstb.2007.2184.
863

864 Soil Survey Staff (2013) Rapid Carbon Assessment (RaCA) project. United States Department of
865 Agriculture, Natural Resources Conservation Service. Available online. June 1, 2013 (FY2013 official
866 release).
867

868 Stockmann, U., Adams, M. A., Crawford, J. W., Field, D. J., Henakaarchchi, N., Jenkins, M., Minasny, B.,
869 McBratney, A. B., Courcelles, V. de R. de, Singh, K., Wheeler, I., Abbott, L., Angers, D. A., Baldock, J.,
870 Bird, M., Brookes, P. C., Chenu, C., Jastrow, J. D., Lal, R., Lehmann, J., O'Donnell, A. G., Parton, W. J.,
871 Whitehead, D. and Zimmermann, M (2013) The knowns, known unknowns and unknowns of sequestration
872 of soil organic carbon, *Agric. Ecosyst. Environ.*, 164, 80–99, doi:10.1016/j.agee.2012.10.001
873

874 Sulman, B. N., Phillips, R. P., Oishi, A. C., Shevliakova, E. and Pacala, S. W.: Microbe-driven turnover

875 offsets mineral-mediated storage of soil carbon under elevated CO₂ (2014) *Nat. Clim. Change*, 4(12),
876 1099–1102, doi:10.1038/nclimate2436

877 Todd-Brown, K.E., Randerson, J.T., Post, W.M., Hoffman, F.M., Tarnocai, C., Schuur, E.A., Allison, S.D.
878 (2013) Causes of variation in soil carbon simulations from CMIP5 Earth system models and comparison
879 with observations. *Biogeosciences* 10 (3).
880

881 Torn, M. S., Chabbi, A., Crill, P., Hanson, P. J., Janssens, I. A., Luo, Y., Pries, C. H., Rumpel, C., Schmidt,
882 M. W. I., Six, J., Schrumpf, M. and Zhu, B. (2015) A call for international soil experiment networks for
883 studying, predicting, and managing global change impacts, *SOIL*, 1(2), 575–582, doi:10.5194/soil-1-575-
884 2015.

885 Trumbore, S.E. (2000) Age of soil organic matter and soil respiration: radiocarbon constraints on
886 belowground C dynamics, *Ecol. Appl.*, 10, 399–411
887

888 Urban, D.L., O’Neill, R.V. & Shugart, H.H. (1987) Landscape ecology: A hierarchical perspective can help
889 scientists understand spatial patterns. *Bioscience*, 37, 119–127
890

891 Vargas, R., D. Alcaraz-Segura, R. Birdsey, N. A. Brunsell, C. O. Cruz-Gaistardo, B. de Jong, J. Etchevers,
892 M. Guevara, D. J. Hayes, K. Johnson, H. W. Loescher, F. Paz, Y. Ryu, Z. Sanchez-Mejia, and K. P.
893 Toledo-Gutierrez. (2017) Enhancing interoperability to facilitate implementation of REDD+: case study of
894 Mexico. *Carbon Management*:1-9. doi: 10.1080/17583004.2017.1285177
895

896 Wang, X., Liu, L., Piao, S., Janssens, I. A., Tang, J., Liu, W., Chi, Y., Wang, J. and Xu, S. (2014) Soil
897 respiration under climate warming: differential response of heterotrophic and autotrophic respiration, *Glob.*
898 *Chang. Biol.*, 20(10), 3229–3237.
899

900 Wieder, W.R., Cleveland, C.C., Smith, W.K., Todd-Brown, K. (2015) Future productivity and carbon
901 storage limited by terrestrial nutrient availability. *Nature Geoscience* doi: 10.1038/ngeo2413
902
903