

**RADIOCARBON AND VARVE CHRONOLOGIES OF ANNUALLY  
LAMINATED LAKE SEDIMENTS OF GOŚCIAŻ LAKE,  
CENTRAL POLAND**

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**ABSTRACT.** A sequence of annually laminated sediments of the Gościąg Lake spans ca 13,000 yr and is actually the longest known continuous sequence in the world.  $^{14}\text{C}$  age measurements were performed on organic and carbonate fractions of bulk samples of laminated sediments from core G0. Accurate measurements of varve thickness performed on the lower part of cores G1 and G2 were used to establish a floating varve chronology covering ca 10,000 yr. Matching of cores G0, G1 and G2 permits comparison of  $^{14}\text{C}$  dates with varve chronology. Good agreement of calibrated  $^{14}\text{C}$  dates with the varve time scale suggests annual lamination of the sediment. Analysis of periodicities in varve thickness indicates solar 11- and 22-yr cycles, as well as a 200-yr cycle over a good part of the investigated sequence. Results of  $^{14}\text{C}$  measurements of carbonate fractions are used to study changes in the water depth of the lake during its history.

INTRODUCTION

In recent years there has been an increasing interest in annually laminated lake sediments from various parts of the world. As discussed in comprehensive reviews by O'Sullivan (1983) and Saarnisto (1986), such sediments may be used for studying various aspects of the Quaternary, including geochronology, climatic changes and the history of vegetation. The importance of laminated lake sediments and glacial varves for studies of natural  $^{14}\text{C}$  changes was realized ca 20 years ago, and the early results of Stuiver (1970) and Tauber (1970) were recently reassessed in light of new calibration data (Stuiver *et al.*, 1986). Important results on  $^{14}\text{C}$  and varve chronology were also obtained from studies of laminated lake sediments of, eg, Schleinsee, Germany (Geyh, Merkt & Müller, 1971) and Lake Lampelönjärvi, South Finland (Tolonen, 1980).

The sequence of laminated sediments of Lake Gościąg, central Poland, spanning >12,500 yr (Ralska-Jasiewiczowa, Wicik & Więckowski, 1987) is actually the longest known in the world (*cf.* Saarnisto, 1986, Table 17.1, p 344). It presents a unique opportunity to study environmental change during a significant part of the Late Glacial and Holocene, with potential significance for studying natural  $^{14}\text{C}$  variations and calibration of the  $^{14}\text{C}$  time scale beyond 10,000 BP, using AMS-dated terrestrial macrofossils. This paper presents a short summary of results of chronologic studies on cores collected in 1985 and 1987.

RADIOCARBON AND VARVE CHRONOLOGY

*Cores*

The chronologic studies are based on four cores (Fig 1), taken in the deepest part of the lake at water depth, ca 26m. Core G0 was taken in 1985 in 1m segments; cores G1, G2 and G3 were taken in 1987 in 2m segments

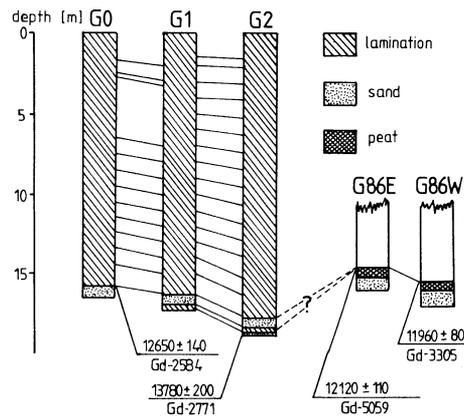


Fig 1. Correlation of cores (G0, G1 and G2 with G86E and G86W)

in approximately the same place as core G0. The distance between cores G1, G2 and G3 was ca 5m. All the cores have reached the sandy layer below the laminated sediment. In cores G1 and G2, a short laminated sequence (ca 30cm thick) was found below a sandy layer, ca 50cm thick. Corings in other parts of the lake provided two cores with thin peat layers below the lake marl; unfortunately, lamination in these cores was not visible.

#### $^{14}\text{C}$ dating

$^{14}\text{C}$  age determinations were performed on carbonate and organic fractions of bulk samples from core G0. The results are shown in Figure 2 as functions of the number of couplets in this core, estimated by K Więckowski (*cf* Pazdur *et al*, 1987). Three characteristic features should be noted: 1)  $^{14}\text{C}$  dates of both organic and carbonate fractions are greater than the corresponding number of couplets, N; 2)  $^{14}\text{C}$  dates of carbonate fraction are older than those obtained on organic matter; 3) the difference between  $^{14}\text{C}$  dates and the number of couplets decreases with increasing age. According to the  $\delta^{13}\text{C}$  values, shown in the lower part of Figure 2, the core may be regarded as isotopically bipartite. These observations lead to the conclusion that  $^{14}\text{C}$  dates of organic matter, which show higher regularity, can be corrected for initial apparent age in the lower and upper parts of the core. Corrected conventional  $^{14}\text{C}$  dates of organic fractions in the uppermost part of the core, obtained by subtracting a constant correction of 1600 yr, were then calibrated according to Stuiver and Pearson (1986), Pearson and Stuiver (1986) and Pearson *et al* (1986). A similar procedure was applied to dates from the lowermost part of the core, *ie*, subtracting a constant value of 1400 yr and using the calibration curve of Stuiver *et al* (1986). Age correction of 1600 yr was determined by assuming zero age of the youngest sample; the value of 1400 yr was determined by trial and error. The calibrated ages are shown as solid squares in Figure 2. The least squares line (Fig 2) has the equation:

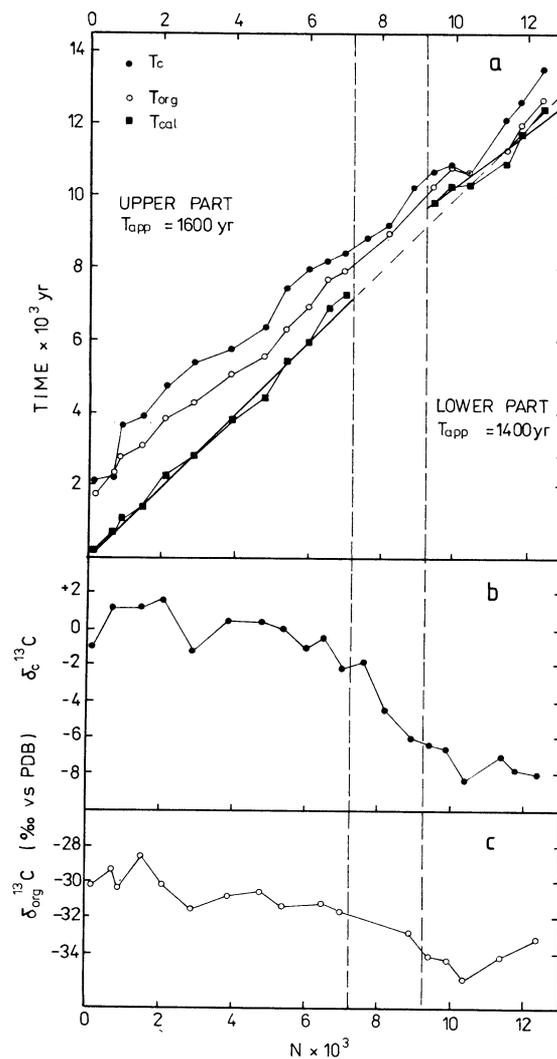


Fig 2.  $^{14}C$  dates and  $\delta^{13}C$  values obtained on core G0 in function of varve number acc to K Więckowski (Ralska-Jasiewiczowa, Wicik & Więckowski, 1987)

$$T_{cal} = (-8 \pm 110) + (1.00 \pm 0.02)N.$$

Correlation coefficient  $r=1.00$ , the scatter of individual data points with respect to the least squares line is equal to  $s=260$  yr, *ie*, slightly exceeds the dating error.

### *Varve chronology*

A floating varve chronology, covering ca 10,000 yr, was established from cores G1 and G2. The surface of both cores was thoroughly cleaned and photographed; measurements of thickness of individual light (summer) and dark (winter) layers were performed on negatives using the apparatus designed for dendrochronological studies (Goslar, 1987). Apparatus accuracy is equal to 0.01mm; however, the boundaries between light and dark layers were identified with greater uncertainty, estimated to ca 0.05mm, and the resulting accuracy of layer thickness measurements is estimated to be ca 10%.

Partial sequences obtained on cores G1 and G2 were combined into a floating varve chronology consisting of four segments which jointly cover 9682 couplets. The thicknesses of the light and dark layers were measured and treated separately, and the resulting chronology was constructed in three sequences, corresponding to light, dark, and total annual (light+dark) increments. There are three gaps in this sequence: the oldest gap is caused by a sandy layer, ca 0.5m thick, overlying the oldest sequence with 294 couplets, two other gaps are represented by blue-gray massive layers (ca 5 and 15mm thick), which occur in both cores. There are also several other gaps in the sequence obtained on cores G1 and G2; these gaps were filled with data available from core G0.

### *Correlation of <sup>14</sup>C and varve chronologies*

Direct comparison of <sup>14</sup>C and varve chronologies of sediments of Gościąg Lake is complicated by the reservoir effect and resulting aging of <sup>14</sup>C dates of core G0. However, it is possible to gather the evidence provided by different cores shown in Figure 1 into a consistent chronologic picture: 1) <sup>14</sup>C dates were obtained on core G0, where two parts with approximately constant  $\delta^{13}\text{C}$  values may be distinguished; 2) peat layers found at the base of cores G86E and G86W provide the same dates (within limits of error), which may be used as estimates of the beginning of lake marl sedimentation; 3) frozen cores taken by M Saarnisto from the uppermost unconsolidated sediment show lamination until the present; 4) available photographic documentation of core G0 permits exact matching of this core with cores G1 and G2, in spite of errors in varve number in the middle part of the core (*cf.* Goslar, Pazdur & Walanus, 1988).

### PERIODICITIES IN THE VARVE SEQUENCE

Periodicities in the sequence of varve thickness obtained on cores G1 and G2 were sought for using running phase analysis, developed by Walanus (1988a). This method is based on Fourier analysis performed in a running manner, strictly for defined values of period, equal to 11, 22, 35 and 200 yr, as suggested by results of other studies. The results indicate superimposed 11- and 22-yr cycles in the entire analyzed sequence, although, in some intervals, they cannot be regarded as significant, as the amplitude is rather low and the phase varies randomly. The 35- and 200-yr cycles occur over large parts of the laminated sequence, the last cycle is very distinct

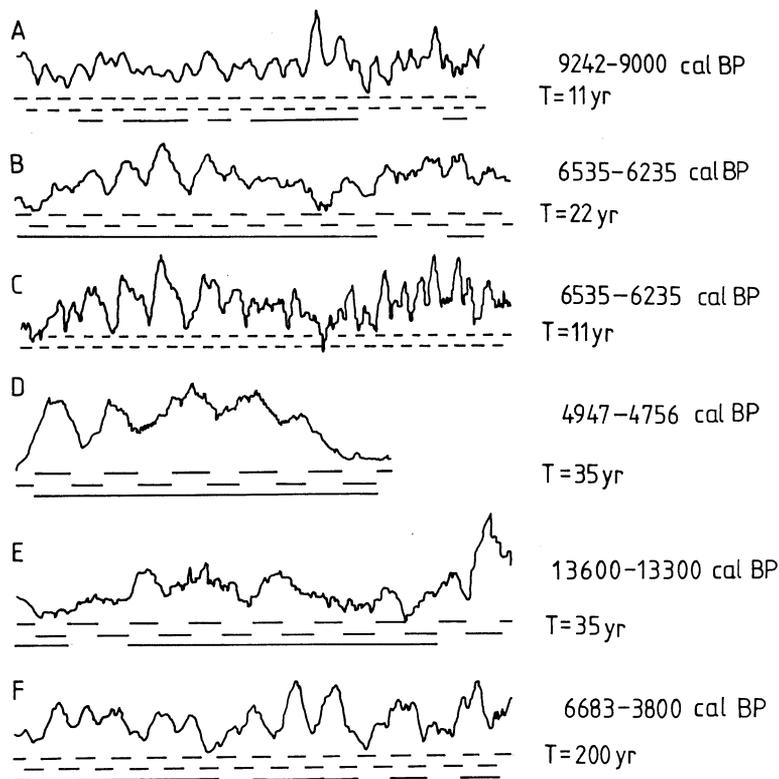


Fig 3. Examples of periodicities in sequence of varve thicknesses of core G1

from ca 7000 to 4500 cal BP, with relatively high amplitude and stable phase (Walanus, 1988b). Illustrative examples are shown in Figure 3. The analysis of periodicities in this sequence is very difficult, as the data are noisy, and moreover, at least three sources of noise could be distinguished: 1) noise in the extraterrestrial signal; 2) noise connected with unknown environmental disturbances of the lake ecosystem; 3) noise introduced by coring. The influence of each of the quoted sources of error on the analyzed sequence of input data should be carefully studied and evaluated before attempting better understanding and interpretation of results of periodicity analysis in terms of environmental or extraterrestrial processes. We expect that the methods developed for analyzing tree-ring width data, and their results will be very important in studying varve sequences.

#### FLUCTUATIONS OF WATER LEVEL IN GOŚCIAŻ LAKE

Mass balance equations, describing the carbon and  $^{14}\text{C}$  content of a lake, introduced by Broecker and Walton (1959), provide the relations between  $^{14}\text{C}$  concentration in bicarbonates dissolved in lake water and the

ratio of lake volume to its area,  $V/A$ . These equations were applied by Benson (1978) and Peng, Goddard and Broecker (1978) to derive corrected  $^{14}\text{C}$  ages of carbonate sediments using estimated values of  $V/A$  ratio and other parameters incorporated in this model.

Laminated sediments of Gościqz Lake provide an independent absolute chronology and thus permit calculation of initial  $^{14}\text{C}$  activity of bicarbonates from measured  $^{14}\text{C}$  ages of carbonate fractions of lake marl. The equation of the Broecker-Walton model may be, in this case, rewritten to provide estimates of the ratio  $V/A$ , which may be interpreted as mean water depth of the lake. After some transformation, we get a simple relation between mean water depth,  $H$ , and  $^{14}\text{C}$  dilution factor,  $q_L$ , of lake carbonates at the moment of their sedimentation

$$H = Aq_L^{-1} - B, \quad (1)$$

where  $A$  and  $B$  are combinations of model parameters, assumed constant during the entire history of the lake. Limnologic and geochemical considerations, accounting for data quoted by Benson (1978) and Peng, Goddard and Broecker (1978) lead to the estimate for the ratio of parameters  $A$  and  $B$ :  $0.90 < A/B < 0.98$ . Results of model calculations are shown in Figure 4 (Pazdur & Starkel, 1988) as relative changes of lake water depth (with respect to AD 1830–1850 level),

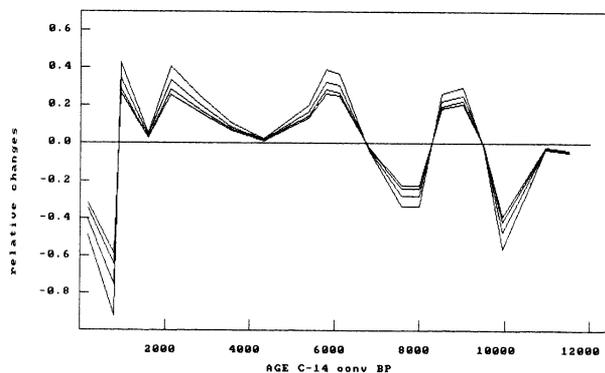


Fig. 4. Fluctuations of lake level reconstructed from  $^{14}\text{C}$  measurements of carbonate fraction of laminated sediment (according to Pazdur & Starkel, 1988)

$$\frac{H}{H_0} = \frac{(A/B)q_L^{-1} - 1}{(A/B)q_0^{-1} - 1} \quad (2)$$

for selected values of the ratio  $A/B = 0.9, 0.93, 0.96, 0.98$ . Comparison of the obtained curve with paleohydrologic and paleoenvironmental data, gathered by Ralska-Jasiewiczowa and Starkel (1988) from studies of lakes, mires and rivers in Poland, shows reasonable coincidences. Moreover, a dis-

tinct shift of rhythmic variations of mean water depth of Gościąg Lake, with respect to generalized paleohydrologic changes, is also visible. This shift in time may be explained by periodic floods of the lake basin and incising of the Ruda creek, draining this lake system, as well as changes of the Vistula River channel, which forms the basal level for Ruda creek and other small tributaries. These concepts will be verified in future studies of sediment sequences in the littoral area of Gościąg Lake and in the valley of Ruda creek.

#### SUMMARY AND CONCLUSIONS

Our results were obtained during the first two years of study of laminated sediments of Gościąg Lake, based on 3 cores, taken in 1985 and 1987. Measurements of varve thickness permit exact matching of the cores; good agreement between  $^{14}\text{C}$  dates obtained on core G0 (calibrated after correcting for  $^{14}\text{C}$  deficiency) and a varve chronology established on cores G1 and G2 suggests annual lamination in the lake. Additional evidence for annual lamination is provided by distinct periodical changes of varve thickness, especially by solar 11- and 22-yr cycles in the complete sequence, covering almost 10,000 yr. Definite proof of the annual character of lamination and exact dating of the floating varve chronology is expected from a correlation of varve thickness sequence with dendrochronologic records. Additional corings are necessary to extend this chronology to the recent three millennia where lamination is seriously disturbed. After completing this initial stage of the investigation with a reliable absolute varve chronology, the laminated sediments of Gościąg Lake will provide an important tool for studying geophysical and astrophysical processes during the Late Glacial and early Holocene, first of all through AMS dating of terrestrial macrofossils. We also expect that patterns of solar activity variations during this time will be reconstructed from the varve record, as was recently suggested by Bracewell (1988).

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