

RADIOCARBON AND STABLE CARBON ISOTOPE ANALYSES OF LAND SNAILS FROM THE CHINESE LOESS PLATEAU: ENVIRONMENTAL AND CHRONOLOGICAL IMPLICATIONS

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ABSTRACT. Paired radiocarbon and stable carbon analyses have been carried out on aragonite shells and organic soft bodies of snails from the Chinese Loess Plateau in order to explore the possibility of using these kinds of samples as environmental and chronological indicators. Results show that the soft bodies exhibit ^{14}C concentrations similar to those of plant leaves, indicating that carbon in the soft bodies is fixed from organic diets. The aragonite shells are depleted in ^{14}C compared to the soft bodies due to ingestion of ^{14}C -depleted carbonate. This depletion shows a consistent pattern across the Chinese Loess Plateau, implying a good potential for the snail shells to be applicable for ^{14}C dating with a simple correction. The $\delta^{13}\text{C}$ values measured for aragonite shells display a linear relationship with those obtained for the soft bodies with a constant offset. In addition, the carbon derived from organic diets accounts for more than 70% of the total shell carbon. This fact suggests that stable carbon isotope composition of aragonite shells mainly reflects that of organic diet, and could be used as a reliable indicator of paleodiet in the Chinese Loess Plateau.

INTRODUCTION

Land snail shells have been used in paleoenvironmental (Goodfriend and Ellis 2002; Stott 2002; Balakrishnan et al. 2005) and chronological (Mook and Vogel 1968; Magaritz and Heller 1980; Brennan and Quade 1997; Pigati et al. 2004; Mastronuzzi and Romaniello 2008) studies in the past decades. However, the environmental and chronological significance of aragonite shells is far from being understood because the sources and dynamics involved in the deposition of the aragonite shells needs to be further defined. Some authors indicate that organic diets are the major source and control the carbon isotope composition of the shell carbon (DeNiro and Epstein 1978; Goodfriend and Ellis 2002; Stott 2002; Balakrishnan et al. 2005). Magaritz et al. (1981) emphasize the land-air boundary CO_2 in controlling the isotopic composition of the shell carbonate with a minor effect from diet. Radiocarbon studies suggest that carbonate carbon (most are ^{14}C -depleted) also contribute to the shell formation (Goodfriend and Hood 1983; Goodfriend 1987; Pigati et al. 2004; Quarta et al. 2007; Romaniello et al. 2008).

The possible sources for the shell carbon, including atmospheric CO_2 , inorganic carbonate, and organic diets, have distinct ^{14}C and stable isotopic signatures. Thus, paired analysis of ^{14}C and stable carbon isotopes of the shell, organic bodies, and organic diets of the snail can provide valuable information on the processes and sources associated with the formation of aragonite shells, particularly for the definition of the proportions of the different sources, which has great significance for unraveling the paleoenvironmental and chronological implications of the carbon isotopes of snail shells. However, previous studies concentrate mostly on the carbon isotopes of snail shells, or stable carbon isotope alone.

In this study, the ^{14}C activities and $\delta^{13}\text{C}$ of both the bodies and shells of living snails collected across the Chinese Loess Plateau were measured. The main object of this study is to evaluate the chronological and environmental significance of the aragonite shell in this region by estimating the proportions of the different carbon origins incorporated into the shells.

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MATERIALS AND METHOD

The live land snails were collected mainly from the Chinese Loess Plateau (Figure 1), and cover different geographic settings across the region in order to have a wide representation of the study area. Each sample consists of at least 40 individuals, thus reducing the effect of individual anomalies. After collection, living snails were drowned in boiled deionized water and the soft parts were separated from the shells using forceps. Shells were not powdered during pretreatment to minimize the potential for adsorption of atmospheric ^{14}C (Samos 1949). At the same time, the leaves of deciduous poplars were collected for determining the ^{14}C activity of atmospheric CO_2 . Nineteen samples, composed only of adult *Cathaica*, were selected for coupled stable and ^{14}C isotopic analysis because of its wide distribution and abundance over the loess plateau. For the other samples (30), only $\delta^{13}\text{C}$ was analyzed.

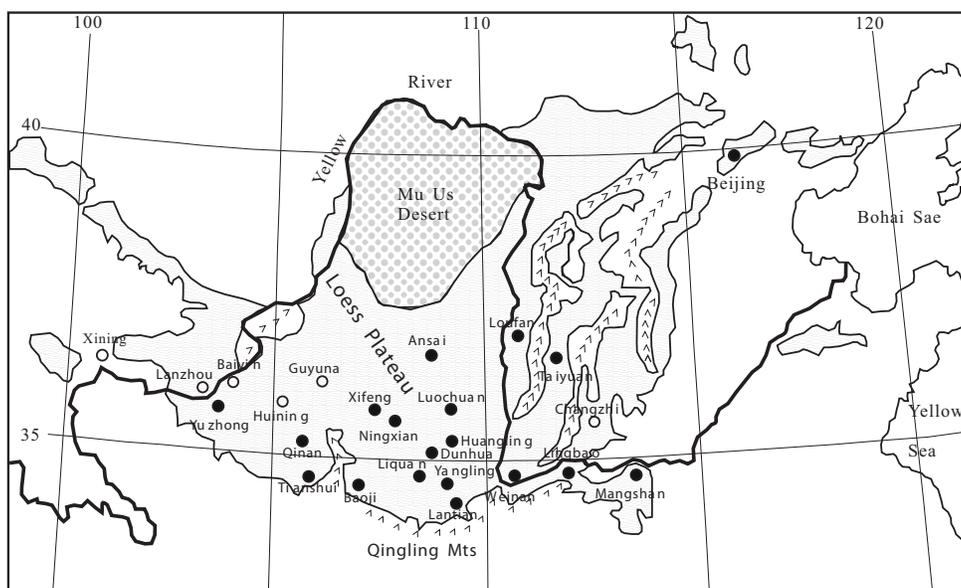


Figure 1 Locations of the sampling sites in the Chinese Loess Plateau. Filled circles represent the samples for both radiocarbon and stable carbon isotope analysis; empty circles indicate those only for $\delta^{13}\text{C}$ analyses.

The live snail shells were treated with 6% NaOCl for 48–72 hr to remove organic matter, washed repeatedly, and sonicated for 5–10 min to remove adhering materials. The cleaned shells were briefly washed with diluted HCl to remove adhering carbonate, rinsed with distilled water, then dried using a filter vacuum, and dried further in a vacuum oven overnight at 70 °C.

The soft bodies, and plant leaves as well, were treated with 10% HCl for 6 hr at 25 °C, washed repeatedly, sonicated for 5–10 min, dried using a filter vacuum, and further dried in a vacuum oven at 70 °C overnight. The dried soft bodies and leaves were crushed and amalgamated using a mortar and pestle for carbon isotopic analyses.

The ^{14}C activities of the snail shells and soft bodies were measured by liquid scintillation spectrometry with benzene as the scintillation solvent. The shells were converted into CO_2 by acid hydrolysis by using HCl under vacuum, and the organic matter powder was combusted in an oxygen stream under a temperature of 900 °C. The CO_2 produced was purified, collected with liquid nitrogen, and then reacted with lithium under 800 °C to be converted into Li_2C_2 . After cooling, C_2H_2 was obtained

through hydrolysis of the Li_2C_2 , and finally catalyzed into benzene. The benzene was then mixed with the scintillator, a mixture of MSB+Butyl PDB. ^{14}C activity was measured by liquid scintillation spectrometry (QuantulusTM 1220).

For stable carbon isotope analysis, the powders of aragonite shells and organic matter (soft bodies and leaves) were reacted with 100% phosphoric acid and combusted offline, respectively. The CO_2 produced was collected with liquid nitrogen, and isotopic ratios of $^{13}\text{C}/^{12}\text{C}$ were measured with a MAT 252 mass spectrometer. The isotopic data were reported in the conventional notation as per mil (‰) deviations relative to the PDB standard with uncertainties (1σ) of $\pm 0.1\text{‰}$.

RESULTS

Carbon isotope compositions of the aragonite shells and soft bodies of the snail samples from the Chinese Loess Plateau are plotted in Figure 2. The stable isotopic composition of snail shells ($\delta^{13}\text{C}_{\text{shell}}$) varies from -13.1‰ to -4.3‰ . The soft bodies have an isotopic value ($\delta^{13}\text{C}_{\text{body}}$) ranging from -26.8‰ to -18.1‰ . The $\delta^{13}\text{C}_{\text{shell}}$ displays a linear relationship with $\delta^{13}\text{C}_{\text{body}}$ with an offset of $14.0 \pm 0.7\text{‰}$ (Figure 2).

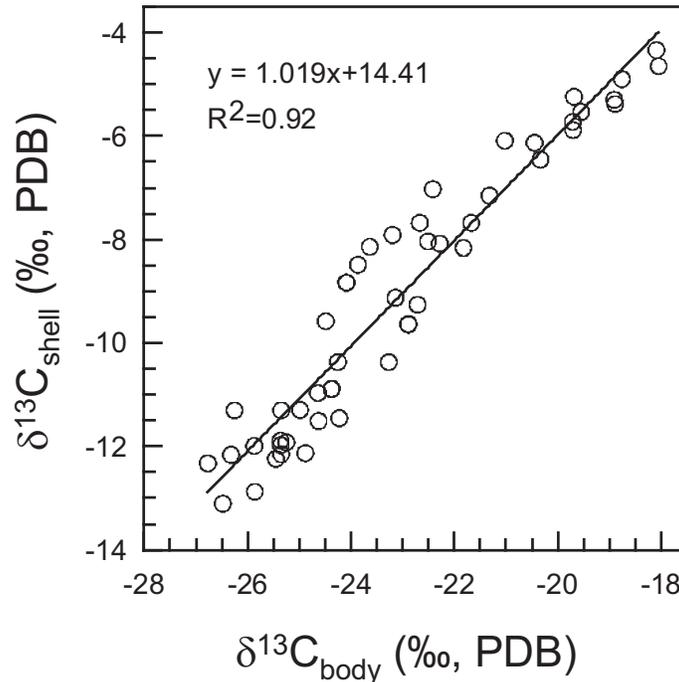


Figure 2 Stable carbon isotope relationship between aragonite shells and soft bodies of the living land snails from the Chinese Loess Plateau.

^{14}C activities of the snail soft bodies range from 106.40 to 108.89 pMC with an average of 107.9 ± 0.6 pMC. These results display no distinguishable difference from those measured for the tree leaves (108.08 ± 1.87 pMC; see Table 1), indicating that the carbon isotope compositions of the snail bodies are derived from organic diets without discrimination. In contrast, the aragonite shells are depleted in ^{14}C (93.75–99.57 pMC) compared to the bodies (Table 1), suggesting a contribution of ^{14}C -depleted carbonate to the shell.

Table 1 Measured and calculated results for live snails across the Chinese Loess Plateau.

Location	Species	$\delta^{13}\text{C}_{\text{body}}$	$\delta^{13}\text{C}_{\text{shell}}$	$A_{\text{shell}}^{\text{a}}$	$A_{\text{body}}^{\text{a}}$	f_{atm} (%)	f_{calc} (%)	f_{met} (%)	Age anomaly
Yuzhong	<i>C. cunlunensis</i>	-21.7	-7.7	95.50 ± 0.70	108.16 ± 1.34	20	17	64	1256 ± 112
Luochuan	<i>C. pulveratrix</i>	-24.3	-10.4	96.44 ± 0.71	106.63 ± 0.86	14	12	74	1070 ± 88
Weinan	<i>C. pulveratrix</i>	-26.3	-12.2	93.80 ± 0.69	106.40 ± 1.29	12	14	74	1285 ± 114
Mangshan	<i>C. pulveratrix</i>	-22.9	-9.6	96.05 ± 0.81	108.25 ± 1.27	12	14	74	1211 ± 116
Taiyuan	<i>C. pulveratrix</i>	-21.3	-7.2	98.28 ± 0.92	108.19 ± 1.07	21	12	67	1029 ± 109
Ansai	<i>C. cunlunensis</i>	-26.3	-11.3	96.68 ± 0.96	107.12 ± 1.13	17	13	71	1106 ± 116
Beijing	<i>C. pulveratrix</i>	-22.3	-8.1	95.77 ± 0.87	108.15 ± 1.07	20	14	66	1237 ± 108
Fenhe	<i>C. pulveratrix</i>	-25.4	-12.0	93.75 ± 0.89	107.66 ± 1.07	9	15	76	1371 ± 110
Xifeng	<i>C. pulveratrix</i>	-22.4	-7.0	97.83 ± 0.91	108.31 ± 0.96	27	13	60	1095 ± 103
Qinan	<i>C. pulveratrix</i>	-25.5	-12.2	97.78 ± 0.88	108.20 ± 0.97	8	12	80	1071 ± 102
Huangling	<i>C. pulveratrix</i>	-23.9	-8.5	98.35 ± 0.88	108.20 ± 0.96	24	12	64	1048 ± 101
Liquan	<i>C. pulveratrix</i>	-24.6	-11.5	97.42 ± 0.83	107.89 ± 0.94	9	12	79	1074 ± 98
Ningxian	<i>C. pulveratrix</i>	-20.3	-6.5	96.97 ± 1.03	108.88 ± 0.93	22	14	64	1181 ± 109
Tianshui	<i>C. pulveratrix</i>	-25.9	-12.0	98.86 ± 0.81	108.00 ± 1.04	12	11	77	978 ± 102
Baoji	<i>C. pulveratrix</i>	-22.5	-8.0	95.45 ± 1.03	107.93 ± 1.03	21	14	65	1252 ± 116
Yangling	<i>C. pulveratrix</i>	-24.4	-10.9	96.06 ± 0.84	107.70 ± 0.81	11	13	76	1177 ± 93
Lantian	<i>C. pulveratrix</i>	-25.2	-11.9	99.57 ± 0.82	107.32 ± 1.05	9	10	81	861 ± 103
Lingbao	<i>C. pulveratrix</i>	-23.1	-9.1	95.34 ± 0.92	107.51 ± 0.96	16	14	70	1226 ± 106
Dunhua	<i>C. pulveratrix</i>	-24.1	-8.8	95.39 ± 0.88	108.20 ± 0.96	22	15	63	1292 ± 103
Leaf		-28.17			108.08 ± 1.87				

^a A uncorrected for isotopic fractionation; PMC = % modern carbon (AD 1950).

DISCUSSION

The mass-balance model was often used to estimate the fractions of carbon of different origins and thus interpret the ^{14}C depletion in snail shells (Goodfriend and Hood 1983; Romaniello et al. 2008). This model is based on the following assumptions: 1) isotope fractionation occurs when atmospheric CO_2 passes through the mantle into hemolymph, and the aragonite shell is deposited from bicarbonate pool; and 2) there is no fractionation between the metabolic CO_2 and HCO_3^- pool. Based on this model, as suggested by Stott (2002), the $\delta^{13}\text{C}$ of snail shells depends on the habits of snails, their dietary intake, and the $p\text{CO}_2$ at ground level, and thus it would be difficult to interpret the $\delta^{13}\text{C}$ of shells solely in terms of dietary intake. Obviously, this contradicts the linear relationship found between the $\delta^{13}\text{C}$ of shells and bodies/diets of both cultured (DeNiro and Epstein 1978; Stott 2002; Metref et al. 2003) and field-collected snails (this study), and contradicts a number of observations that show the $\delta^{13}\text{C}$ of land snail shell reflects the carbon isotope composition of their diets (Magaritz et al. 1981; Goodfriend and Ellis 2002; Balakrishnan et al. 2005).

In fact, the isotopic fractionations involving the metabolic CO_2 fractionation can well explain the $\delta^{13}\text{C}$ linear relationship and constant offset between the shells and bodies of the cultured snails (Stott 2002). The total fractionation between the metabolic CO_2 and aragonite shells is 12.7‰, 10‰ from the fractionation between the metabolic CO_2 and HCO_3^- pool from which the shell is deposited (Mook et al. 1974; Zhang et al. 1995) and 2.7‰ from that between aragonite and bicarbonate pool (Rubinson and Clayton 1969). In our study, the average offset between the aragonite shells and soft bodies is around 14.0‰, which is close to the total isotopic fractionations between the metabolic CO_2 and aragonite shells, supporting a significant fractionation between the metabolic CO_2 and HCO_3^- pool that cannot be ignored during estimation of the proportions of the shell carbon with various origins.

Taking into account the mass balance, the following equations can be established:

$$R_{\text{total}} = f_{\text{met}}R_{\text{met}} + f_{\text{calc}}R_{\text{calc}} + f_{\text{atm}}R_{\text{atm}} \quad (1)$$

$$f_{\text{met}} + f_{\text{calc}} + f_{\text{atm}} = 1 \quad (2)$$

where R and f are the ¹³C/¹²C or ¹⁴C/¹²C ratio and carbon fraction, respectively, and the subscripts *total*, *met*, *calc*, and *atm* represent respectively the integrate, metabolic, soil carbonate (mostly calcite), and atmospheric carbon sources for the CO₂-H₂O-CaCO₃ system in the snail body fluid from which the shell is deposited.

In the CO₂-H₂O-CaCO₃ system, the carbon isotopic ratio is dominated by HCO₃⁻ and CO₂ including gaseous and dissolved phases. When isotopic fractionation reaches equilibrium, there is

$$R_{\text{total}} = fR_{\text{HCO}_3^-} + (1-f)R_{\text{CO}_2} = R_{\text{HCO}_3^-} \left(f + \frac{1-f}{\alpha} \right) \quad (3)$$

where α is the average value of fractionation factors between HCO₃⁻ and dissolved and gaseous CO₂ ($\approx \left(\alpha_{\text{CO}_2(\text{aq.})}^{\text{HCO}_3^-} + \alpha_{\text{CO}_2(\text{gas.})}^{\text{HCO}_3^-} \right) / 2$), and $R_{\text{HCO}_3^-}$ and R_{CO_2} are the carbon isotopic ratios for HCO₃⁻ and CO₂ in the system, respectively. f is the molar fraction of HCO₃⁻.

In the system CO₂-H₂O-CaCO₃, there are 2 moles of HCO₃⁻ for every 1 mole of Ca²⁺ at normal pH. Therefore, the molar concentration of Ca (f_{Ca}) is only half that of HCO₃⁻ (f), i.e. $f_{\text{Ca}} = 1/2(f)$, or $f = 2f_{\text{Ca}}$ (and $f = 2f_{\text{calc}}$). Thus, Equation 3 becomes

$$R_{\text{total}} = R_{\text{HCO}_3^-} \left(2f_{\text{calc}} + \frac{1-2f_{\text{calc}}}{\alpha} \right) \quad (4)$$

Combing Equations 1 and 4, we get Equation 5:

$$R_{\text{HCO}_3^-} = \alpha \frac{f_{\text{met}}R_{\text{met}} + f_{\text{calc}}R_{\text{calc}} + f_{\text{atm}}R_{\text{atm}}}{1 + 2f_{\text{calc}}(\alpha - 1)} \quad (5)$$

Introducing $\delta^{13}\text{C}_{\text{sample}} = \left(\frac{R_{\text{sample}}}{R_{\text{sample}}} - 1 \right) \times 1000$, Equation 5 then becomes

$$\delta^{13}\text{C}_{\text{HCO}_3^-} + 1000 = \alpha_{13} \frac{f_{\text{met}}\delta^{13}\text{C}_{\text{met}} + f_{\text{calc}}\delta^{13}\text{C}_{\text{calc}} + f_{\text{atm}}\delta^{13}\text{C}_{\text{atm}} + 1000}{1 + 2f_{\text{calc}}(\alpha_{13} - 1)} \quad (6)$$

where α_{13} stands for the ¹³C-¹²C fractionation factor. At 20 °C, $\alpha_{13} = 1.0091$ (Mook et al. 1974) and $\delta^{13}\text{C}_{\text{HCO}_3^-} = \delta^{13}\text{C}_{\text{shell}} - 2.7$ (Mook and Vogel 1968; Rubinson and Clayton 1969), where $\delta^{13}\text{C}_{\text{shell}}$ is the measured isotope composition of the snail shell. $\delta^{13}\text{C}_{\text{met}}$ can be obtained by measuring the $\delta^{13}\text{C}$ of the soft bodies ($\delta^{13}\text{C}_{\text{body}}$). $\delta^{13}\text{C}_{\text{atm}}$ is taken as -8‰ (Wahlen 1994). $\delta^{13}\text{C}_{\text{calc}}$ is the value of carbonates in the surface soil of the Chinese Loess Plateau.

Introducing $A_{\text{sample}} = {}^{14}\text{C}/{}^{12}\text{C}$ for the ¹⁴C activity, we modify Equation 5 as follows:

$$A_{\text{HCO}_3^-} = \alpha_{14} \frac{f_{\text{met}}A_{\text{met}} + f_{\text{calc}}A_{\text{calc}} + f_{\text{atm}}A_{\text{atm}}}{1 + 2f_{\text{calc}}(\alpha_{14} - 1)} \quad (7)$$

where α_{14} stands for the ¹⁴C-¹²C fractionation factor and $\alpha_{14} \approx \alpha_{13}^2$ (Wigley and Muller 1981). $A_{\text{HCO}_3^-}$, A_{met} , A_{atm} , and A_{calc} are ¹⁴C activities of the HCO₃⁻ pool, metabolic and atmospheric CO₂,

and soil carbonate, respectively. Since old carbonate is completely ^{14}C -depleted in most cases, $A_{\text{calc}} = 0$ can be applied in Equation 7. $A_{\text{HCO}_3^-} = (1 + 0.0027)^{-2} A_{\text{shell}}$, A_{shell} and A_{met} , and A_{met} are the measured ^{14}C activities of shells and bodies (A_{body}).

As the carbon isotopic values of the carbonate in surface soil ($^{13}\delta_{\text{calc}}$), ^{14}C activities, and carbon isotopic ratios can be obtained from previous work or measurement, the terms f_{atm} , f_{met} , and f_{calc} become known through resolving the combined Equations 2, 6, and 7. According to Gu et al. (1991), the $^{13}\delta_{\text{carb}}$ of the surface soil ranges from -3‰ to -9‰ , with an average of -6‰ in the Chinese Loess Plateau. In this study, the average value is used for $^{13}\delta_{\text{carb}}$. This could cause a 1–2% shift for the calculated f_{atm} and f_{met} , but no effect for f_{calc} is observed.

The calculated results of the different carbon sources are presented in Table 1. The shell carbon is dominated by its ingested diet, which accounts for 60–81%. The fractions from atmospheric CO_2 vary from 8% to 27%; the proportions of carbon originated from carbonate display a narrow variation with an average of $13 \pm 2\%$. These results suggest that the organic diet is the main contributor for the shell carbon.

Concerning ^{14}C , estimating the proportion of ^{14}C -depleted carbonate in the snail shells allows calculation of the corresponding age anomaly. In order to calculate the age anomaly, fractionation correction of the A_{shell} term is necessary. The fractionation-corrected $A_{\text{corrected}}$ term can be calculated from the measured term (A_{measured}) by using the standard formula of Stuiver and Polach (1977):

$$A_{\text{corrected}} = A_{\text{measured}} \left[1 - \frac{2(\delta^{13}\text{C} + 25)}{1000} \right] \quad (8)$$

This equation is valid only for the component of shell carbon derived directly or indirectly from atmospheric CO_2 (Pigati 2003). Therefore, Equation 8 should be

$$A_{\text{corrected}} = A_{\text{measured}} \left[1 - \frac{2(\delta^{13}\text{C}_{\text{nlc}} + 25)}{1000} \right] \quad (9)$$

where $\delta^{13}\text{C}_{\text{nlc}}$ is the carbon isotope of the shell carbonate derived from non-carbonate sources. Substituting $\delta^{13}\text{C}_{\text{HCO}_3^-} = \delta^{13}\text{C}_{\text{shell}} - 2.7$ and $\alpha_{13} = 1.0091$ in Equation 6, the following is obtained

$$\delta^{13}\text{C}_{\text{shell}} = \frac{f_{\text{met}}(1.0091\delta^{13}\text{C}_{\text{met}} + 11.8) + f_{\text{atm}}(1.0091\delta^{13}\text{C}_{\text{atm}} + 11.8) + f_{\text{calc}}(1.0091\delta^{13}\text{C}_{\text{calc}} - 6.35)}{1 + 2f_{\text{calc}}(1.0091 - 1)}$$

and the proportion of the contribution of the non-carbonate sources (P) is

$$P = \frac{f_{\text{met}}(1.0091\delta^{13}\text{C}_{\text{met}} + 11.8) + f_{\text{atm}}(1.0091\delta^{13}\text{C}_{\text{atm}} + 11.8)}{f_{\text{met}}(1.0091\delta^{13}\text{C}_{\text{met}} + 11.8) + f_{\text{atm}}(1.0091\delta^{13}\text{C}_{\text{atm}} + 11.8) + f_{\text{calc}}(1.0091\delta^{13}\text{C}_{\text{calc}} - 6.35)}$$

Thus, Equation 9 becomes

$$A_{\text{corrected}} = A_{\text{shell}} \left[1 - \frac{2(P_{\text{nlc}}\delta^{13}\text{C}_{\text{shell}} + 25)}{1000} \right] \quad (10)$$

P_{nlc} displays a rather consistent variation across the Chinese Loess Plateau, ranging from the 0.71 to 0.86, with an average of 0.79 ± 0.04 . Therefore, the fractionation correction can be reduced to

$$A_{\text{corrected}} = A_{\text{shell}} \left[1 - \frac{2(0.79\delta^{13}\text{C}_{\text{shell}} + 25)}{1000} \right]$$

and the age anomalies (Δ) can be calculated by

$$\Delta = -8033 \ln \left(\frac{A_{\text{shell}} \left[1 - \frac{2(0.79\delta^{13}\text{C}_{\text{shell}} + 25)}{1000} \right]}{A_{\text{atm}}} \right)$$

The results in Table 1 show that the age anomalies range from 860 ± 66 to 1371 ± 134 yr. The measured age anomalies are rather convergent (1148 ± 127 yr), demonstrating that the ¹⁴C of the *Cathaica* shells has the potential for chronological application after a simple correction in the Chinese Loess Plateau. The convergent ¹⁴C age anomalies of the shell carbonate across the Chinese Loess Plateau might be associated with the similarities of both the ecological habits of the studied snails and the calcium sources. According to Goodfriend (1987), the age anomalies correspond closely with the ecology of the snails, and the difference in availability of calcium for shell growth. Snails with a continual availability of calcium for growth do not need to take up as large an amount of carbonate as plant-dwelling species, to which carbonate is available only by ground feeding (Goodfriend 1987), and have less ¹⁴C-depletion of the shell carbonate. In our study, the samples for ¹⁴C analysis are composed of the same species, *Cathaica*, which should have similar ecological habits. Likewise, topsoil in the Chinese Loess Plateau, which has uniform origination and is enriched in calcium minerals, can provide a continual supply of calcium for snail growth. As pointed out by Pigati (2003) and Quarta et al. (2007), the age anomaly correction can result in a significant enlargement of uncertainty, which may limit its applications for higher chronological precision, such as archaeological chronology.

CONCLUSION

The stable and ¹⁴C isotope composition of aragonite shells and soft bodies of the *Cathaica* collected in different sites of the Chinese Loess Plateau have been analyzed in order to estimate the possible carbon sources, paleovegetation and chronological implications of the shells. The ¹⁴C activities of snail bodies cannot be discriminated from that of the plant leaves. The ¹⁴C activities of shell aragonite are significantly lower than those measured for bodies, indicating the incorporation of ¹⁴C-depleted carbon from carbonate into the shell. The age anomalies of the shells caused possibly by ingested carbonate concentrated around 1148 ± 127 yr, suggesting the potential of the *Cathaica* shell for chronological application in the Chinese Loess Plateau. The $\delta^{13}\text{C}$ of shell aragonite displays a linear relationship with the carbon isotopic composition of bodies with a constant offset of 14.0‰. The fractions of the different carbon sources in snail shells have been estimated by a model, which considers both mass balance and isotope fractionations between different carbon phases during the formation of the snail shells. The results show that the fraction of the carbon from organic diets accounts for 60–81%, and the proportions of the carbon from carbonate and atmospheric CO₂ ranges from 8% to 27% and 10% to 17%, respectively, indicating the shell $\delta^{13}\text{C}$ of the *Cathaica* in the Chinese Loess Plateau mainly reflect that of organic diets.

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