

The Importance of Nutrients for Microbial Priming in a Bog Rhizosphere

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Abstract

Wetlands host anaerobic microbes which convert organic carbon into methane (CH₄), a powerful greenhouse gas. Wetland plants can influence which carbon compounds are available for microbial processing by exuding freshly fixed carbon from their roots. Notably, exudation of carbon from plant roots can trigger microbial priming: the process of new carbon stimulating the microbial community into processing more soil carbon than they otherwise would have. This study utilized high resolution Fourier transform ion cyclotron mass spectrometry (FT-ICR-MS) analysis to probe the chemical composition of soil organic compounds from the rhizosphere of *Carex aquatilis*, a common wetland sedge, which is known to have stimulated microbial priming within a peat soil. The goal was to identify what types of molecules are created or lost during microbial priming in the wetland rhizosphere and thus advance mechanistic understanding of the process. FT-ICR-MS analysis demonstrated that more microbial transformations of carbon occurred among water-soluble compounds than among hydrophobic compounds, but that some hydrophobic compounds were processed as well. Crucially for understanding microbial priming, the root exudates triggered increased processing

28 of high molecular weight molecules regardless of nutrient content but processed low molecular
29 weight compounds only if they contained essential nutrients. The pattern of which molecules
30 are processed and which are not is evidence for a selective priming effect in which some types
31 of carbon compounds are used at an increased rate, while others are not.

32 Introduction

33 Wetlands host anaerobic microbes which convert organic carbon into methane (CH₄), a
34 powerful greenhouse gas. Globally, this wetland-generated CH₄ is responsible for
35 approximately a third of total CH₄ emissions (Ciais et al. 2013), and CH₄ in the atmosphere is
36 responsible for 15-19% of total greenhouse gas radiative forcing (Intergovernmental Panel on
37 Climate Change 2014). This substantial contribution to the greenhouse effect is dynamic and
38 responds to environmental changes; wetland CH₄ emissions are the primary driver of
39 interannual variability in global emissions (Ciais et al. 2013). The microbial communities that
40 produce CH₄ are sensitive to a variety of variables including temperature (Westermann 1993;
41 Yvon-Durocher et al. 2014), pH (DeLaune et al. 1986; Kotsyurbenko et al. 2007), redox
42 environment (Fetzer and Conrad 1993; Fritz et al. 2011; Boye et al. 2017), and the composition
43 of the organic carbon available to them (Updegraff et al. 1995; Girkin et al. 2018). Two of those
44 variables can in turn be dynamically impacted by the presence of plants: redox environment
45 through gas transport effects and carbon availability through root exudation. As such,
46 influencing the microbial environment makes plants a key factor in wetland CH₄ emissions.

47 The types of carbon molecules present in the soil is an important variable influencing
48 wetland CH₄ emissions because it determines what metabolic processes microbes can conduct.
49 It was once thought that the structure of a molecule determined whether it was “labile” or
50 “recalcitrant” to microbial degradation (Sollins et al. 1996; Hedges et al. 2000; Lützow et al.
51 2006), and authors proposed various means by which the “recalcitrance” could be established
52 from the molecular structure (Lützow et al. 2006; D’Andrilli et al. 2015). Over time, the scientific
53 consensus has shifted to recognize that all carbon compounds are ultimately capable of being
54 microbially processed, albeit at different rates, and that the order of microbial preference can
55 depend on a variety of factors including the composition of the microbial community and the

56 redox state of the environment (Schmidt et al. 2011; Lehmann and Kleber 2015). While authors
57 continue to use the language of “recalcitrance” and “lability”, sometimes acknowledging the
58 changing understanding of these terms (e.g., Zhang et al. 2018) and sometimes not (e.g.,
59 Barcellos et al. 2018), these studies still provide valuable insights into situations where certain
60 classes of molecules are more or less bioavailable to the microbes present.

61 The situational bioavailability of carbon compounds can depend on traits of the
62 molecules as well as on environmental factors. One fundamental constraint on microbial use of
63 carbon is the limited pore size of cell walls, which dictates that any molecule that is either too
64 large or not water soluble must be degraded by exo-enzymes before the microbe can take it up
65 (Benz and Bauer 1988; Fenchel et al. 2012). However, microbes have also been shown to
66 preferentially utilize high molecular weight compounds in a variety of environments despite the
67 extra step of extracellular degradation (Amon and Benner 1996; McArthur and Richardson
68 2002; Antony et al. 2012). As microbes use large compounds they metabolize them into smaller
69 compounds, thus shifting the size distribution of molecules (Amon and Benner 1996; McArthur
70 and Richardson 2002; Antony et al. 2012; Pracht et al. 2018). Another property that has been
71 studied is aromaticity. More aromatic compounds have historically been thought to be less
72 bioavailable (Lützow et al. 2006), but new evidence indicates that is likely only true in oxic
73 environments; when oxygen is limiting, thermodynamic limitations dominate and aromaticity is
74 not controlling (Pracht et al. 2018). In those oxygen-limited or highly reduced environments,
75 the nominal oxidation state of carbon (NOSC) can be the most important factor (Keiluweit et al.
76 2016; Boye et al. 2017).

77 Wetland plants can influence what carbon compounds are available for microbial
78 processing by exuding freshly fixed carbon from their roots in the form of sugars,
79 carbohydrates, organic acids, amino acids, and a variety of other small, water-soluble
80 compounds (Smith 1976; Lugtenberg et al. 1999; Jones et al. 2009; Dommergues 2012). These
81 compounds stimulate the growth of beneficial microbial and fungal communities (Broeckling et
82 al. 2008). Plants also introduce higher molecular weight compounds to the soil through
83 sloughing of root cells and release of mucilage, but these compounds generally are less utilized
84 by the soil microbial community (Bais et al. 2006). Studying root exudates in soil is challenging

85 because root exudates often are quickly utilized by microbes and thus these compounds do not
86 remain in the soil for measurable periods of time. Analytical techniques can only measure what
87 remains in the soil. This under-detection of quickly metabolized compounds may bias results
88 towards compounds that are less bioavailable in the rhizosphere.

89 Root exudates must provide some benefit to the plant, as they are a carbon cost
90 (Marschner 1974). One way that root exudation can change the rhizosphere in a manner that
91 potentially benefits plants is through microbial priming. Priming is the mechanism by which the
92 addition of a new carbon source to soils stimulates microbes to degrade more of the existing
93 soil carbon than they would have in the absence of the new carbon source (Fontaine et al.
94 2007; Kuzyakov 2010; Ruirui et al. 2014; Ye et al. 2015). Priming can occur because the fresh
95 carbon causes changes to the microbial ecosystem, increasing the demand for carbon
96 compounds and/or the ability of microbes to process carbon compounds (Fontaine et al. 2004;
97 Kuzyakov 2010; Ye et al. 2015). Priming could benefit plants through nitrogen mining (N-
98 mining), a process by which the addition of nutrient-poor carbon compounds causes microbes
99 to degrade existing soil compounds which contain nitrogen (Craine et al. 2007; Ruirui et al.
100 2014). Wetland types such as bogs and some fens are often nitrogen limited (Aerts et al. 1992)
101 despite the large amount of nitrogen bound in slowly decaying plant matter (Turunen et al.
102 2004; Moore et al. 2005; Drewer et al. 2010). These wetlands are therefore an environment
103 where it is highly beneficial for plants to access the nitrogen already in the soil, using tactics
104 such as N-mining. Other studies have also found evidence of other types of nutrient mining, for
105 example for sulfur (Creamer et al. 2014). However, some recent studies have failed to find
106 evidence for priming in peat soils, such as Girkin et al. (2018), though that study used root
107 exudate analogs instead of actual plants.

108 This manuscript presents results from a plant-growth experiment that used peat from a
109 boreal bog and a typical wetland sedge, *Carex aquatilis*. Waldo et al (2019) previously
110 demonstrated, using isotopic tracers, that *Carex* in this experiment triggered microbial priming.
111 That result was based upon the finding that in the presence of plants the emission of soil-
112 derived CH₄ was increased by an order of magnitude, which was enough that in planted boxes
113 there was more CH₄ being derived from soil carbon than derived from root carbon (Waldo et al.

114 2019). Here, we present the results of Fourier transform ion cyclotron resonance mass
115 spectrometry (FT-ICR-MS) analysis of organic compounds in root, rhizosphere, and unplanted
116 soil samples from the Waldo et al. (2019) experiment. These results address the question of
117 what soil organic compounds were processed when priming was active, with the goal of
118 elucidating the factors controlling the occurrence of microbial priming. Understanding the
119 causes and mechanisms of microbial priming will help researchers better predict the fate of
120 wetland soil carbon in a future with predicted elevated boreal plant productivity.

121 **Materials and Methods**

122 *Sample Collection*

123 This study utilized samples collected from a previously described experiment (Waldo et
124 al. 2019). Briefly, we grew *Carex aquatilis*, a common boreal wetland sedge, for ten weeks in
125 rhizoboxes filled with bog peat collected from Interior Alaska. The bog site was described in
126 detail by Neumann et al. (2019). Control boxes consisted of peat with no plants. We monitored
127 CH₄ emissions throughout the experiment. During weeks 5 and 10 of the experiment, we
128 exposed 4 plants each to ¹³CO₂ by placing a clear fluxing hood on each rhizobox and injecting 15
129 mL of 99 atom% ¹³CO₂ into the headspace over a period of five consecutive days. This ¹³CO₂ was
130 photosynthesized and isotopically labeled the plants. Following labelling, we anaerobically
131 collected root and soil samples, destroying the plants.

132 For both box types (real plants and control) we collected soil samples from the same
133 locations, but the method of sample collection depended on box type. For planted boxes, we
134 collected roots simultaneously with the rhizosphere soil. We collected the samples at depths of
135 approximately 5 cm, 20 cm, and 35 cm. At each depth, we took three samples: one in the
136 center and one six cm from either edge of the box. For planted boxes, we cut root sections from
137 each location. We sonicated the root sections in phosphate-buffered solution (PBS) then moved
138 the root to a new container. All the soil that fell off the root during sonication we considered
139 rhizosphere, which is an operational definition of the rhizosphere which has been used
140 previously (White et al. 2015). We saved both roots and rhizosphere soil for analysis. For
141 unplanted control boxes, we sampled cubes of soil from the 9 standard locations. We stored all

142 samples at -20 °C before shipping them to the Environmental Molecular Sciences Laboratory in
143 Richland, WA, USA, where they were stored at -80 °C until analysis. In total, we analyzed 73
144 rhizosphere samples from 9 planted boxes (4 harvested after week 5 and 5 harvested after
145 week 10), 43 bulk soil samples from 5 unplanted boxes, and 6 root sections from 4 plants.

146 *Fourier transform ion cyclotron resonance mass spectrometry (FT-ICR-MS)*

147 A 12 Telsa FT-ICR-MS (Bruker Daltonics) was used for analyzing the extracts of carbon
148 compounds from the soil samples, based on the methods of Tfaily et al. (2015). We performed
149 three sequential extractions on each soil sample to extract compounds that dissolve in different
150 solvents: a water extraction for hydrophilic compounds, methanol extraction for compounds of
151 intermediate hydrophobicity, and chloroform extraction for hydrophobic compounds. The
152 extractions were performed sequentially in the stated order.

153 To perform the extractions, we thawed the samples at room temperature. We then
154 centrifuged (VWR Symphony 4417) samples at a rotational centrifugal force of 2720g for 5 min
155 and collected the supernatant containing the water-soluble compounds. To extract the
156 remainder of the water-soluble compounds, we used four cycles of centrifugation in which we
157 added 1 mL of water to the sample, shook it for 30 min, centrifuged it, and collected the
158 supernatant. The combined supernatant from the five centrifugations (the initial porewater and
159 four water additions) was acidified to pH 2 using nitric acid. This solution contained all of the
160 water-soluble compounds in the sample, in addition to unwanted salts. The water extraction
161 was coupled with a solid-phase extraction (SPE) to remove salts from the solution that
162 interfered with FT-ICR-MS readings according to Thorsten et al. (2008). We activated SPE
163 columns (PPL, 500 mg, Varian Mega Bond Elut, Varian Inc.) with 1 mL of methanol, then ran
164 the acidified water samples through the column. Salts moved through the column while carbon
165 compounds were retained. We then washed each column three times with 10 mM hydrochloric
166 acid and dried it by blowing nitrogen gas through the column. Last, we eluted the carbon
167 compounds off the column using 2 mL of methanol.

168 The methanol extraction was performed on the pellet of peat remaining after the final
169 centrifugation of the water extraction. We added 1 mL of methanol to each sample, shook it for

170 two hours, centrifuged it and collected the supernatant. After the methanol extraction, we
171 immediately added 1 mL of chloroform to the remaining peat pellet and let it sit quiescent
172 overnight at 4 °C. We then centrifuged the samples and collected the supernatant as with the
173 previous extractions. Pure chloroform does not run well in the ICR-MS instrument, so after the
174 extraction was completed the chloroform extracts were blended with methanol in a 1:1 ratio so
175 that they could be run directly on the instrument.

176 The water extraction contained water-soluble compounds, the methanol extraction
177 contained compounds on intermediate solubilities, and the chloroform extraction contained
178 compounds which were not soluble in water. All three extractions were then analyzed on a FT-
179 ICR-MS. We used a Bruker electrospray ionization source in negative ion mode with the
180 following settings: 0.1 s ion accumulation time, 200 scans, range of 98-900 m/z , resolution 260k
181 at 400 m/z . Negative ion mode was selected because it has been shown to be more effective
182 for ionization of organic molecules (Stenson et al. 2003). The instrument was flushed with a
183 mixture of water and methanol between samples, and methanol blanks were run at the
184 beginning and end of each session to check for carryover from the samples. The instrument was
185 calibrated weekly with a standard solution containing $C_2F_3O_2$, $C_6HF_9N_3O$, $C_{12}HF_{21}N_3O$,
186 $C_{20}H_{18}F_{27}N_3O_8P_3$, and $C_{26}H_{18}F_{39}N_3O_8P_3$.

187 We processed the data using DataAnalysis software (v4.2, Bruker Daltonics) to convert
188 the raw spectra into mass to charge (m/z) ratios, which we then assigned molecular formulas
189 using a Compound Identification Algorithm (Kujawinski and Behn 2006; Minor et al. 2012). We
190 assigned formulas to peaks with a signal to noise ratio of at least 7, and mass measurement
191 error of less than one ppm. The only atoms we allowed to be assigned were the most common
192 isotope for each of C, H, O, N, S, and P, plus ^{13}C . Formulas were only assigned which included at
193 least one each of C, H, and O, which had no more than 100 C, 200 H, 30 O, 20 N, 10 S, and 6 P,
194 and which had at least four O per P.

195 *Laser ablation electrospray ionization mass spectroscopy (LAESI-MS)*

196 To analyze compounds present on and in roots, we used laser ablation electrospray
197 ionization mass spectroscopy (LAESI-MS) (Stopka et al. 2017; Stopka et al. 2019). We mounted

198 2-4 cm long root sections on microscope slides using double-sided tape and ablated the surface
199 of the root using focused 7-nanosecond pulses from a mid-IR laser (IR Opolette HE 2731) set to
200 a 2.94 μm wavelength. The ablated material was carried into a Thermo Fischer Velos Orbitrap
201 Pro MS using an ion spray generated from a solution of 50% methanol (in DI water), 0.1% acetic
202 acid . The spray flowrate was at 500 nL/min minute and aligned on-axis to the MS orifice. Both
203 positive ion mode and negative ion mode were used to analyze the ions captured. We sampled
204 roots using a grid scan where we ablated the surface of the root in a rectangular grid with
205 between 300 and 400 μm between the centers of each ablation target. All raw spectra were
206 assigned molecular formula using the HMDB 4.0 database (Guo et al. 2017) through the
207 METASPACE platform (Palmer et al. 2017). This system only assigned formulas to m/z ratios
208 which correspond to an identified metabolite in the HMDB 4.0 database.

209 *Soil Extraction Data Organization for Analysis*

210 We combined FT-ICR-MS formula lists from all extractions of each soil sample to form a
211 single list of all compounds extracted from that sample. Due to variability in ionization
212 efficiency, we only considered compound presence, not intensity, in our analysis. We then
213 compiled compounds from all the rhizosphere soil samples to form one list of compounds
214 present in the rhizosphere, and we compiled compounds from all un-planted soil samples to
215 form a second list of compounds present in unplanted soil. By comparing these two lists and
216 removing any compounds which appeared on both, we isolated those compounds which were
217 unique to either the rhizosphere or unplanted soil. Using the assumption that the unplanted
218 soil was comparable to the initial condition of the soil before the influence of plants, we
219 considered the compounds unique to the rhizosphere to have been created and those
220 compounds unique to the unplanted soil to have been lost in the presence of roots.
221 Compounds which were common to both the rhizosphere and unplanted soil were classified as
222 persistent. For each sample we further identified compounds in the water-extracted fraction
223 and in a lumped category of hydrophobic compounds containing both methanol extracted, and
224 chloroform extracted compounds. Within the water-extracted or hydrophobic fraction of each
225 sample we identified the created, lost, and persistent compounds based on the master lists.

227 Compounds from both soil (FT-ICR-MS data) and roots (LAESI-MS data) were classified
228 according to their molecular formula using two different classification schemes. The first, a
229 frequently used atom-ratio classification scheme (Kim et al. 2003; Mopper et al. 2007; Minor et
230 al. 2012; D’Andrilli et al. 2013; D’Andrilli et al. 2015), involved sorting compounds using only O-
231 to-C and H-to-C ratios, producing classes with properties similar to lipids, lignins, proteins,
232 condensed hydrocarbons, amino-sugars, unsaturated hydrocarbons, carbohydrates, tannins,
233 and “other” undefined compounds. The second was a more recently described
234 multidimensional stoichiometric compound classification (MSCC) scheme that takes into
235 account molecular size and heteroatom (N, S, and P) composition (Rivas-Ubach et al. 2018). The
236 compound classes described in the MSCC are lipids, proteins, amino-sugars, carbohydrates,
237 nucleotides, phytochemicals, and “other” undefined compounds. Classification resulting from
238 both schemes does not indicate the compound is necessarily a lipid, lignin, protein, etc., but
239 rather that the compound has properties similar to that of a lipid, lignin, protein, etc.

240 We used two metrics to describe chemical properties: aromaticity index (AI) (Koch and
241 Dittmar 2006; Koch and Dittmar 2016) and nominal oxidation state of carbon (NOSC) (LaRowe
242 and Van Cappellen 2011). Both metrics are unitless. Calculations for these metrics are given in
243 Equations 1 & 2, with the elemental symbols representing the number of atoms of that element
244 present in the molecular formula. Aromaticity Index values calculated below zero are defined to
245 be zero, because negative aromaticity has no physical meaning (Koch and Dittmar 2006).

$$246 \quad 1) AI = \frac{1+C-0.5*O-S-0.5*(N+P+H)}{C-0.5*O-S-N-P}$$

$$247 \quad 2) NOSC = 4 - \frac{4*C+H-3*N-2*O+5*P-2*S}{C}$$

248 *Statistics – Compound Class and Metrics*

249 For the heteroatom and compound classifications, we analyzed the total count of each
250 classification within each soil sample (all three extractions combined). Box and whisker plots of
251 these data show the quartiles for total counts within samples. For metrics that calculate values
252 per compound (AI, NOSC, and atom ratios), we separated hydrophobic and water-extracted

253 compounds into separate lists and took the mean values for analysis. Therefore, each sample
254 produced six values: the mean of created, lost, and persistent compounds in both the water
255 extraction and hydrophobic extractions. For all statistical comparisons we used a mixed-effects
256 ANOVA with the box from which the sample came as a random effect variable, and the
257 treatment type (planted or unplanted) as the fixed effect, which we tested for significance. We
258 performed this test using the “fitlme” function in MATLAB (R2018b).

259 *Microbial Metabolisms and Transformation - FT-ICR-MS*

260 To assess involvement of identified compounds in microbial metabolic pathways in the
261 soil we compared assigned molecular formulae from the FT-ICR-MS data to the KEGG database
262 (release 86.1) to identify metabolic pathways involving the detected compounds (Kanehisa and
263 Goto 2000). We used the full list of KEGG compounds detected in our samples, not just created
264 or lost ones as with other analyses. This choice was made because we wanted to assess the full
265 pathway, and some compounds may be present on multiple pathways. Because metabolic
266 pathways can include members of various hydrophobicities, we combined all three extractions
267 for KEGG analysis.

268 We also identified biochemical transformations in the soil using the methodology of
269 Breitling et al. (2006) modified with an updated list of transformations from Graham et al.
270 (2017). This technique finds the difference in mass between all pairwise combinations of mass
271 peaks in the sample and compares them to a list of mass changes caused by known biochemical
272 transformations (Table S1). Because the transformation analysis relies on the relationships
273 between compounds/peaks we used all data and did not use the created/lost/persistent
274 categorization used for other analyses. If an observed mass difference was within 10^{-4} Da of a
275 transformation on the list then we counted it as a transformation. We analyzed the number of
276 transformations in each extraction normalized to the number of peaks and performed a
277 principal component analysis (PCA) on the percentage of total transformations that each
278 individual transformation accounted for within each extraction. We performed the PCA using
279 the “pca” function in MATLAB (R2018b). This procedure was repeated using only
280 transformations involving nitrogen atoms and transformations involving sulfur atoms. For the

281 nutrient-specific analysis of number of transformations, results are expressed as a percentage
282 of total transformations.

283 We also performed a variation of transformation analysis by comparing all peaks in
284 unplanted samples with all peaks in rhizosphere samples. This identified which transformations
285 converted compounds present in the unplanted soil into the compounds present in the
286 rhizosphere. The sign (positive or negative) of the mass difference was used to determine
287 whether the transformation represented the addition or removal of a group.

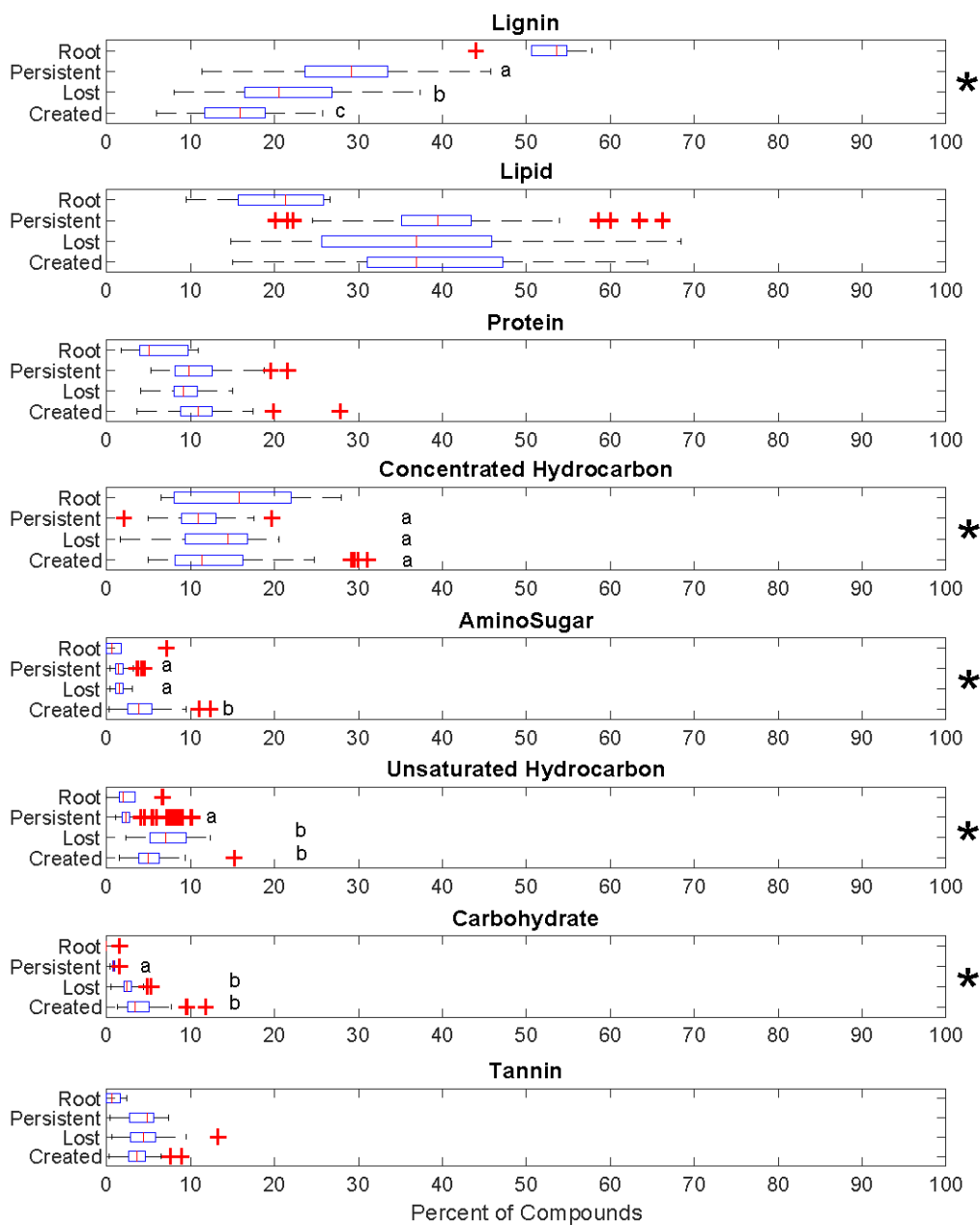
288 **Results**

289 Results describing the characteristics of persistent, lost, and created compounds are described
290 individually first and comparisons between these groups are made at the end of the results
291 section.

292 *Persistent compounds*

293 Using the traditional classification scheme, the carbon compounds which were present
294 in both unplanted and rhizosphere soil consisted largely of lipid-like and lignin-like compounds,
295 with these two categories comprising a mean of over 50% of compounds (Fig 1, uncategorized
296 compounds not shown). The MSCC only classified one quarter of persistent compounds as part
297 of a set category and placed three quarters of persistent compounds in an “other” category (SI
298 Fig S1). The MSCC is intentionally designed to be more precise about class assignments (Rivas-
299 Ubach et al. 2018), but the large number of unclassified compounds demonstrates how, in

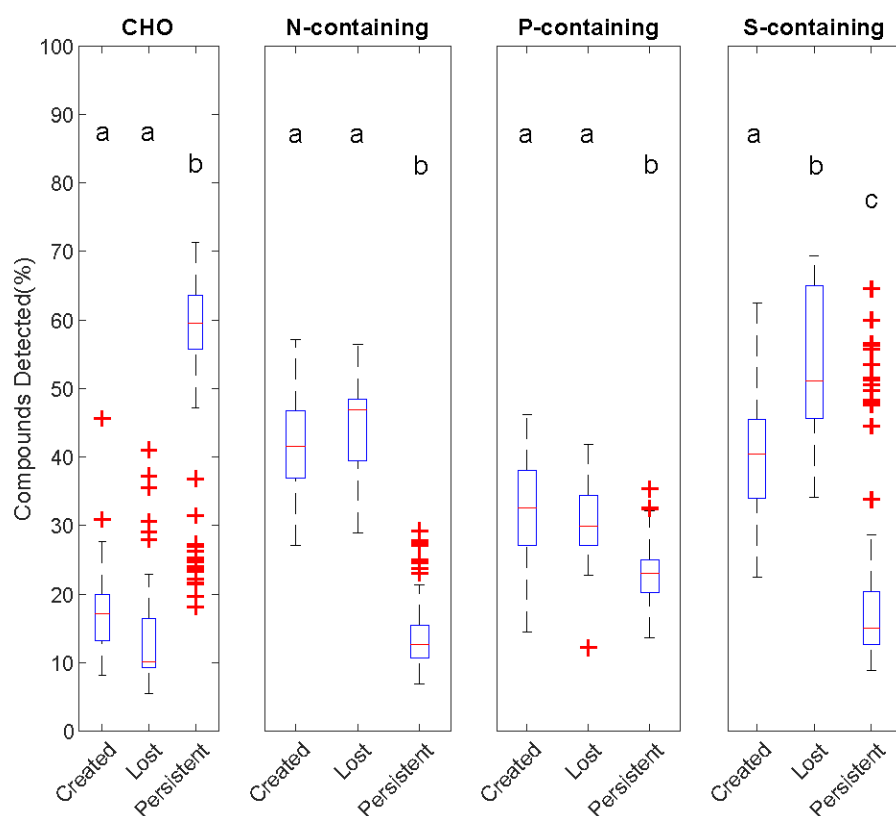
300 complex environmental media where compounds degrade and are reprocessed, many
 301 molecules fail to fall within the stringent categories.



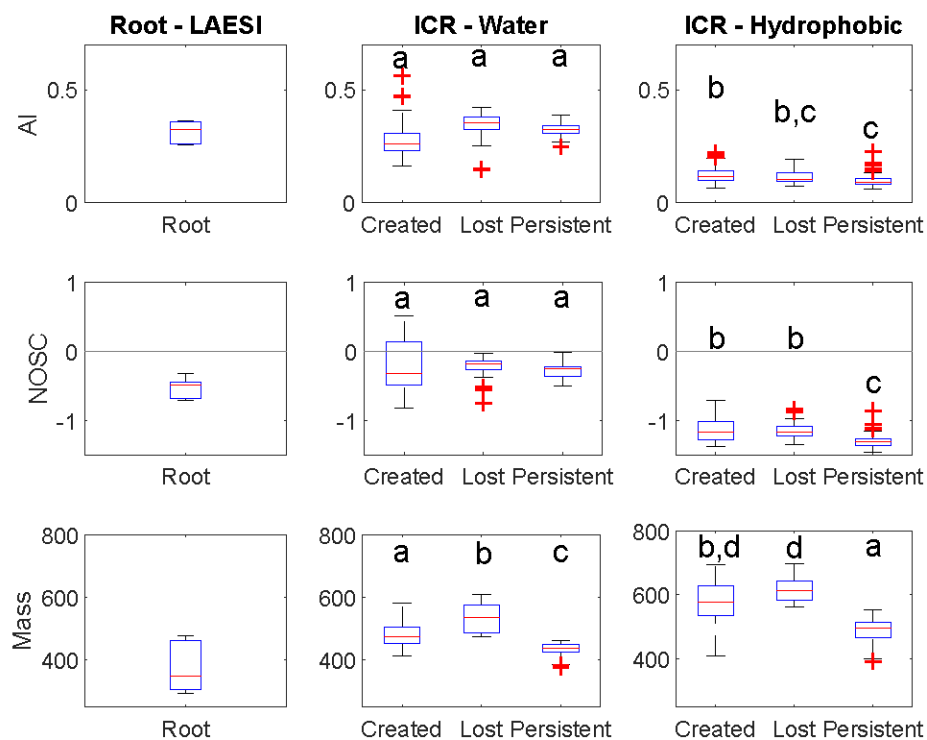
302
 303 Figure 1. Atom-ratio compound classifications for persistent, lost, and created soil compounds
 304 as well as root surfaces. Boxes show median with upper and lower quartiles, and tails show all data
 305 within 2.7σ of the mean. Outliers outside 2.7σ are marked with "+". Created n=73, lost n=40, persistent
 306 n=112, and roots n=6. The asterisks (*) indicate that soil compound category (persistent, lost, or

307 created) is a statistically significant ($p < 0.05$) effect in a mixed-effects model. The letters indicate which
308 categories are significantly ($p < 0.05$) different or not according to a Kruskal-Wallis ANOVA. Root
309 compounds were not included in the statistical analysis. Uncategorized compounds not shown.

310 The persistent compounds had few heteroatoms, and were mostly comprised of only
311 carbon, hydrogen, and oxygen (Fig 2). Graphs showing the ratio of each heteroatom to carbon
312 can be found in SI Figure S2. Persistent compounds had significantly higher NOSC and AI in the
313 water-soluble fraction than the hydrophobic fraction (Fig 3). The median mass was 436 Da for
314 water-extracted persistent compounds and 615 Da for hydrophobic persistent compounds.



315
316 Figure 2, number of compounds in each sample that fall into different heteroatom groups for persistent,
317 lost, and created soil compounds. All heteroatom groups had significant ($p < 0.05$) differences;
318 differences within each type are marked with letters. Distributions are of the means of individual
319 samples. Boxes show median with upper and lower quartiles, and tails show all data within 2.7σ of the
320 mean. Outliers outside 2.7σ are marked with "+". For created n=73, for lost n=39, and for persistent
321 n=112.



322

323 Figure 3. Boxplots of molecular indices. The rows show: aromaticity index (AI), nominal oxidation state
 324 of carbon (NOSC), and the molecular mass in Daltons. Letters indicate groups which are significantly
 325 different ($p < 0.05$) from other groups for that index. The analysis spans the water and hydrophobic
 326 classification, and does not include root compounds.

327

328 *Compounds processed in the rhizosphere — Created and Lost*

329 Compounds affected by microbial processing, those that were either created or lost,
 330 were classified mostly as lipid-like and lignin-like compounds, similar to persistent compounds,
 331 with these two categories comprising a mean of over 50% of compounds using the traditional
 332 scheme (Fig 1), and were largely unclassified using MSCC (SI Fig S1). Unlike persistent
 333 compounds, a majority of both created and lost compounds contained heteroatoms, with less
 334 than 20% CHO compounds (Fig 2). Bogs are typically nutrient limited (Aerts et al. 1992), and it is

335 important to note that even though these heteroatoms existed they were not necessarily
336 available to plants or microbes, as will be addressed in the discussion below.

337 Similar to persistent compounds, processed compounds had significantly higher NOSC
338 and AI in the water-soluble fraction than the hydrophobic fraction (Fig 3). The median masses
339 were between 475 Da and 615 Da. The masses varied both between water-extracted and
340 hydrophobic fractions and between created and lost compounds (Fig 3), as will be discussed in
341 the next sections.

342 *Comparison between hydrophobic and water-extracted compounds*

343 The hydrophobic compounds had lower AI and NOSC than water-extracted compounds
344 across lost, created, and persistent compounds (Fig 3). For mass, each category in the water
345 extraction was smaller than the corresponding category among hydrophobic compounds, e.g.,
346 lost compounds in the water extraction were smaller than lost compounds in hydrophobic,
347 even though the lost water-extracted compounds were still larger than the persistent
348 hydrophobic compounds.

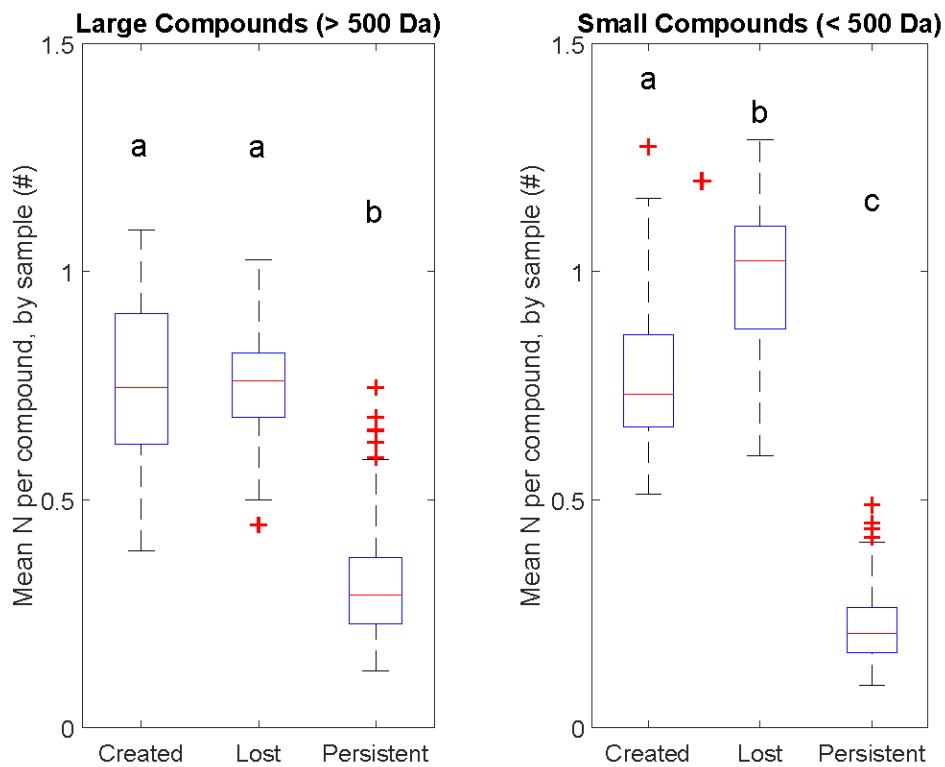
349 *Comparison between processed and preserved compounds.*

350 The compound classes were broadly similar between created, lost, and preserved
351 compounds, but there were small yet statistically significant ($p < 0.05$) differences in lignins,
352 amino-sugars, and carbohydrates (Fig 1). Preserved compounds had the highest percentage
353 lignins, followed by lost then created. Created compounds were comprised of more amino
354 sugars than either persistent or lost compounds. Created compounds were also most likely to
355 be comprised of carbohydrates, followed by lost compounds, and persistent compounds were
356 the least likely to be carbohydrates.

357 Heteroatom groupings revealed significant ($p < 0.05$) differences between the persistent
358 and microbially processed molecules. The fraction of nitrogen- and sulfur-containing
359 compounds was higher in both lost and created compounds than in persistent, with a smaller
360 but still significant difference in phosphorus (Fig 2). Compounds comprised of only CHO had the
361 opposite pattern; they were most abundant in persistent compounds.

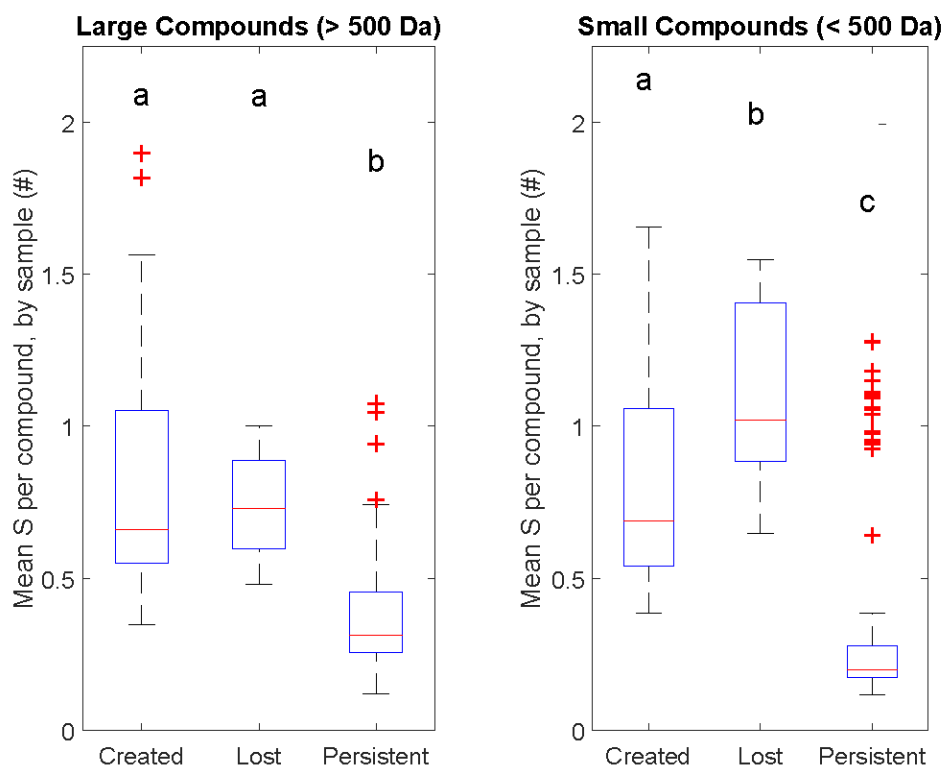
362 Of the three metrics shown in Figure 3, mass had the most salient differences across
363 both processing type and hydrophobicity. In both hydrophobic and hydrophilic fractions, the
364 lost compounds had the highest mass followed by created compounds with persistent
365 compounds having the lowest mass. The NOSC data showed no statistically significant
366 difference between preserved, created, and lost compounds in the water extracted fraction,
367 but in the hydrophobic fraction the created compounds had higher ($p < 0.05$) NOSC values than
368 persistent compounds, with lost compounds having an intermediate value not significantly
369 different ($p > 0.05$) from either (Fig 3). There were statistically significant differences in AI; lost
370 compounds in the water extraction had the highest AI, followed by persistent then created
371 compounds. In the hydrophobic extraction the pattern was different: created compounds had
372 the highest AI followed by lost then persistent compounds.

373 Because molecular mass was different between the groups of molecules, and microbes
374 break larger compounds down into smaller compounds (Amon and Benner 1996; McArthur and
375 Richardson 2002; Antony et al. 2012; Pracht et al. 2018), it is useful to analyze high and low
376 molecular mass compounds separately. Among low molecular mass compounds (<500 Da), in
377 the water extraction, lost compounds had the highest average NOSC and AI, and the persistent
378 compounds had the lowest values for both metrics, with the differences being statistically
379 significant ($p < 0.05$) for both metrics (Fig S4). Among high mass (>500 Da) compounds there
380 were no statistically significant differences in AI or NOSC in either water-extracted or
381 hydrophobic compounds (Fig S5). Both high and low molecular mass compounds had largely
382 similar classifications to the full data set (Figs S6 & S7). However, small created compounds had
383 significantly less nitrogen and less sulfur than small lost compounds ($p < 0.05$, Fig 4 and Fig 5), a
384 difference that did not exist among the large compounds ($p > 0.05$, Figs 4 and 5). This pattern
385 was true among both hydrophobic and hydrophilic compounds (Fig S8). When examining
386 compound classes by molecular size, the small created molecules had significantly ($p < 0.05$)
387 more proteins, amino sugars, and carbohydrates than lost compounds (Fig S7).



388

389 Figure 4. Nitrogen to carbon ratios among both large (left) and small (right) compounds. Small
 390 compounds created in the rhizosphere were depleted in nitrogen relative to those lost. No difference in
 391 nitrogen content was detected between created and lost large compounds.



392
 393 Figure 5. Sulfur to carbon ratios among both large (left) and small (right) compounds. Small compounds
 394 created in the rhizosphere were depleted in sulfur relative to those lost. No difference in sulfur content
 395 was detected between created and lost large compounds.

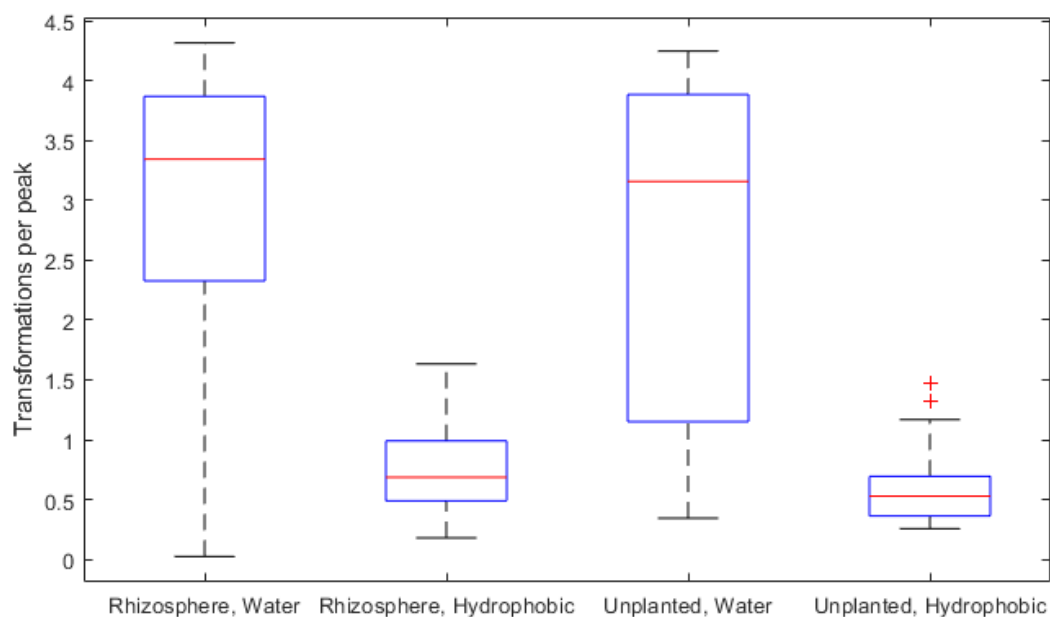
396
 397 *Root compounds*

398 The compounds detected on the root surface by LAESI-MS are consistent with
 399 fragments of structural lignins comprising the root itself. The traditional compound
 400 classification scheme identified over half the compounds as lignin-like, with lipids and
 401 concentrated hydrocarbons comprising most of the rest (Fig 1). As with the soil compounds,
 402 most root compounds could not be classified with the MSCC (Fig S1). Around three quarters of
 403 the detected compounds contained only C, H, and O (Fig S2), which is typical of lignins or other
 404 structural molecules such as polysaccharides. The AI was high, which is again consistent with
 405 structural lignins and polysaccharides that contain many aromatic rings.

406 Note that these measurements were made of the root material, not the root exudates
407 which are typically comprised of amino acids and sugars (Smith 1976; Lugtenberg et al. 1999;
408 Jones et al. 2009; Dommergues 2012) and are believed to stimulate microbial activity. However,
409 dead root cells sloughed into the soil may have been introduced into the rhizosphere.

410 *Microbial Transformations*

411 Microbial transformation analysis involved all compounds present in a sample, not just those
412 identified as being persistent, created, or lost in the presence of roots. There were significantly
413 more ($p < 0.05$) microbial transformations among the water extracted compounds than in the
414 hydrophobic compounds, but there were no differences ($p > 0.05$) between the number of
415 transformations in rhizosphere or unplanted extractions (Fig 6). These results indicate that the
416 compounds in the water fraction were more involved in microbial metabolisms than those in
417 the hydrophobic fractions. Similarly, the PCA analysis of the transformation data showed that
418 the first two principle components together explained over 95% of variation and showed that
419 the water extraction was offset from hydrophobic compounds, but there was no difference
420 between rhizosphere and unplanted soil samples (Fig S8).



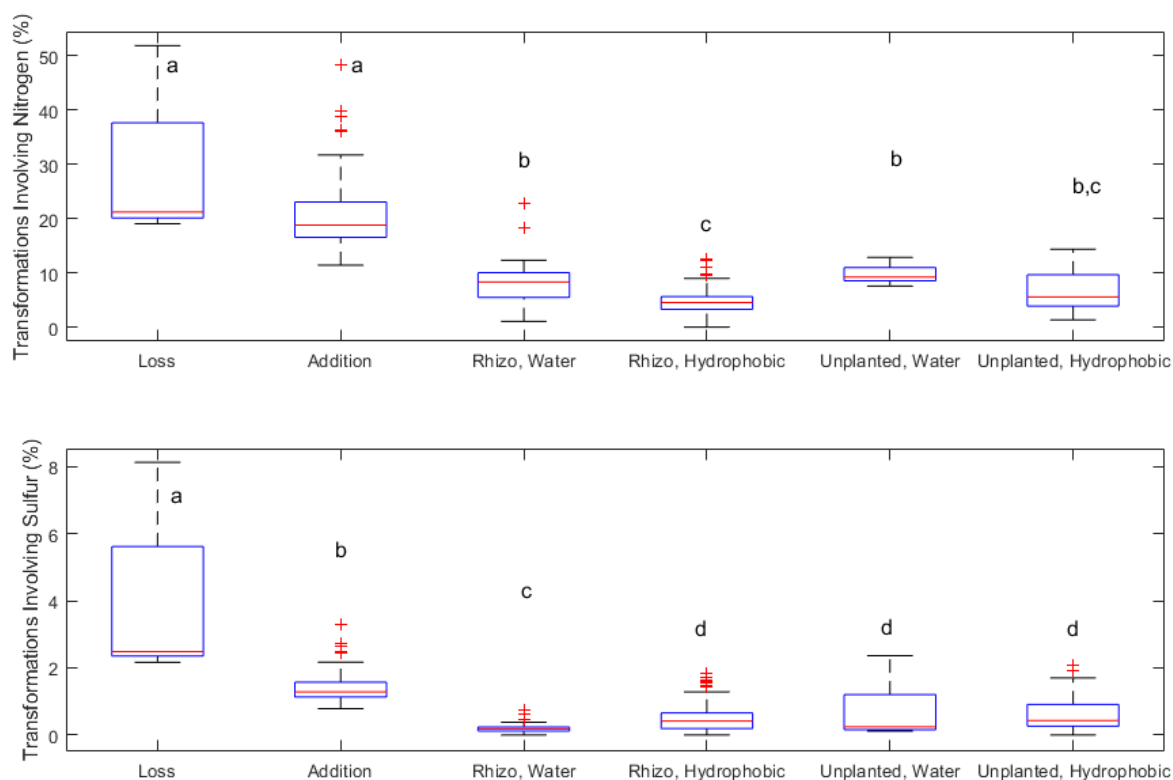
421
422 Figure 6. Comparison of number of microbial transformations in each extraction from both sample types
423 (rhizosphere and unplanted), including all compounds in each sample, not just lost or created. These

424 data show that more microbial transformations occurred with water-soluble compounds than with
425 hydrophobic compounds. However, there is no difference in transformations between rhizosphere and
426 unplanted soil.

427

428 Breaking the transformations down by nutrient showed that reactions involving
429 nitrogen were similarly likely in rhizosphere and unplanted soil (Figure 7). However,
430 transformations converting unplanted soil compounds to compounds found in the rhizosphere
431 were more likely to involve nitrogen than the transformation within an extraction, whether the
432 nitrogen was added during the transformation or removed (Figure 7).

433 Transformations involving sulfur were overall less common than those involving
434 nitrogen, always accounting for less than 10% of all transformations. Unlike nitrogen, the PCA
435 did reveal differences between the types of sulfur transformations happening within each
436 extraction (Fig S10). Statistically significant ($p < 0.05$) differences in PC1 (explaining 49% of
437 variation) and PC2 (explaining 16% of variation) point towards the extractions having unique
438 transformation profiles. The strongest correlation for any one transformation is in PC1 where a
439 sulfate for water substitution was four and a half times more predictive than any single other
440 reaction. When transforming from unplanted soil to rhizosphere (The Removal and Addition
441 categories in Figure 7), sulfur was significantly more likely to be removed than added ($p < 0.05$).
442 Also, more removal and addition transformations involved sulfur than the transformations
443 within any extraction ($p < 0.05$, Figure 7).



444

445 Figure 7. Comparison of the portion of transformations that involved nitrogen (top) and
 446 sulfur (bottom) among transformations within and between carbon pools. The “Removal”
 447 category comprises transformations from unplanted soil to rhizosphere where a functional
 448 group is removed, the “Addition” category comprises transformations where a group is added.
 449 For both nutrients, the transformation type that was most likely to involve the nutrient was the
 450 loss of compounds from the unplanted soil.

451 KEGG pathway analysis revealed no significant difference ($p > 0.05$) in the number of
 452 metabolites or metabolic pathways between created and lost compounds. A PCA analysis of the
 453 pathways did reveal a statistically significant ($p < 0.05$) difference between created and lost
 454 compounds, but the difference was small. The first principle component explained 99.9 % of
 455 the variation, and its interquartile range among created compounds was .050 to .055, while
 456 among lost compounds the interquartile range was .052 to .060. There was no correlation
 457 between PC 1 and the flux of CH_4 out of the box ($p > 0.05$, $R^2 = .03$).

458

459 **Discussion**

460 *A note on FT-ICR-MS data*

461 Data obtained by FT-ICR-MS describe those compounds that were present in a sample
462 during collection, not all those which ever existed. This distinction is important because those
463 compounds which are most bioavailable and therefore used very rapidly may never show up in
464 the results. In the context of this study, some root exudates fall into this category. Many types
465 of root exudates may be added to the soil, but if some of those are processed quickly then only
466 the less-bioavailable fraction will remain to be measured. Additionally, because the method
467 only indicates whether compounds are present or not, there is no way to tell if a compound
468 was exuded by a root or created by microbes in the soil.

469 *Differences between Created and Lost Compounds*

470 Compounds lost in the presence of roots must have been processed in a way that does
471 not occur in the absence of roots. Compounds created in the presence of roots could have
472 come from either root-influenced processes (e.g., metabolic by-products of microbes that
473 utilize root exudates) or been exuded from the roots themselves. However, direct detection of
474 root exudates is unlikely for several reasons. Some of the sugars and amino acids that comprise
475 root exudates (Smith 1976; Lugtenberg et al. 1999; Jones et al. 2009; Dommergues 2012) were
476 too small to be detected consistently using the FT-ICR-MS technique (<100 Da). Additionally,
477 sugars and amino acids are water-soluble and typically have high NOSC values, making them
478 highly bioavailable in anoxic systems (Keiluweit et al. 2016). Because the created compounds
479 were similar to lost and preserved compounds in terms of NOSC and proportion of protein- and
480 amino-sugar-like compounds (Figs 1 and 3), we conclude that the exudates themselves were
481 not directly detected. Working from the interpretation that the compositional changes we
482 observed in the rhizosphere were *not* due to the direct detection of root exudates, we can draw
483 conclusions about what kind of microbial processing could explain the observed changes in the
484 rhizosphere.

485 The experiment from which the analyzed samples were obtained demonstrated that,
486 based on greater CH₄ emissions from planted than unplanted boxes, the rhizosphere housed a

487 more active microbial population (Waldo et al. 2019). However, neither the KEGG database nor
488 microbial transformation analysis (Fig 6) techniques detected a difference between rhizosphere
489 and bulk soil samples, as those techniques have in other studies (Graham et al. 2017). This lack
490 of difference could indicate that these metabolite-based methods are not very sensitive to
491 differences in the population or activity of microbes in two systems that are similar (e.g., the
492 same highly organic soil material), or that, in the rhizosphere, key metabolites were cycled so
493 quickly that they did not exist long enough to be detected. This rapid cycling would impact the
494 other metrics used to understand soil carbon by biasing the measurements against the most
495 biologically active molecules.

496 For both hydrophobic and hydrophilic compounds, the lost compounds were higher MW
497 than the created compounds (Fig 2), leading to a net decrease in average compound size across
498 the rhizosphere. This decrease in mass is consistent with the breakdown of higher MW
499 molecules into smaller ones, which aligns general understanding of microbial processing. The
500 fact that plants were facilitating the mass decrease aligns with the findings of Chanton et al.
501 (2008) who found smaller compounds in *Carex* dominated bogs than in those where vascular
502 plants were absent. In this experiment, the group with the smallest mean mass was the
503 preserved compounds, indicating that microbial and/or abiotic transformations did not target
504 these small compounds and that they were also not taken up by plants.

505 Surprisingly, neither AI nor NOSC was able to explain why some compounds were used
506 and others were not, even though those metrics have been useful in other studies (e.g.,
507 Keiluweit et al. 2016; Boye et al. 2017; Pracht et al. 2018). Transformation analysis (Fig 6)
508 indicates that the water-extracted compounds were more actively used in microbial
509 metabolisms, but there were no significant differences among created, lost, and preserved
510 compounds in the water-extracted AI or NOSC. There were, however, differences in AI and
511 NOSC between water-extracted and hydrophobic compounds, indicating that these indices
512 were still differentiating between molecules with important properties such as hydrophobicity.
513 One possible explanation for the apparent lack of influence of NOSC is that while there was no
514 detectable oxygen in the soil (Waldo et al. 2019), the flux of oxygen from the roots could have
515 nonetheless generated enough electron acceptors that NOSC was less of a salient factor than in

516 the highly reduced conditions studied by Keiluweit et al. (2016) and others (e.g., Boye et al.
517 2017; Pracht et al. 2018). Instead, molecular mass and nutrient content had more apparent
518 control over microbial processing (Figs. 3 and 4), an observation that is best explained by
519 microbial priming.

520 *Microbial Priming*

521 We demonstrated in our previous analysis of isotopic data from this plant-growth
522 experiment that microbial priming was occurring (Waldo et al. 2019). Priming is broadly defined
523 as an increase in microbial use of soil carbon due to the addition of a fresh carbon source. In
524 this analysis, changes in molecular characteristics such as size are used as proxies of microbial
525 activity. Most notably, higher MW compounds were broken down into smaller compounds in
526 the presence of roots. The roots could have simply accelerated a microbial process which also
527 occurred in the unplanted soil, a situation that is consistent with the hypothesis that priming is
528 caused by increased populations of microbes needing more food (Ye et al. 2015).

529 However, the microbes could have also engaged in N-mining. The N-mining hypothesis
530 predicts that microbes break down molecules that would otherwise be less bioavailable in
531 order to obtain nitrogen (Craine et al. 2007; Ruirui et al. 2014). N-mining could explain why the
532 low molecular mass compounds that were created had significantly less ($p < 0.05$) nitrogen than
533 those that were lost (Fig 4). There was a net removal of nitrogen from small compounds in the
534 presence of roots. Plants likely encouraged this N-mining in two ways: they added nitrogen-
535 poor root exudates which provided microbes with extra energy but not extra nutrients, and
536 they also would have absorbed some of the small nitrogen-containing compounds, decreasing
537 the pool of available nitrogen in the soil. Among large compounds, the created and lost
538 compounds had similar nitrogen content (Fig 4), indicating that these compounds were not
539 being mined for nitrogen. Rather, large compounds were processed for energy regardless of
540 nitrogen content. The molecules that remained and accumulated in the rhizosphere were small,
541 nitrogen-poor compounds that did not meet the needs of the plants or microbes.

542 Sulfur-containing compounds also followed a similar pattern to nitrogen-containing
543 ones (Fig 5), indicating that S-mining may be occurring as well as N-mining. In fact, the

544 transformation analysis provided additional evidence for S-mining; significantly more sulfur-
545 containing groups were removed than added, a difference that appeared to exist for nitrogen
546 transformations, but was not statistically significant (Figure 7). In addition, the transformation
547 analysis found differences in the types of transformations between rhizosphere and unplanted
548 soil in reactions involving sulfur. The single transformation which was most increased in the
549 rhizosphere was a sulfate for water substitution. The fact that sulfate is the primary form of
550 sulfur taken up by plants (Leustek and Saito 1999) indicates that plant uptake may be ultimately
551 driving the sulfate use in the rhizosphere. This observation is consistent with previous studies
552 that have shown that addition of root exudates or other simple substrates can increase
553 microbial sulfur utilization in a variety of soils (Vong et al. 2003; Vong et al. 2003; Creamer et al.
554 2014). Studies of the interaction between soil carbon and nutrient availability have focused on
555 nitrogen because it is both a key nutrient and one with high anthropogenic influence (Creamer
556 et al. 2014). However, the concept can be applied to any limiting nutrient in the soil.

557 *Priming of water soluble versus physically bound carbon*

558 The FT-ICR-MS data analyzed here advances understanding of the mechanisms of
559 priming by demonstrating that the presence of plants triggered increased processing of both
560 large molecules (regardless of nutrient content, Fig 3) and small nutrient-containing molecules
561 (Figs 4&5). Interestingly, none of the evidence for mechanisms of priming varied between
562 hydrophobic and hydrophilic compounds. Both rhizosphere and unplanted soil samples had
563 more microbial transformations and more transformations involving nitrogen in the water-
564 soluble compounds (Fig 6). For all other evidence of microbial activity (e.g., created compounds
565 of smaller size than lost compounds), the same patterns were exhibited between hydrophobic
566 and water-soluble compounds (Fig 3 and Fig S8). This finding is notable because Graham et al.
567 (2017) found that the addition of water-soluble, thermodynamically favorable carbon
568 decreased microbial processing of carbon compounds in the chloroform extraction, referred to
569 as the “physically bound” fraction in that study.

570 While such preservation of the physical bound, hydrophobic fraction did not occur in
571 our system, the nature of our evidence for priming indicates that the results of Graham et al.

572 (2017) do not necessarily contradict priming, but instead elucidate how it occurs. We propose
573 that priming is a selective effect which often increases total carbon use, but which may increase
574 *or decrease* the microbial use of specific carbon pools such as water-soluble versus physically
575 bound or compounds with specific heteroatoms. This selectivity could be the result of several
576 factors. At its most basic level, it could be that priming exacerbates the already-present
577 preferences, such as for soluble compounds (Boye et al. 2017; Graham et al. 2017).
578 Alternatively, selective priming could be the result of favoring the growth of one community of
579 microbes over other communities, or the result of encouraging microbes to change their
580 metabolisms and produce a different set of enzymes. The nutrient-mining hypothesis, which
581 this study strongly supports, is premised upon the idea that microbes can use different pools of
582 carbon for different purposes, i.e., some compounds are utilized for energy and others for
583 nutrients. That effect is demonstrated here by the fact that created and lost large compounds
584 had similar nitrogen content, while smaller compounds had significant amounts of nitrogen
585 removed when processed (Fig 4). This study also found evidence of a similar effect for sulfur,
586 another plant nutrient. At least some portion of the nutrients that disappeared from the ICR
587 data were taken up by plant roots, and as plants removed nutrients from the soil system
588 microbes needed to process more nutrient-containing compounds to ensure their own supply.
589 If root exudates are more favorable for microbes to use for a certain purpose then native
590 compounds in the soil which microbes used for that purpose may be protected, a possible
591 explanation for the results of Graham et al. (2017).

592 **Conclusions**

593 This study built upon prior analysis of Waldo et al (2019), which established that
594 microbial priming occurred in the presence of plants and led to increased CH₄ production and
595 emissions. The two metrics that were most useful in differentiating between compounds
596 processed and preserved in the presence of plants were size and nitrogen content in water-
597 soluble compounds. It is well established that microbes break higher MW molecules down into
598 smaller molecules during processing of organic carbon (Amon and Benner 1996; McArthur and
599 Richardson 2002; Antony et al. 2012; Pracht et al. 2018), and here we demonstrated that the
600 presence of plants led to increased processing of those large molecules, presumably due to the

601 addition of root exudates. This increase in processing of large molecules is an example of
602 priming caused by a more active microbial community utilizing more total soil carbon. Small
603 molecules were also processed to a greater extent in the presence of plants, but only if they
604 contained nutrients, lending support to the nutrient-mining hypothesis (Craine et al. 2007;
605 Ruirui et al. 2014; Creamer et al. 2014) and possibly expanding it to indicate sulfur-mining as
606 well. This selective priming implies a shift in the carbon processing capabilities and/or
607 nutritional needs of the microbial community in the presence of highly preferable compounds.
608 Because boreal bogs tend to be nutrient-limited environments, this laboratory study implies
609 that such nutrient-mining may be favorable in similar systems around the world and the strong
610 priming effect demonstrated in the companion study (Waldo et al. 2019) could occur in natural
611 systems. Changes to the global climate and nutrient cycles are connected (Greaver et al. 2016),
612 and so determining how they interact in this key ecosystem will help us understand future
613 changes to soil carbon storage in northern peatlands, which represents an important climate
614 feedback (Yu 2012).

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